

Water vapor and the transition to strong convection

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- **Issues with precip. simulation, esp. at regional scales, tropics: global warming, El Niño..., Sensitivity to convective schemes**
e.g., IPCC 2001, 2007; Trenberth et al 2003; Maloney and Hartmann 2001; Joseph and Nigam 2006; Biasutti et al. 2006; Dai 2006; Tost et al. 2006; Neelin et al 2007; Bretherton 2007...

1. Sensitivity of convective margin zones

2. Characterizing transition to deep convection

- **dependence on temperature and water vapor**
- **remote sensing statistics and buoyancy calculations from vertical structure**

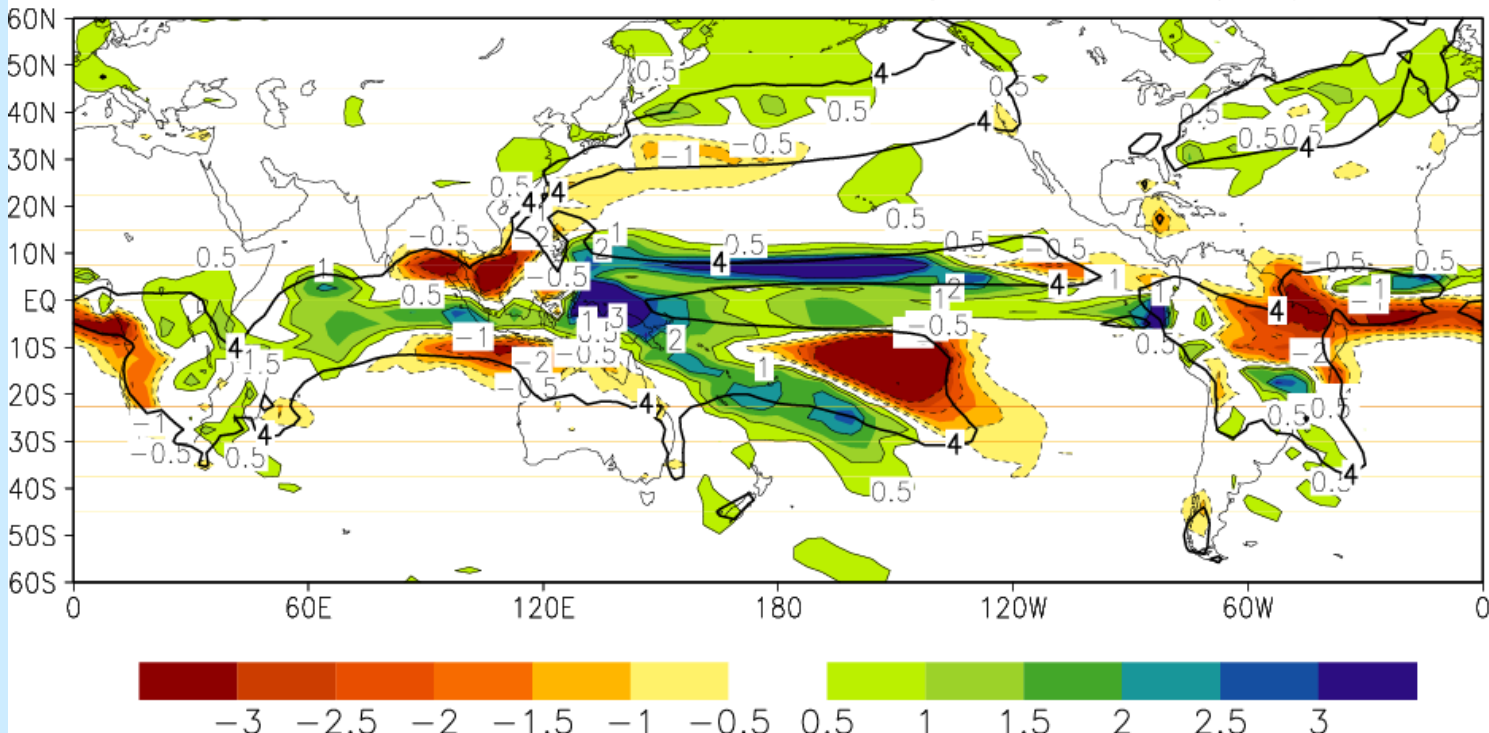
3. Long tails in distributions of column tracers

Issues with precipitation simulation, especially at regional scales, tropics: global warming, El Niño...

- Sensitivity to convective schemes, interaction with large-scale
- [although some agreement on large-scale or amplitude]

Precipitation change: HadCM3, Dec.-Feb., 2070-2099 avg minus 1961-90 avg.

4AR HadCM3 SRESA2 DJF Pa(2070-99) (61-90)



**4 mm/day
model
climatology
black
contour for
reference**

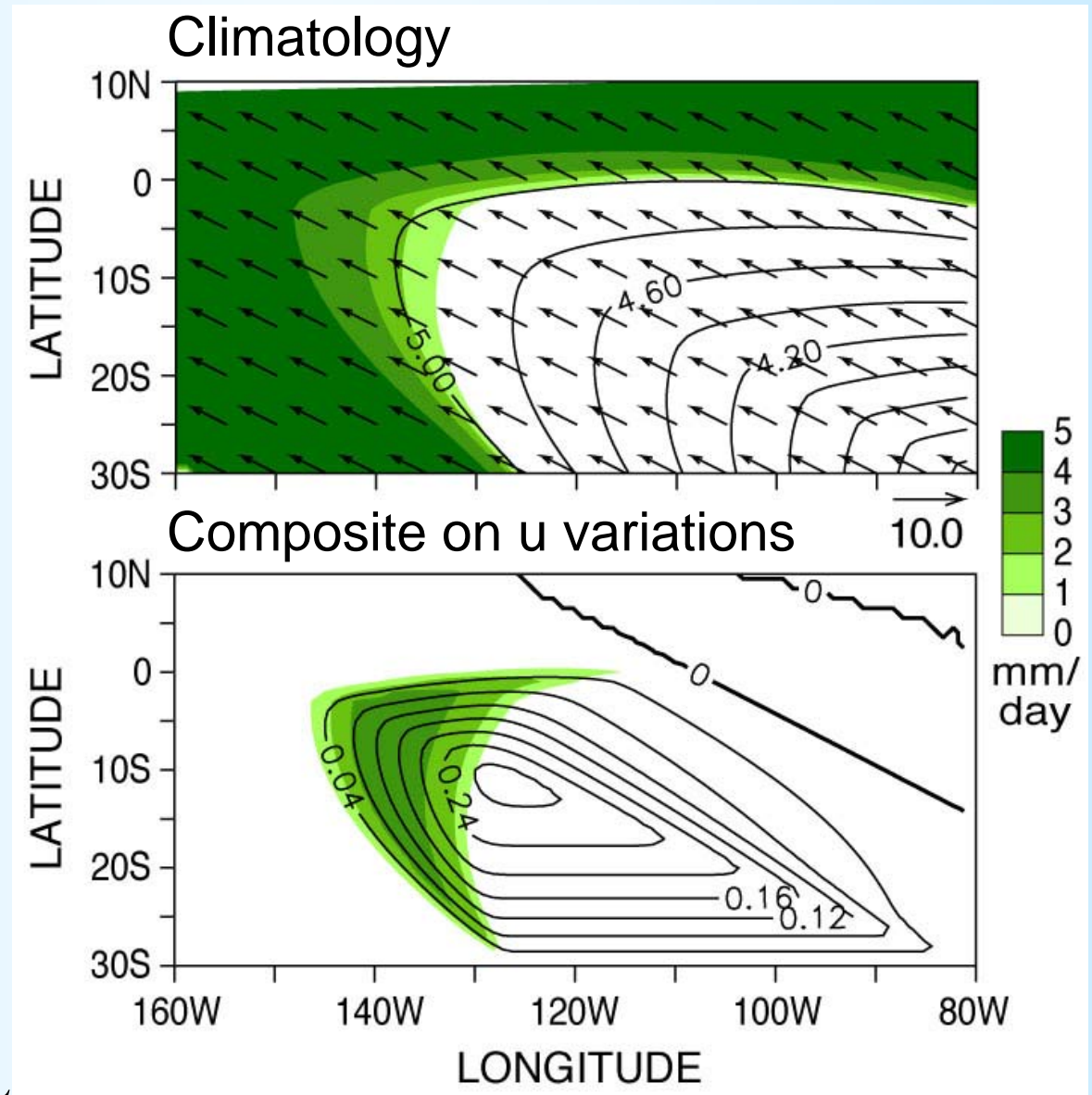
e.g., IPCC 2001, 2007; Trenberth et al 2003; Maloney and Hartmann 2001; Joseph and Nigam 2006; Biasutti et al. 2006; Dai 2006; Tost et al. 2006; Neelin et al 2006; Bretherton 2007...

