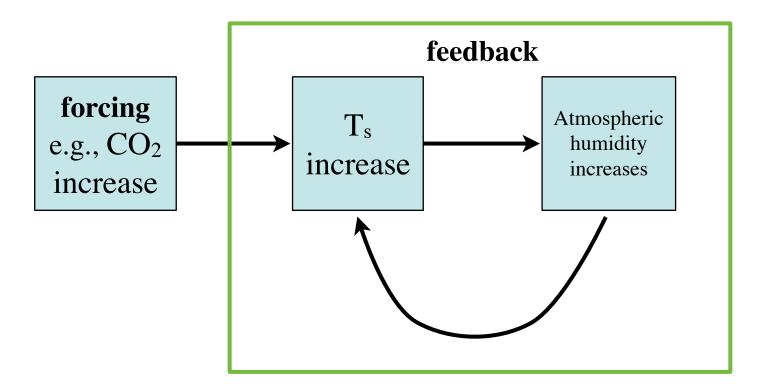
Verifying the water vapor feedback using ENSO

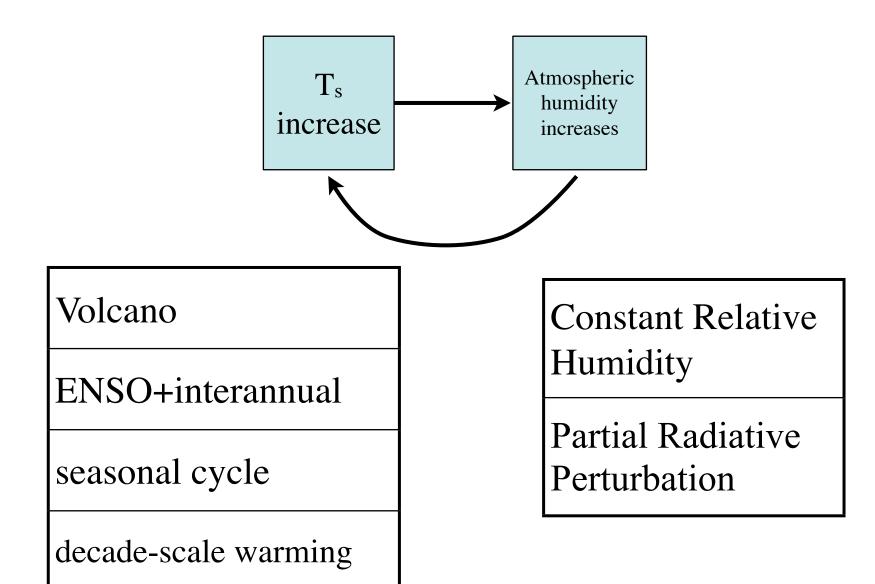
A. E. Dessler and S. Wong Department of Atmospheric Sciences Texas A&M University



How can we measure this?







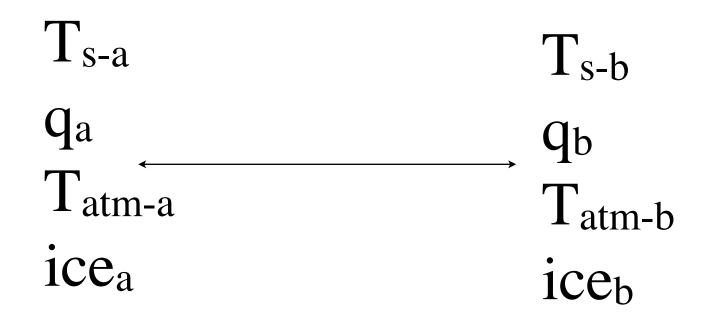


Observational tests of the water vapor feedback

Volcano	Soden et al., 2002; Forster and Collins, 2004
ENSO+interannual	Soden, 1997; Dessler et al., 2008; Minschwaner and Dessler, 2004; Gettelman and Fu, 2008;
seasonal cycle	Inamdar and Ramanathan, 1998; Wu et al., 2008
decade-scale warming	Hall and Manabe, 1999; Soden et al., 2005