1. Sensitivity at convective margin

Prototype model* : dry advection into a precipitating region

Precipitation (green) and moisture (contours) would be constant except for trade wind inflow

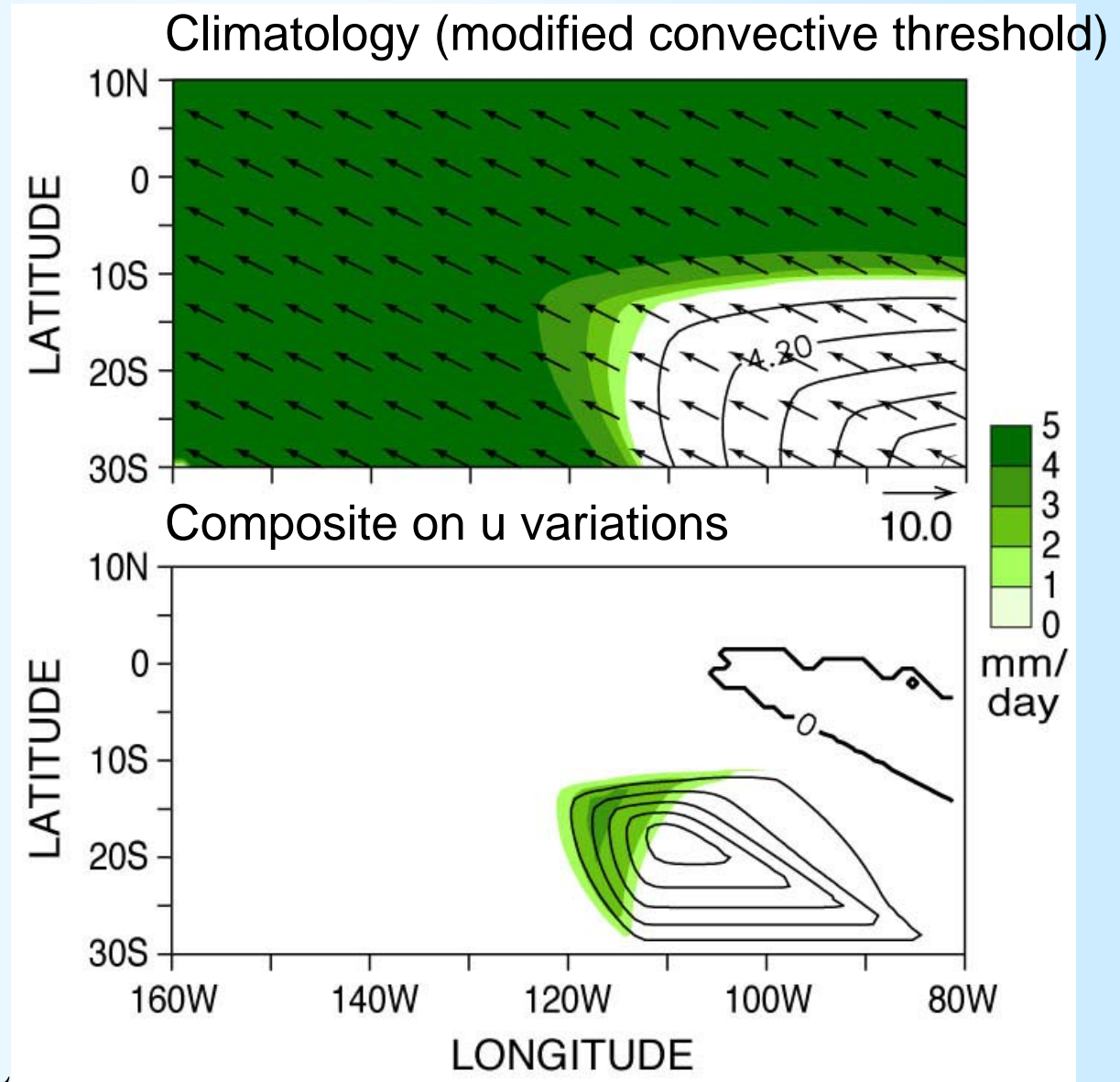
- temp. & moisture equations, specified wind + Gaussian variations;
- Analytic solutions for interplay with local thermodynamics and convective threshold



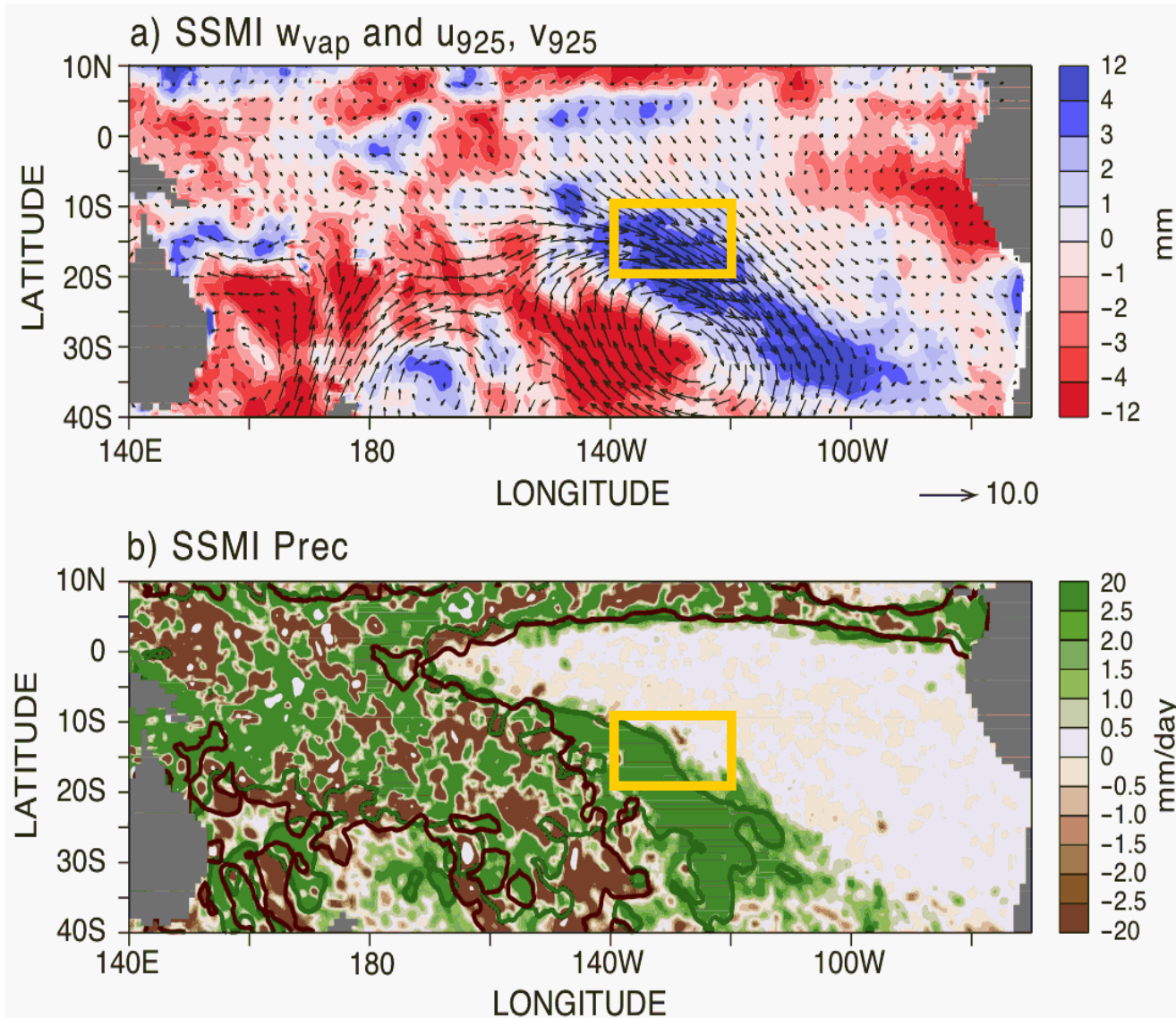
Prototype model: change in threshold for convection

Precipitation
(green) and
moisture (contours)

Substantial impact
of a poorly
constrained aspect
of convective
schemes



South Pacific Convergence Zone (SPCZ) composites: SSMI precip, column water vapor on wind variations



Daily SSMI

Composites on
 u_{925} mb avgd
140°W-120°W,
20°S-10°S.

4 mm day⁻¹ (weaker
trades/less low-level
dry air inflow)

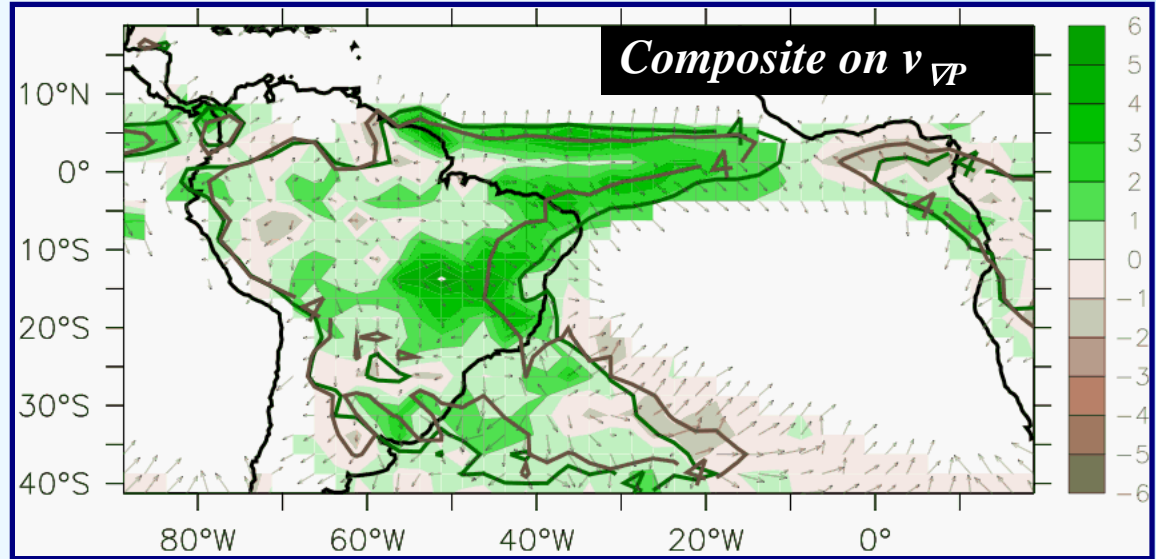
4 mm day⁻¹ (stronger
trades/more low-level
dry air inflow)

Precip. composite on local inflow wind anomaly

Inflow wind $v_{\nabla P}$ across
gradient of mean
precipitation

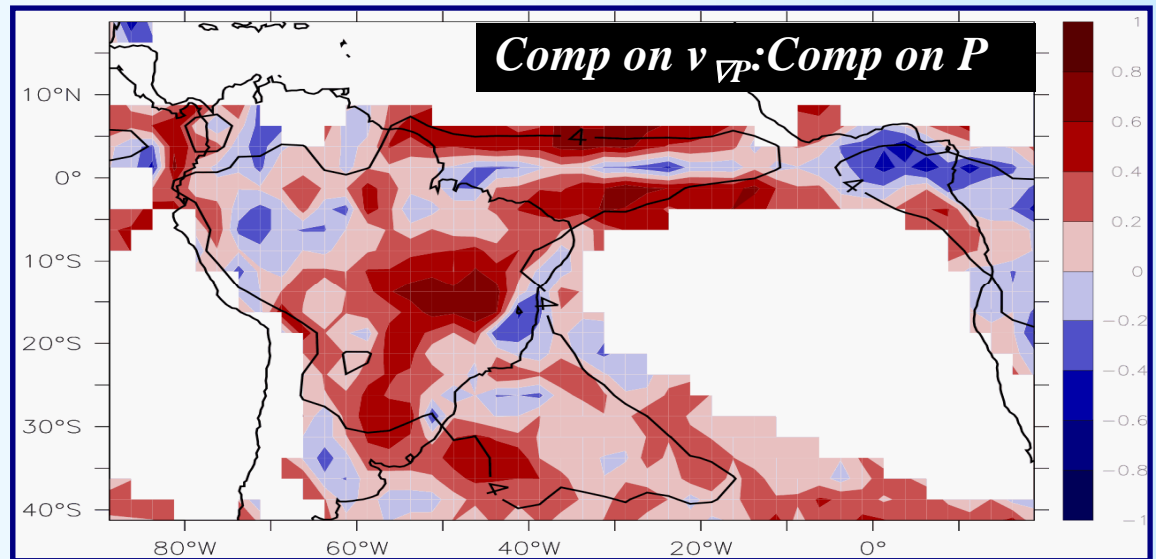
Atmospheric boundary
layer (ABL) wind

Large sensitivity at
margin



Ratio to composite on
precipitation

Locally, monthly
composite precipitation
differences associated
with inflow represent 80-
90% of total composite-
differenced precipitation



2. Transition to strong convection

- **Convective quasi-equilibrium assumptions: Above onset threshold, convection/precip. increase keeps system close to onset** Arakawa & Schubert 1974; Betts & Miller 1986; Moorthi & Suarez 1992; Randall & Pan 1993; Zhang & McFarlane 1995; Emanuel 1993; Emanuel et al 1994; Bretherton et al. 2004; ...
- **Pick up a function of buoyancy-related fields – temperature T & moisture (here column integrated moisture w)**
- **Elsewhere: Onset of strong convection conforms to list of properties for continuous phase transition with critical phenomena** (Peters & Neelin 2006, Nature Physics); **mesoscale implications** (Peters, Neelin & Nesbitt 2009, JAS)
- **Stochastic convective schemes (and old-fashioned schemes too) need to better characterize the transition to deep convection**

Precip. dependence on tropospheric temperature & column water vapor from TMI*

- Averages conditioned on vert. avg. temp. \hat{T} , as well as w (T 200-1000mb from ERA40 reanalysis)

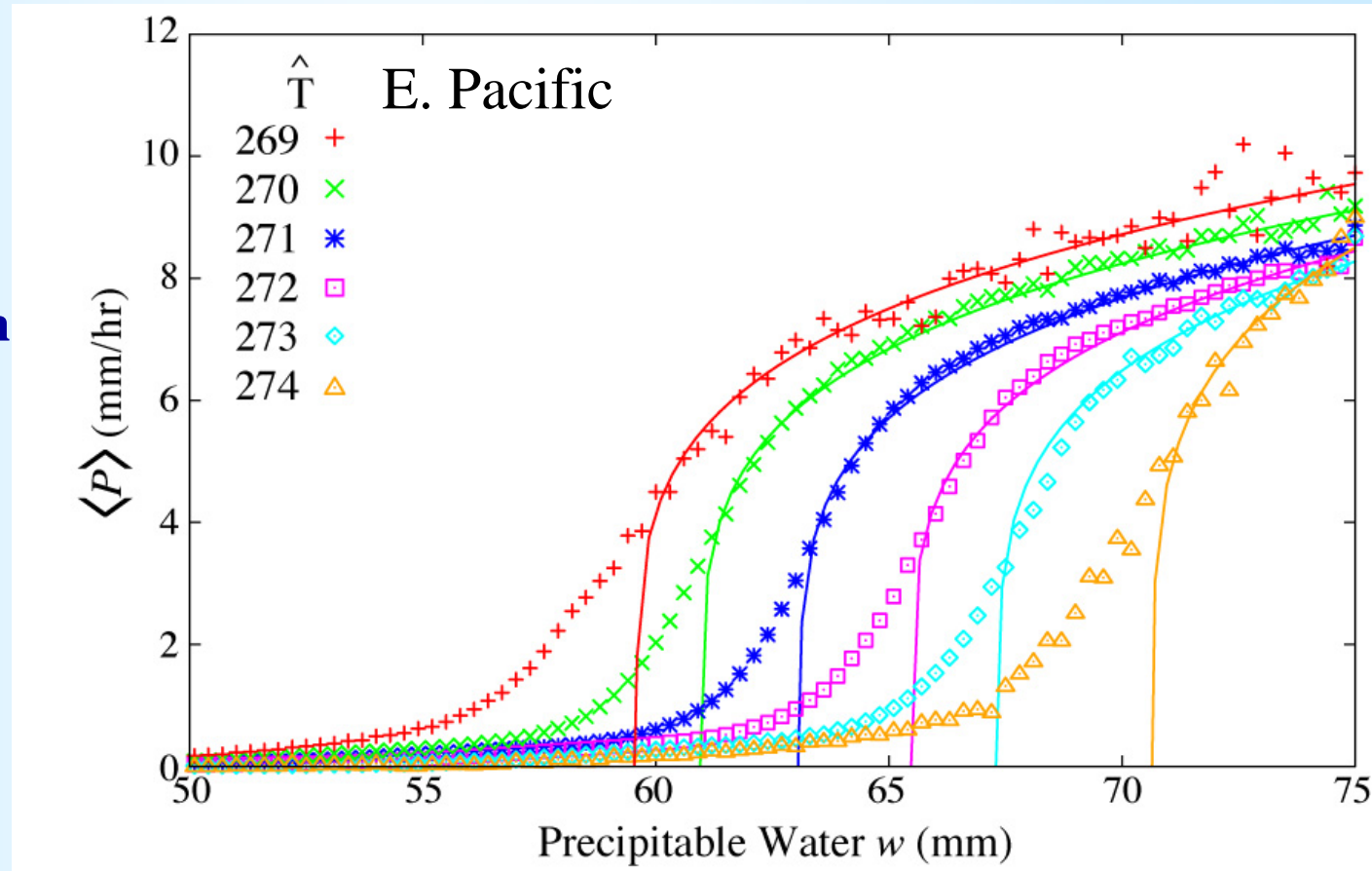
- Power law fits above critical:

$$P(w) = a(w - w_c)^\beta$$

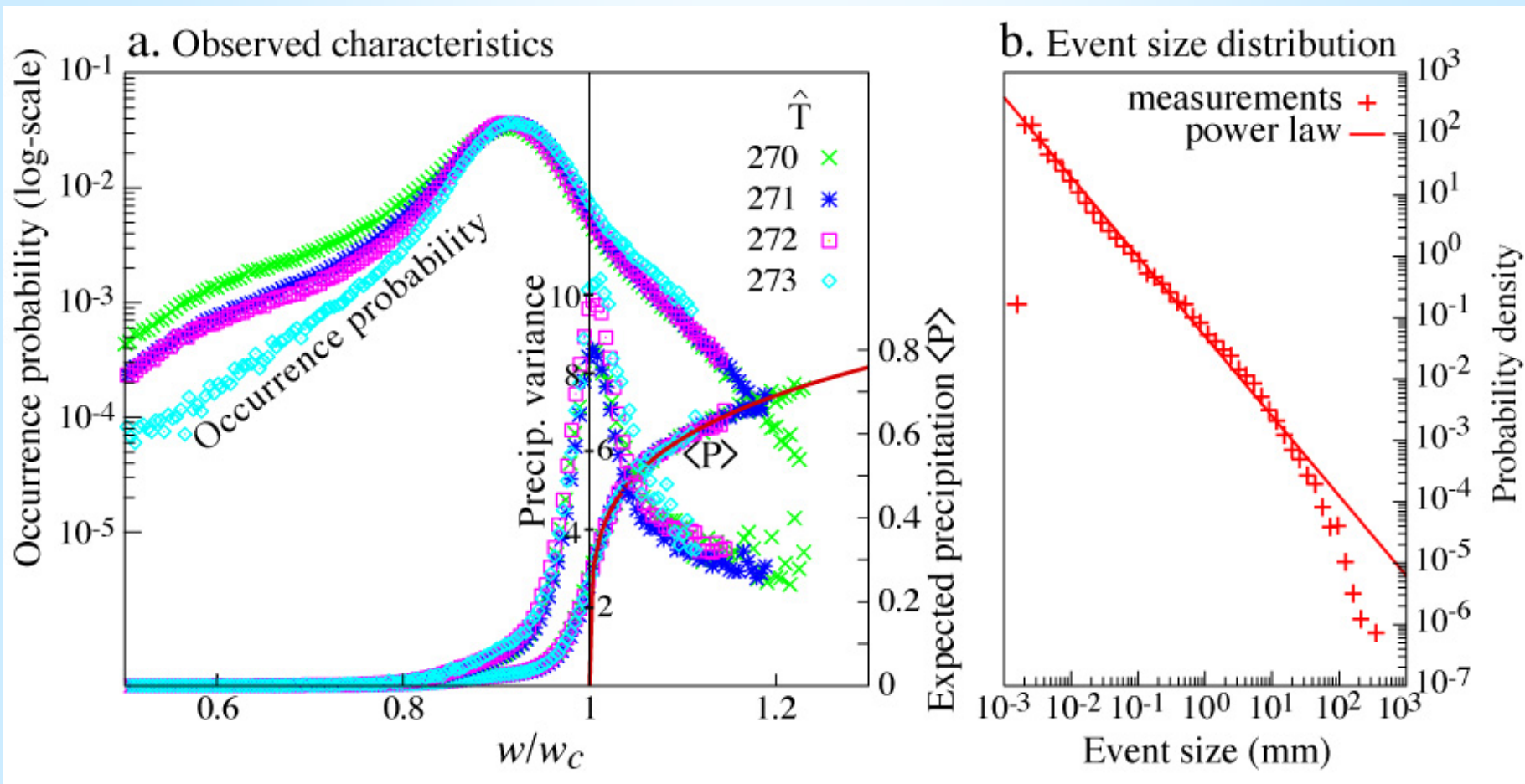
w_c changes, same β

- [note more data points at 270, 271]

*TMI: Tropical Rainfall Measuring Mission Microwave Imager (Hilburn and Wentz 2008), 20N-20S