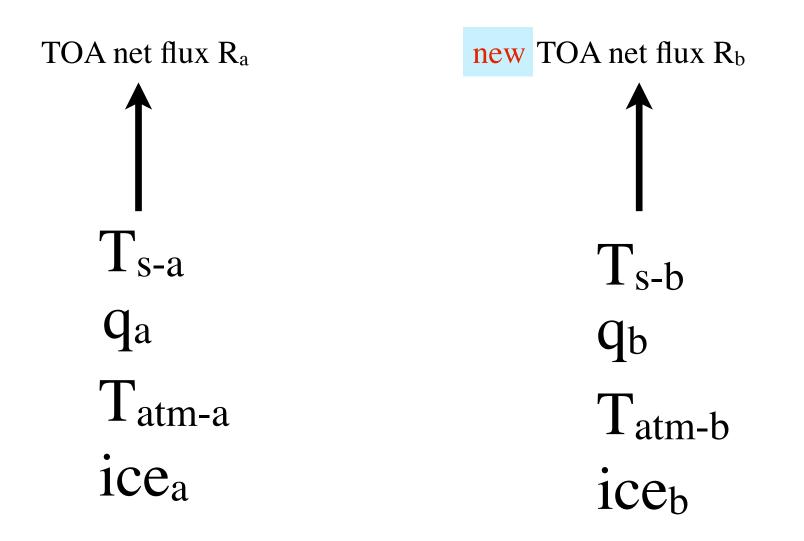


Partial Radiative Perturbation



Wetherald and Manabe, 1988 Colman, 2003; Forster and Collins, 2004; Soden and Held, 2006

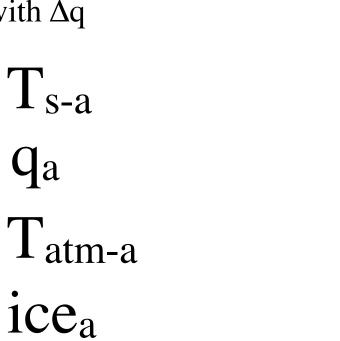


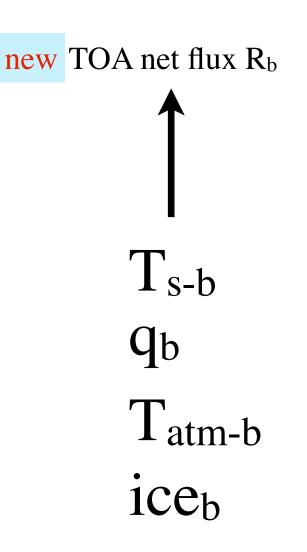


Wetherald and Manabe, 1988? Colman, 2003; Forster and Collins, 2004; Soden and Held, 2006



 ΔR = change in global average TOA flux due to Δq ΔT_s = change in global avg. surface temperature change (T_{s-b}-T_{s-a}) associated with Δq





Wetherald and Manabe, 1988? Colman, 2003; Forster and Collins, 2004; Soden and Held, 2006



the water vapor feedback

$$\lambda = \frac{\Delta R}{\Delta T_S} \qquad T_s = \frac{-G}{\lambda_o + \lambda_1 + \lambda_2 + \cdots}$$

 ΔR = change in global average TOA flux due to Δq between two climate states

 ΔT_s = change in global avg. surface temperature change (T_{s-b}-T_{s-a}) associated with Δq

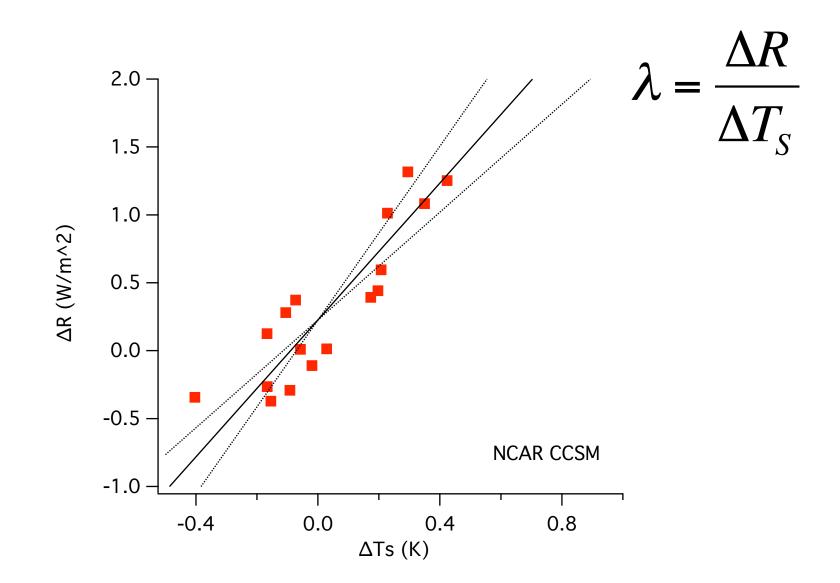
Wetherald and Manabe, 1988? Colman, 2003; Forster and Collins, 2004; Soden and Held, 2006