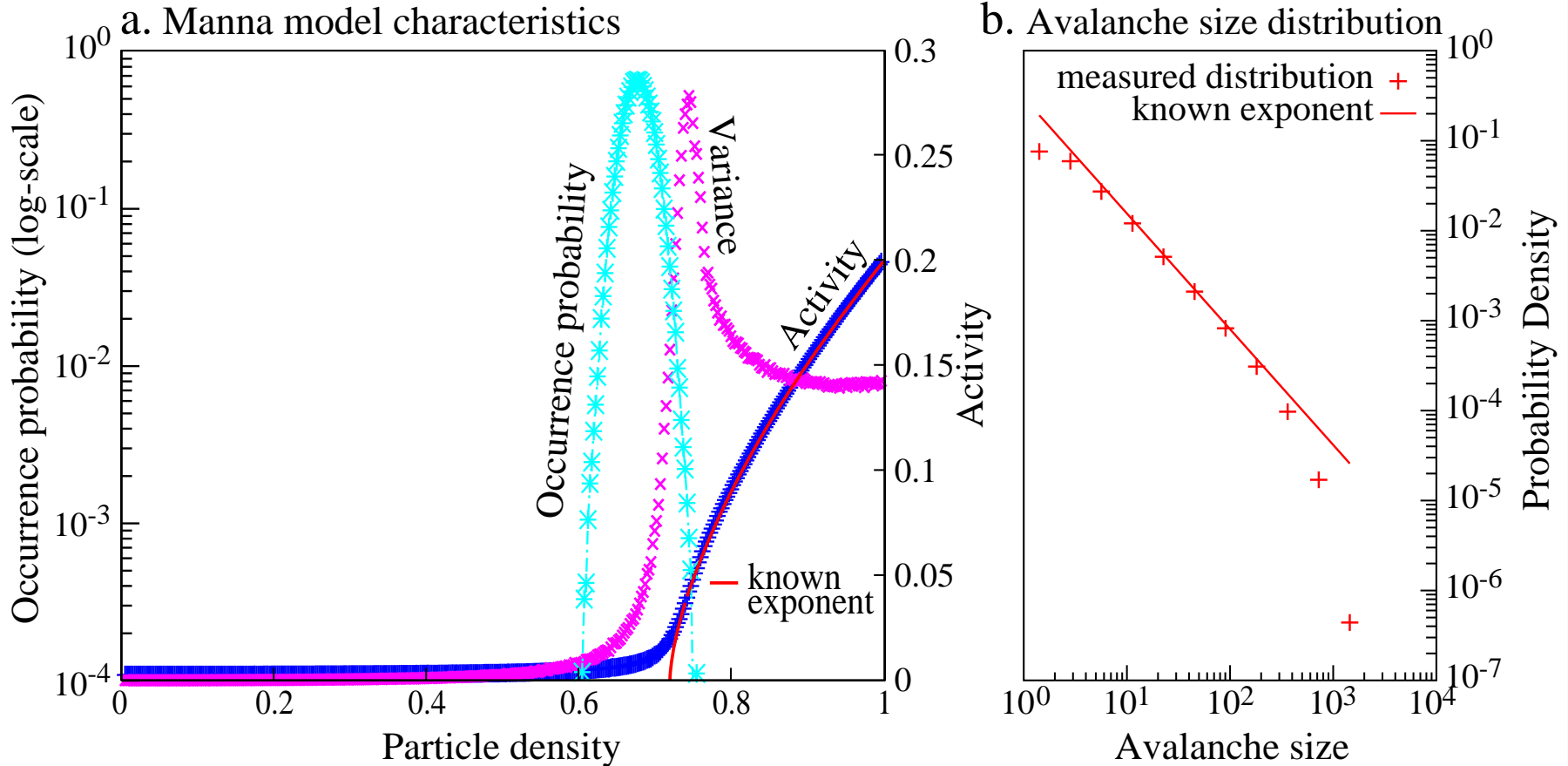


Collapsed statistics for observed precipitation



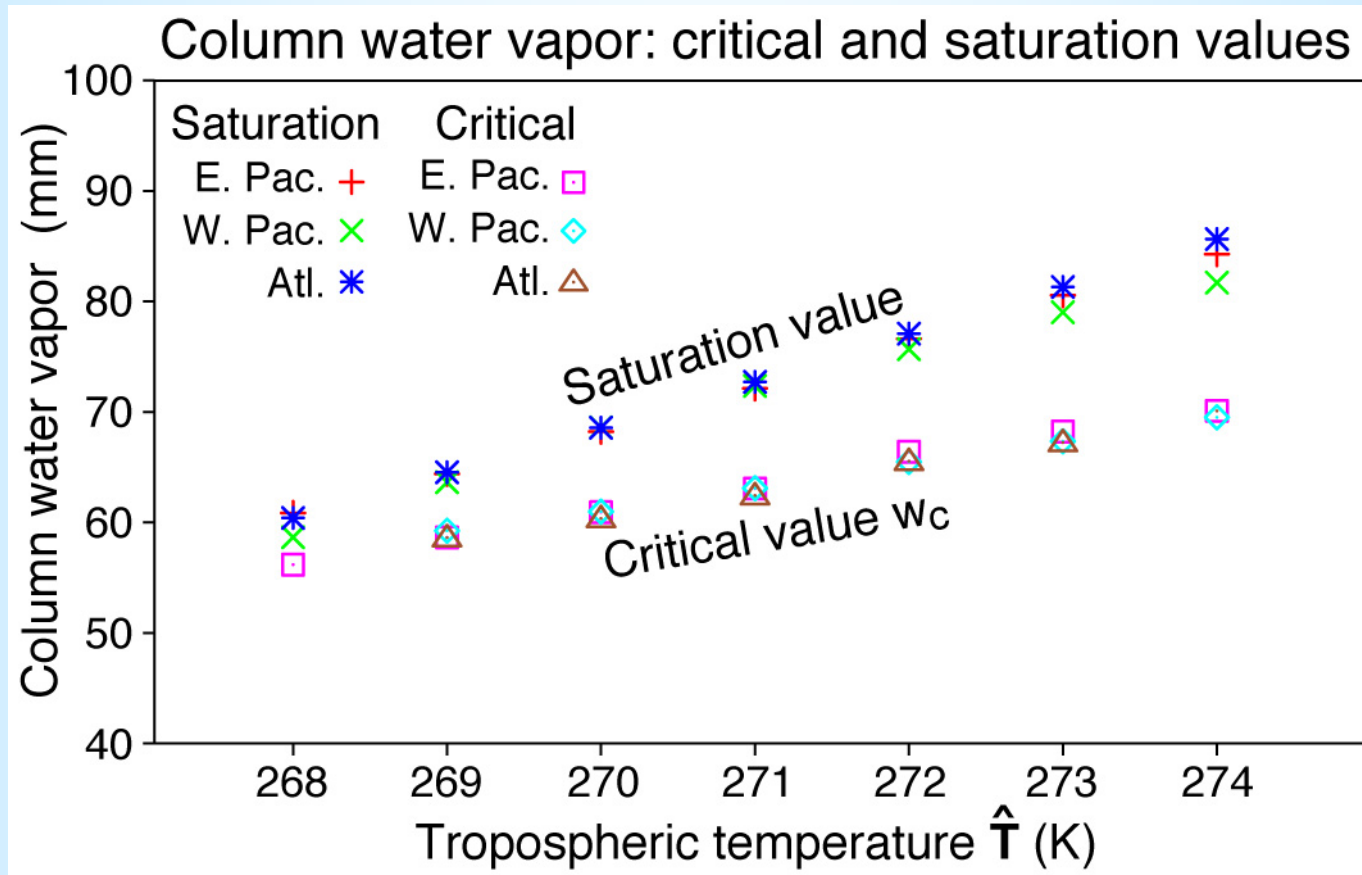
- Precip. mean & variance dependence on w normalized by critical value w_c ; occurrence probability for precipitating points (for 4 T values); Event size distribution at Nauru

Example from Manna (1991) lattice model (hopping particles—not a model of convection! 20x20 grid shown)



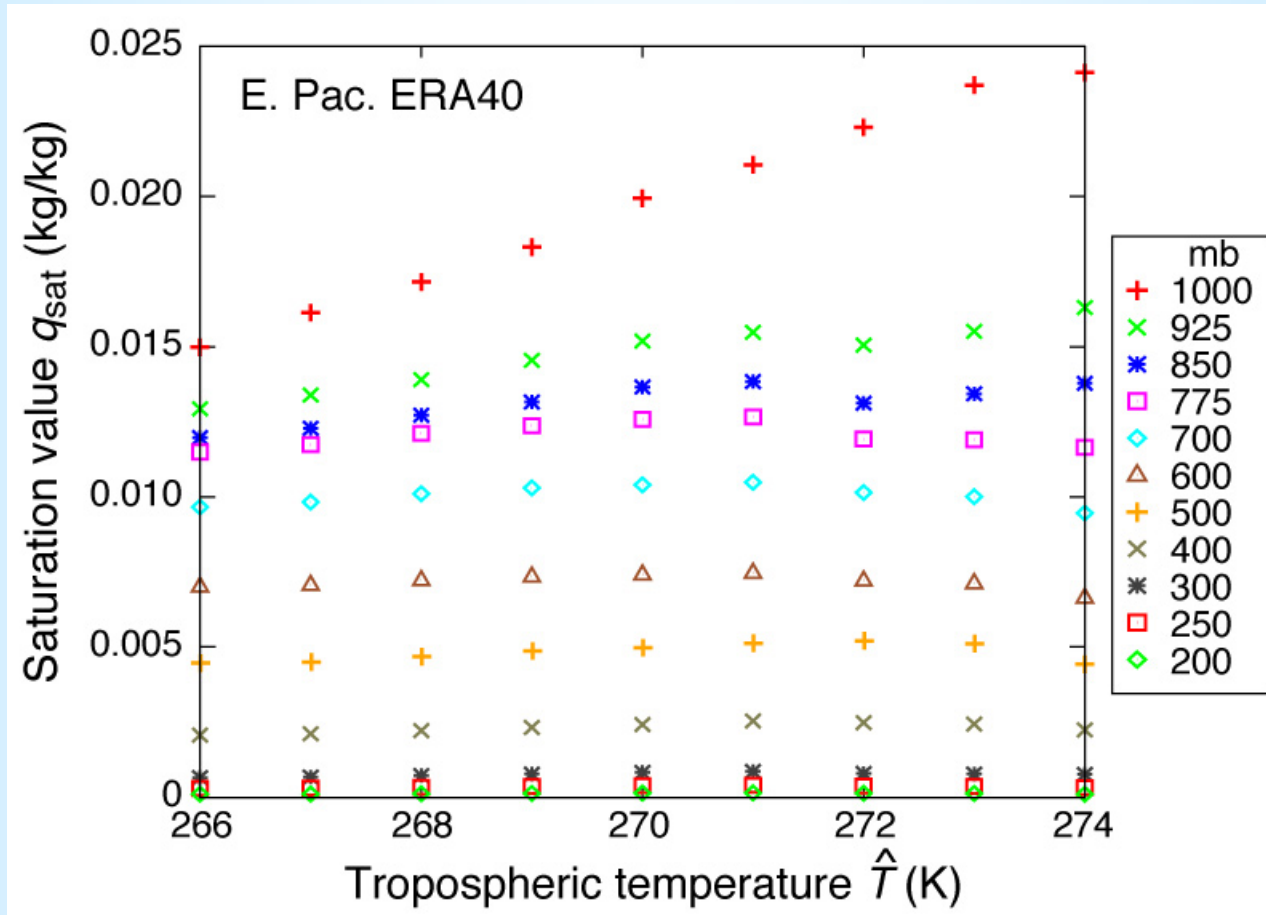
- **Activity (order parameter) & variance dependence on particle density (tuning parameter) [conserving case]**
- **Occurrence probability (log scale; very Gaussian) & event size distribution [self organizing case]**

Critical point dependence on temperature



- Find critical water vapor w_c for each vert. avg. temp. \hat{T}
- Compare to vert. int. saturation vapor value binned by same \hat{T}
- *Not* e.g., a constant fraction of column saturation
- lower tropospheric saturation $q_{\text{sat}}(T)$ binning gives same results

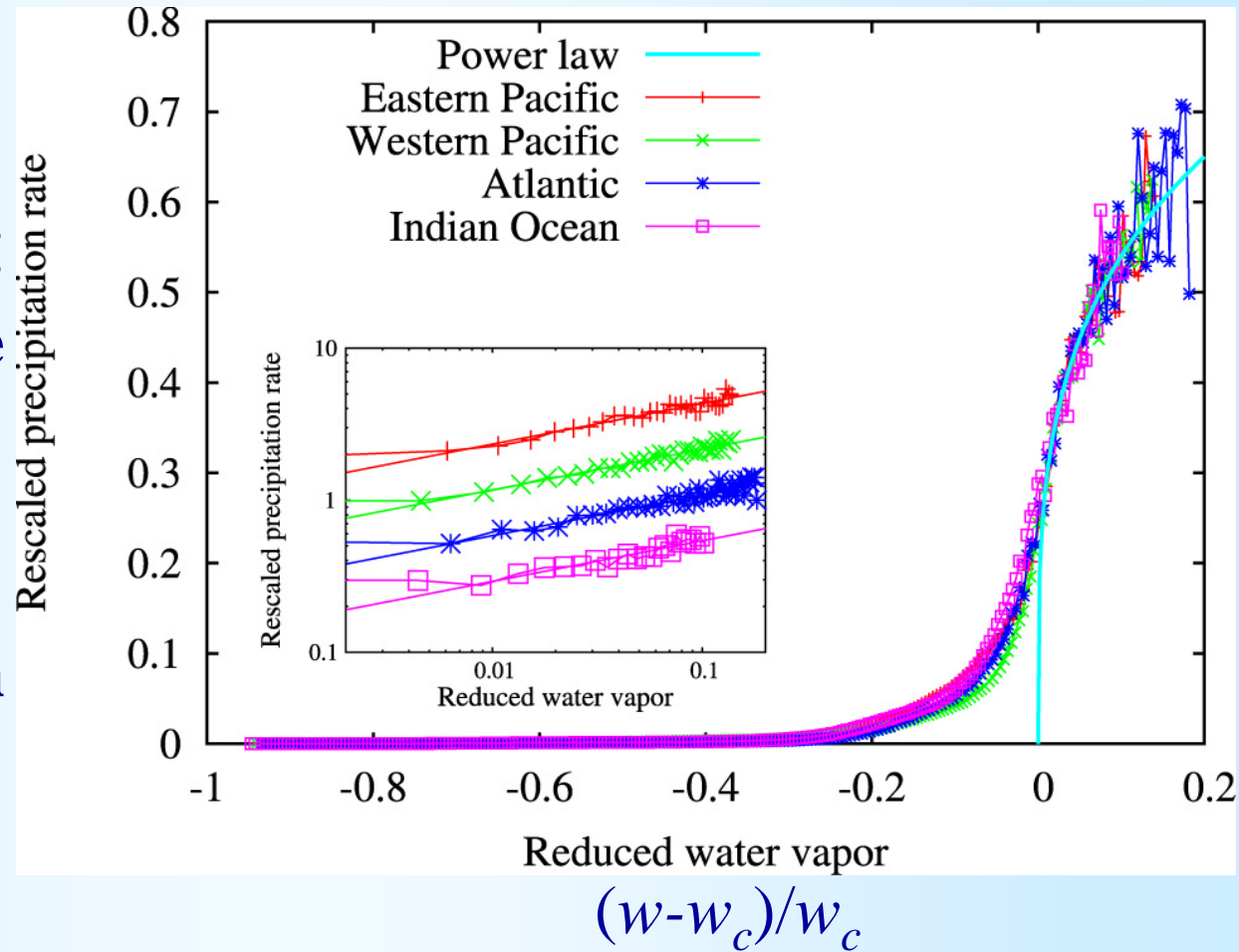
Saturation value $q_{\text{sat}}(T)$ by level



- Saturation mixing ratio by level binned by vert. avg. temp. \hat{T}
- Compare to critical value & vert. int. saturation value vs. \hat{T}
- Appears consistent with substantial control by lower free troposphere proximity to saturation

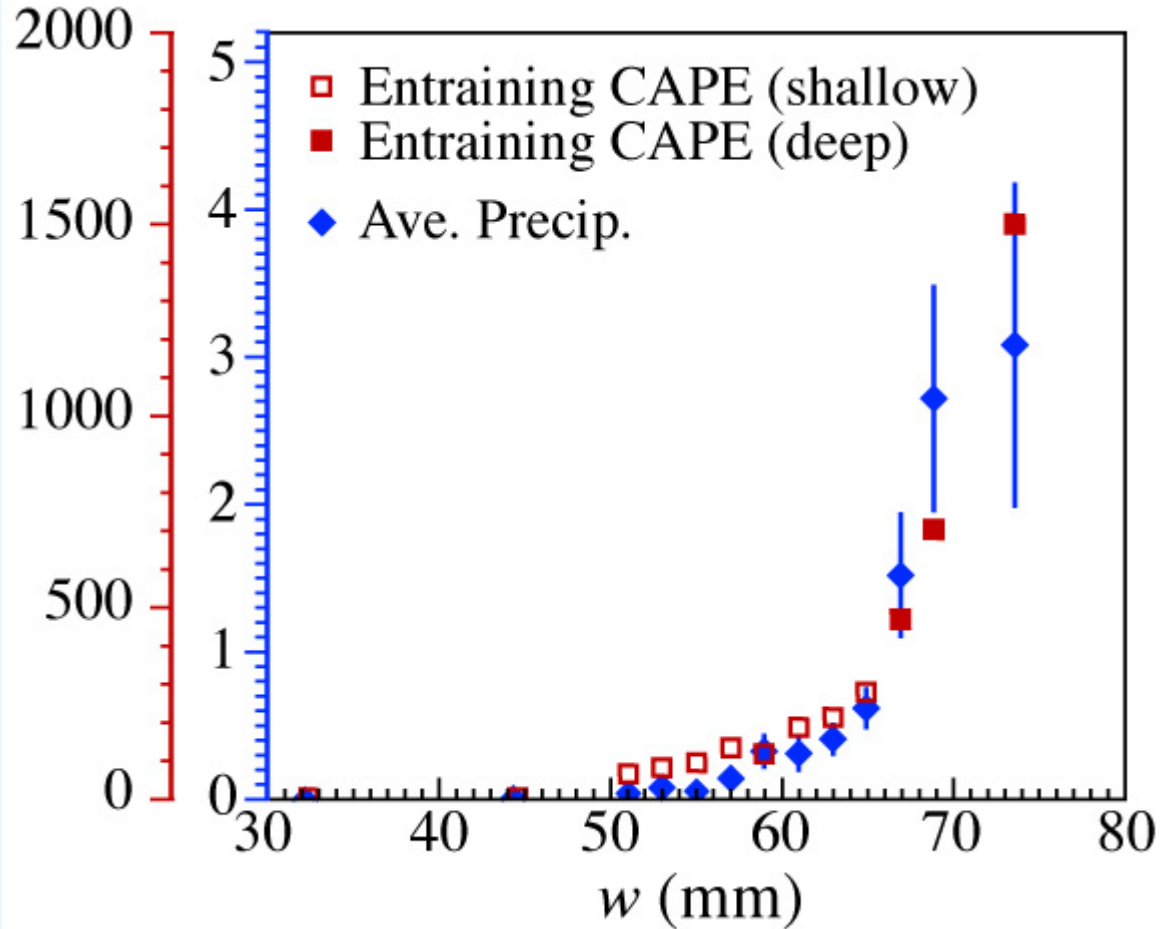
Check pick-up with radar precip data

- TRMM radar data for precipitation
- 4 Regions collapse again with w_c scaling
- Power law fit above critical even has roughly same exponent as from TMI microwave rain estimate
- (2A25 product, averaged to the TMI water vapor grid)