Method

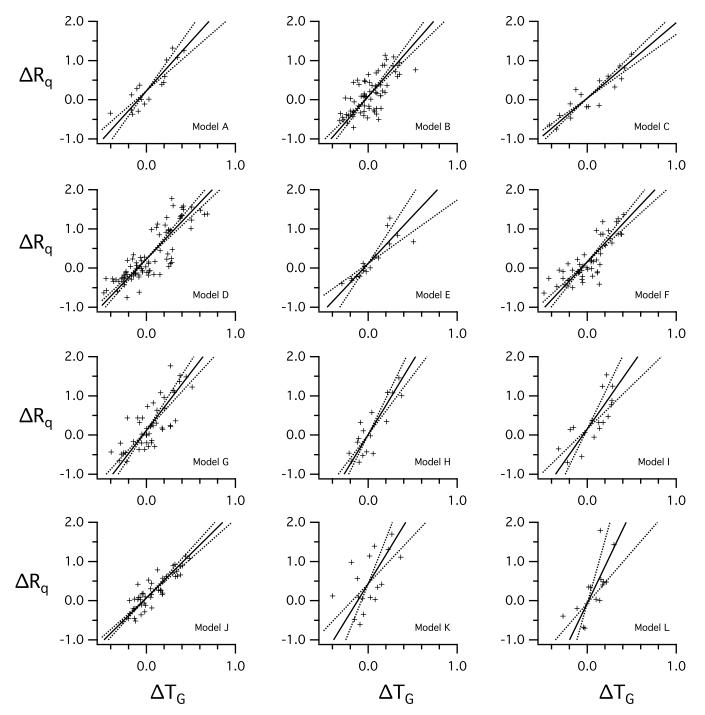
- AMIP models from the PCMDI archive
- Reanalysis: ERA40 & MERRA
- For each strong DJF ENSO month (e.g., Jan. 1983), calculate ΔR and ΔT_s between that month and the longterm average for all Januaries in the model run
- Obtain one estimate of ΔR and ΔT_s for every strong ENSO month



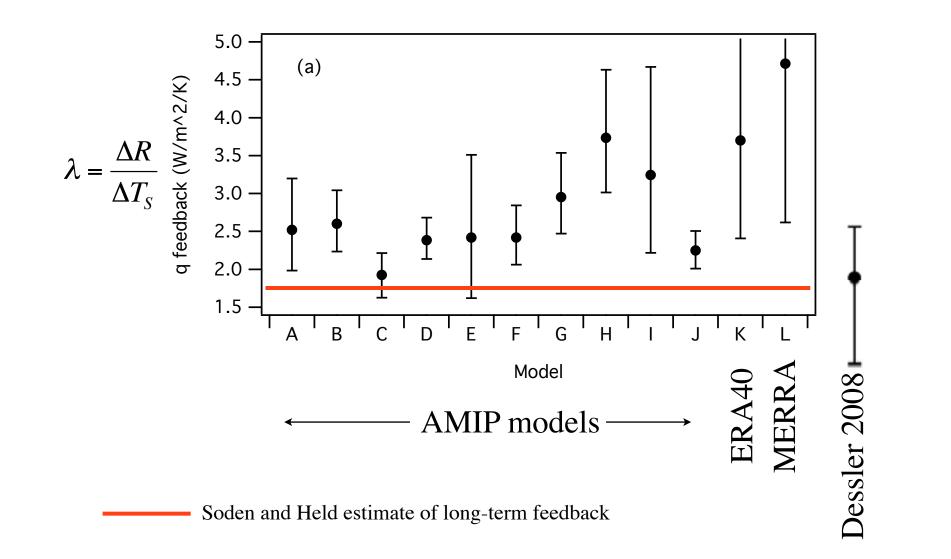


Each point is a strong ENSO DJF









Conclusions 1

- Models, reanalysis, and pure obs. agree that the feedback is positive and strong
- Reanalysis is larger than the models
 - if reanalysis is correct, then models are underestimating the feedback
 - uncertainties generally overlap
- ENSO WV feedback larger than in response to long-term warming (e.g., Colman, Soden and Held)

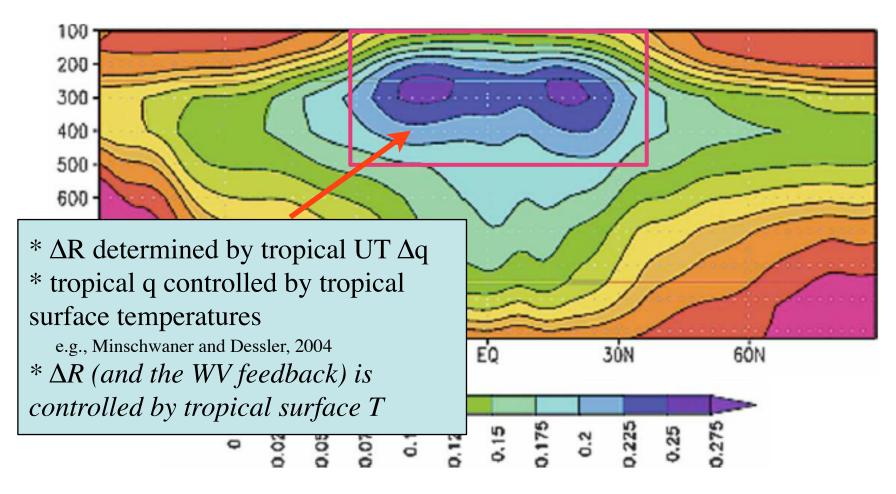


Questions for the rest of the talk

- What is the source of disagreements among the models and between the models and reanalyses?
- Speculation: Why is the ENSO feedback larger than the feedback in response to long-term warming?