Entraining convective available potential energy and precipitation binned by column water vapor, w

- buoyancy & precip. pickup at high w
- boundary layer and lower free troposph. moisture contribute comparably*
- consistent with importance of lower free tropospheric moisture (Austin 1948; Yoneyama and Fujitani 1995; Wei et al. 1998; Raymond et al. 1998; Sherwood 1999; Parsons et al. 2000; Raymond 2000; Tompkins 2001; Redelsperger et al. 2002; Derbyshire et al. 2004; Sobel et al. 2004; Tian et al. 2006)

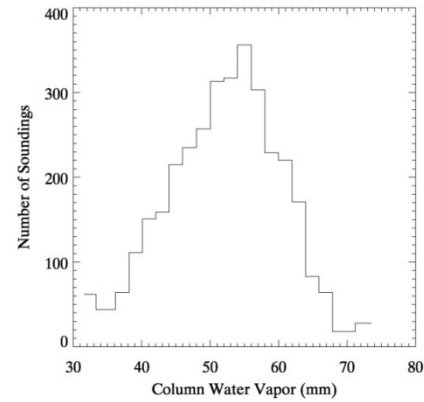
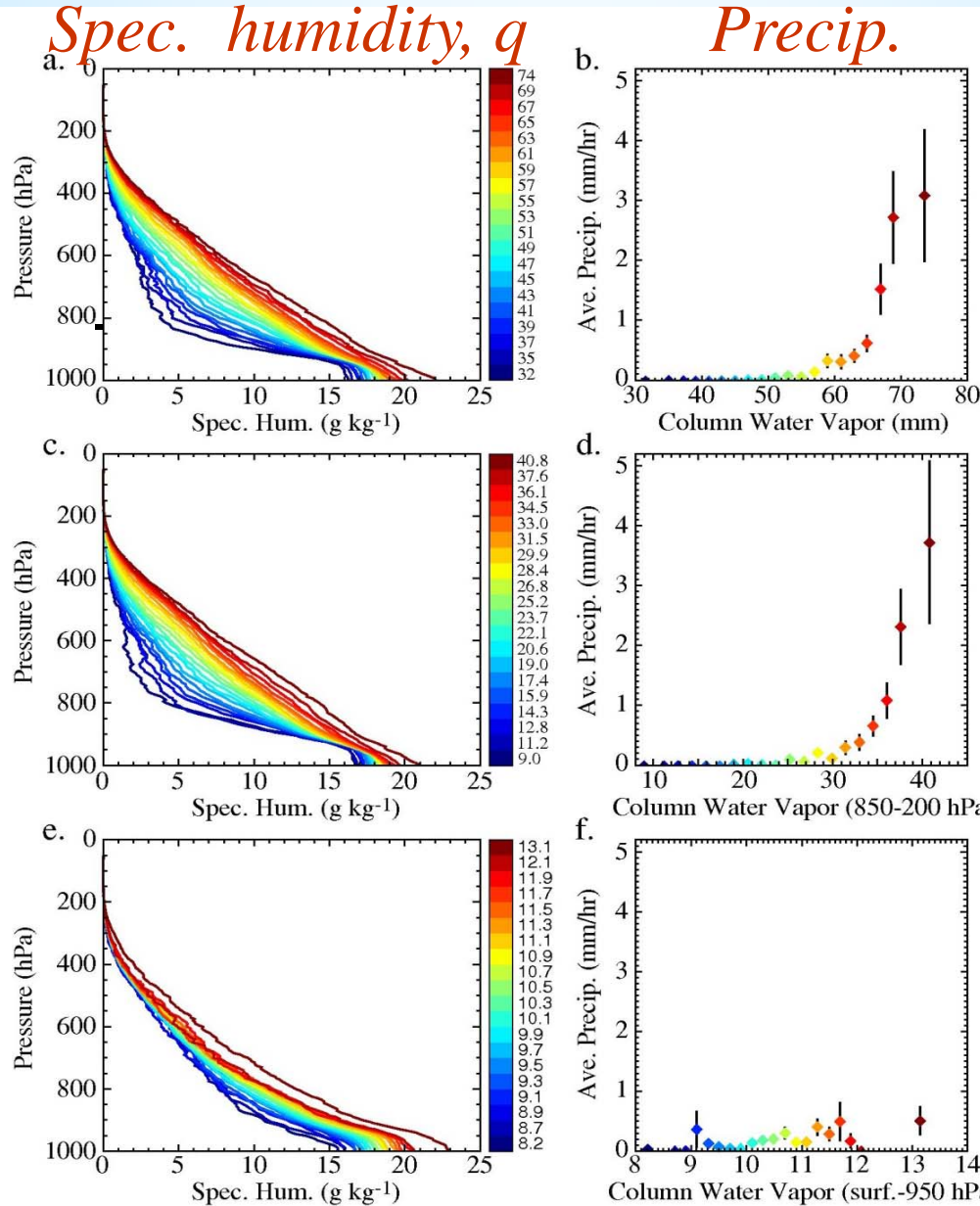


Binning q , precip. on vert. int. water vapor

Binned by:
Column
water vapor

850-
200 mb

Surface-
950mb

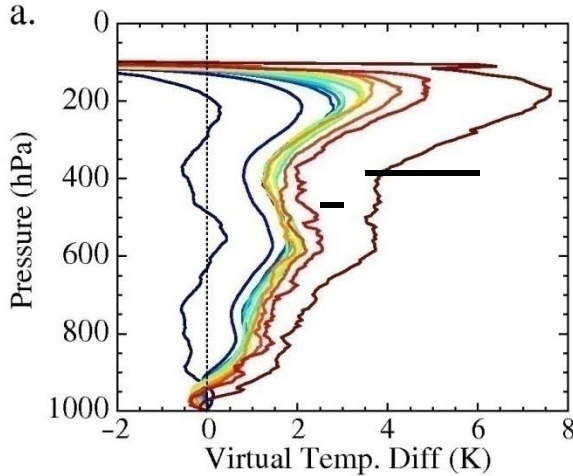


[Note fewer
soundings
in high bins]

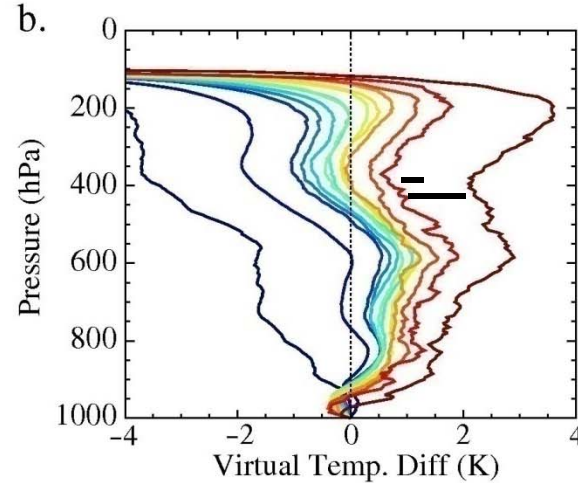
**Nauru ARM
site observations**

Lifted parcel buoyancy by column water vapor bins

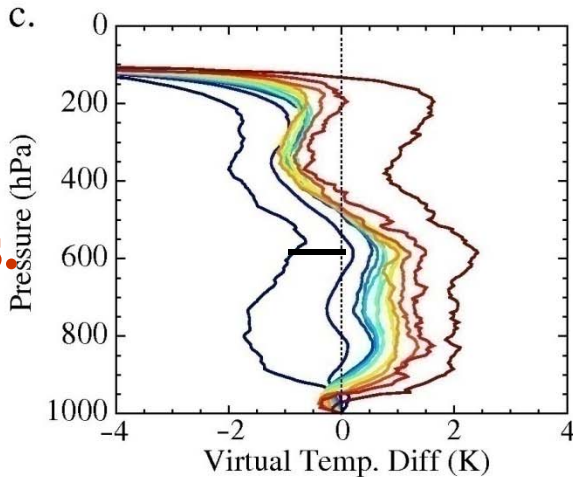
No
mixing



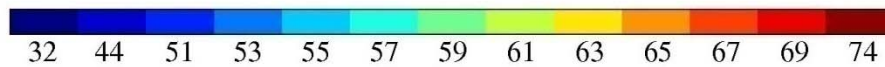
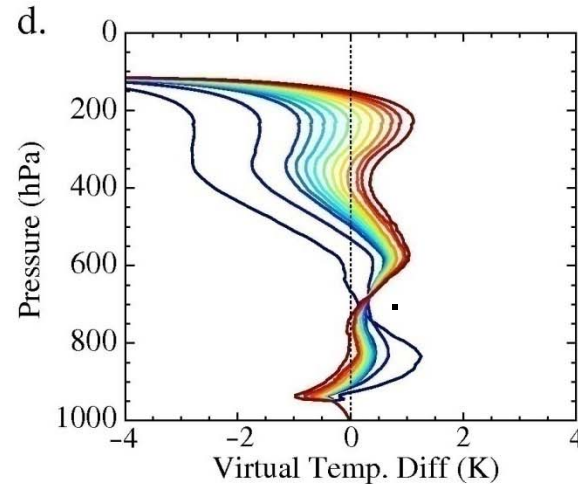
Const.
mixing
(Brown &
Zhang 1997)