Water vapor feedback is primarily a "tropical" phenomenon

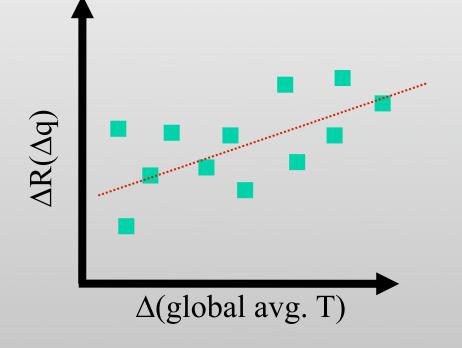


Change in R per unit change in q(x,y,z): $\Delta R/\Delta q(x,y,z)$

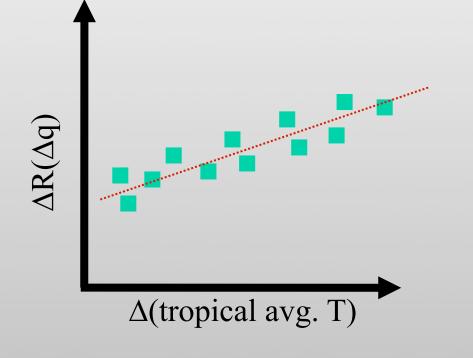
Fig. 2 of Soden et al., 2008

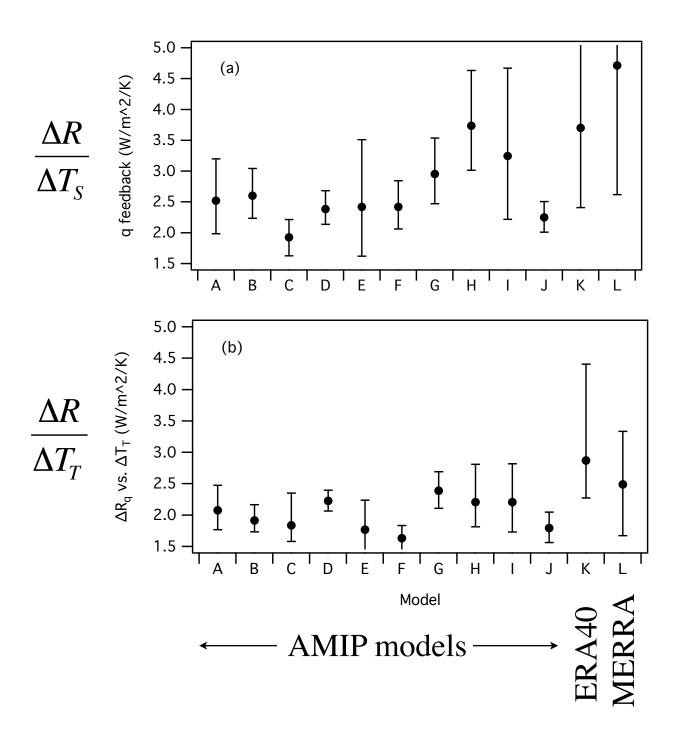


Regress ΔR vs. *global* surface temperature







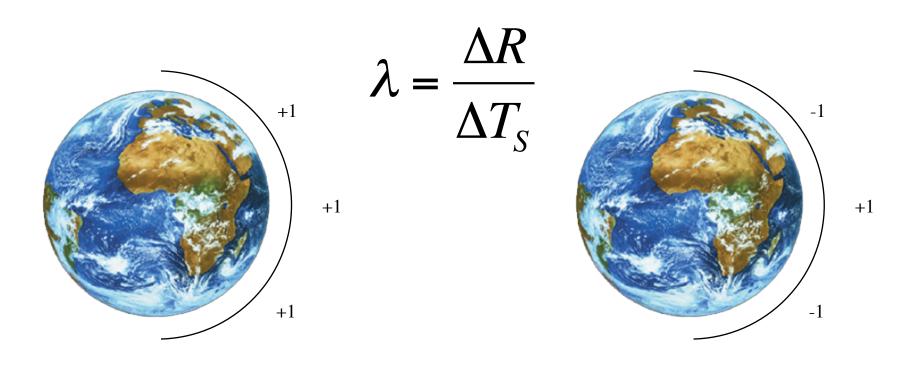




What does this mean?

- A consistent relationship exists between tropical surface ΔT_T and the radiative response to water vapor ΔR
 - GCMs
 - reanalyses
 - pure obs.





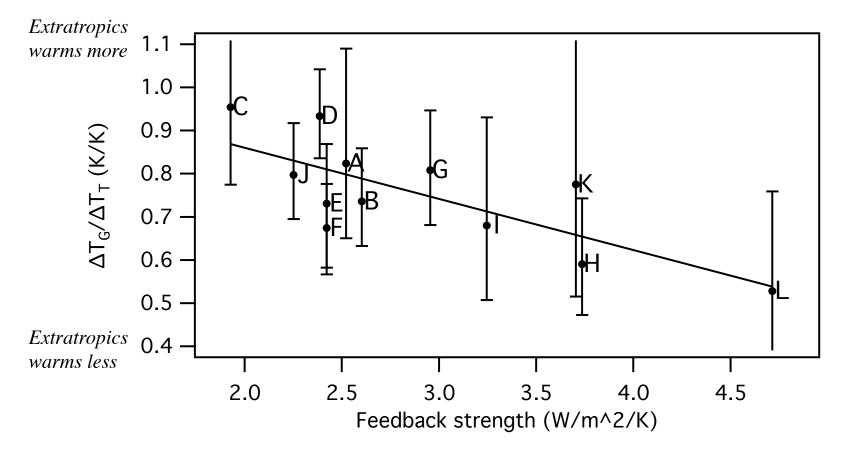
smaller feedback

larger feedback

ΔR for these two worlds is the same ΔT_s is different



$$\lambda = \frac{\Delta R \left(\Delta q \left(\Delta T_T \right) \right)}{\Delta T_G}$$





Summary

- In response to ENSO climate change, models, reanalysis, and pure obs. show a strong and positive water vapor feedback
- Models, reanalysis, and pure obs. show that the radiative response to WV between two climate states is determined by change in the tropical surface T
 - WV feedback determined by tropical UT water
 - Tropical UT water is controlled by surface T described by Minschwaner and Dessler [2004] (see me for reprints)
- Differences among models and between models and reanalyses are due to differing estimates of extratropical surface T changes --- they are NOT due to the radiative response of WV
- This work is in press at J. Climate. See me for preprints.