Const.
mixing,
with q in
free tropos.
constant



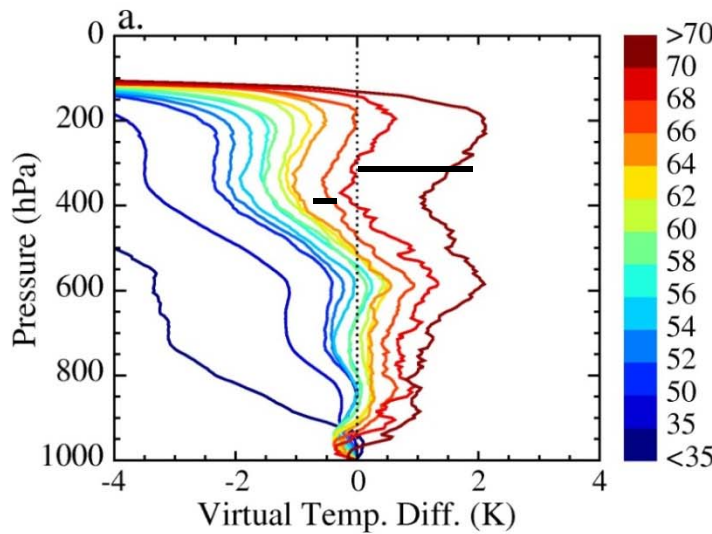
Const.
mixing,
only q in
free tropos.
changes



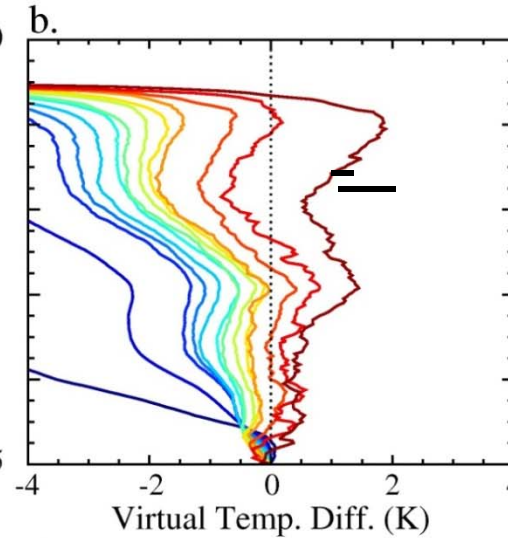
- Highest column water vapor bins most buoyant
- Both boundary layer and lower free troposphere contribute

Lifted parcel buoyancy by column water vapor bins

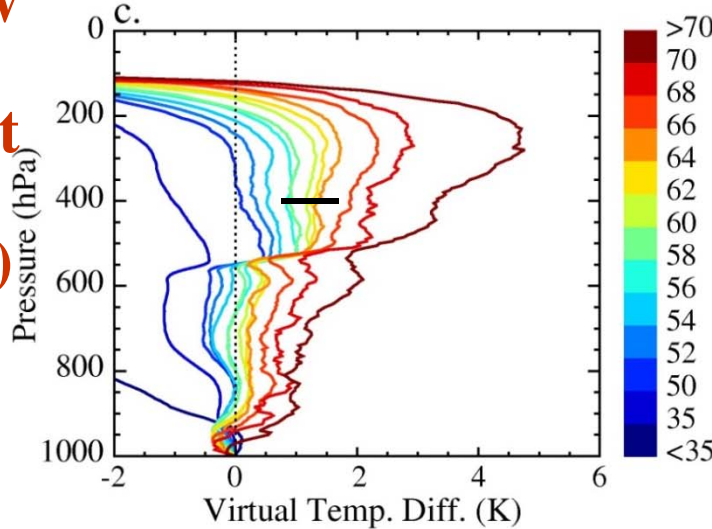
Deep inflow mixing A



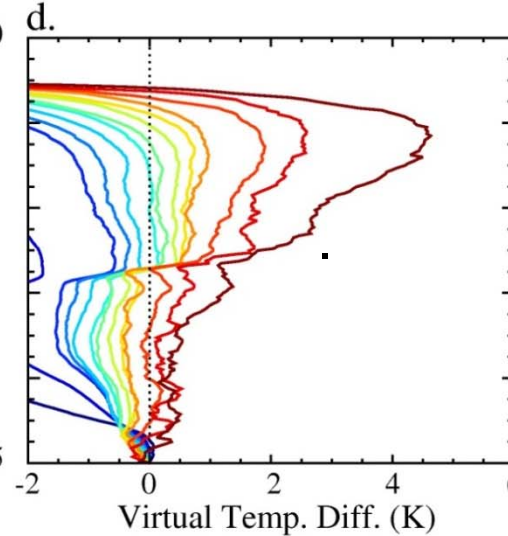
Deep inflow mixing B



Deep inflow mixing A with instant freezing (reversible)



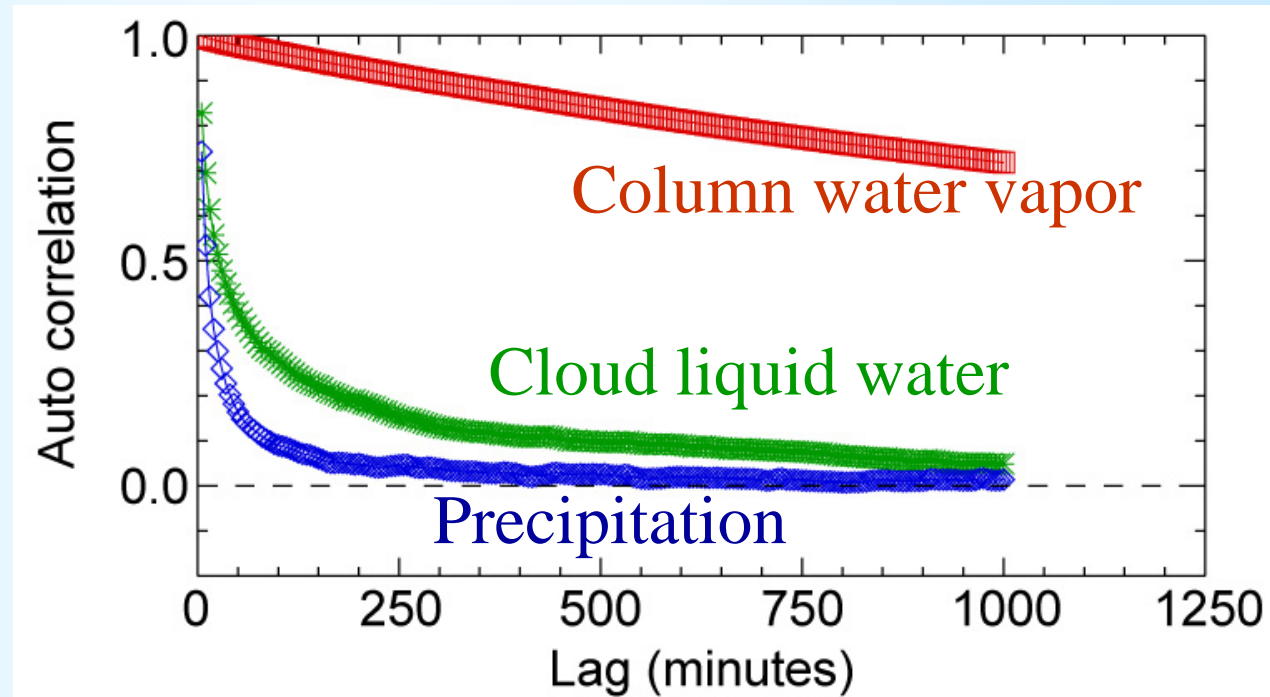
Deep inflow mixing B with instant freezing (reversible)



- Highest few column water vapor bins deep convective
- microphysics between these cases; large potential impact

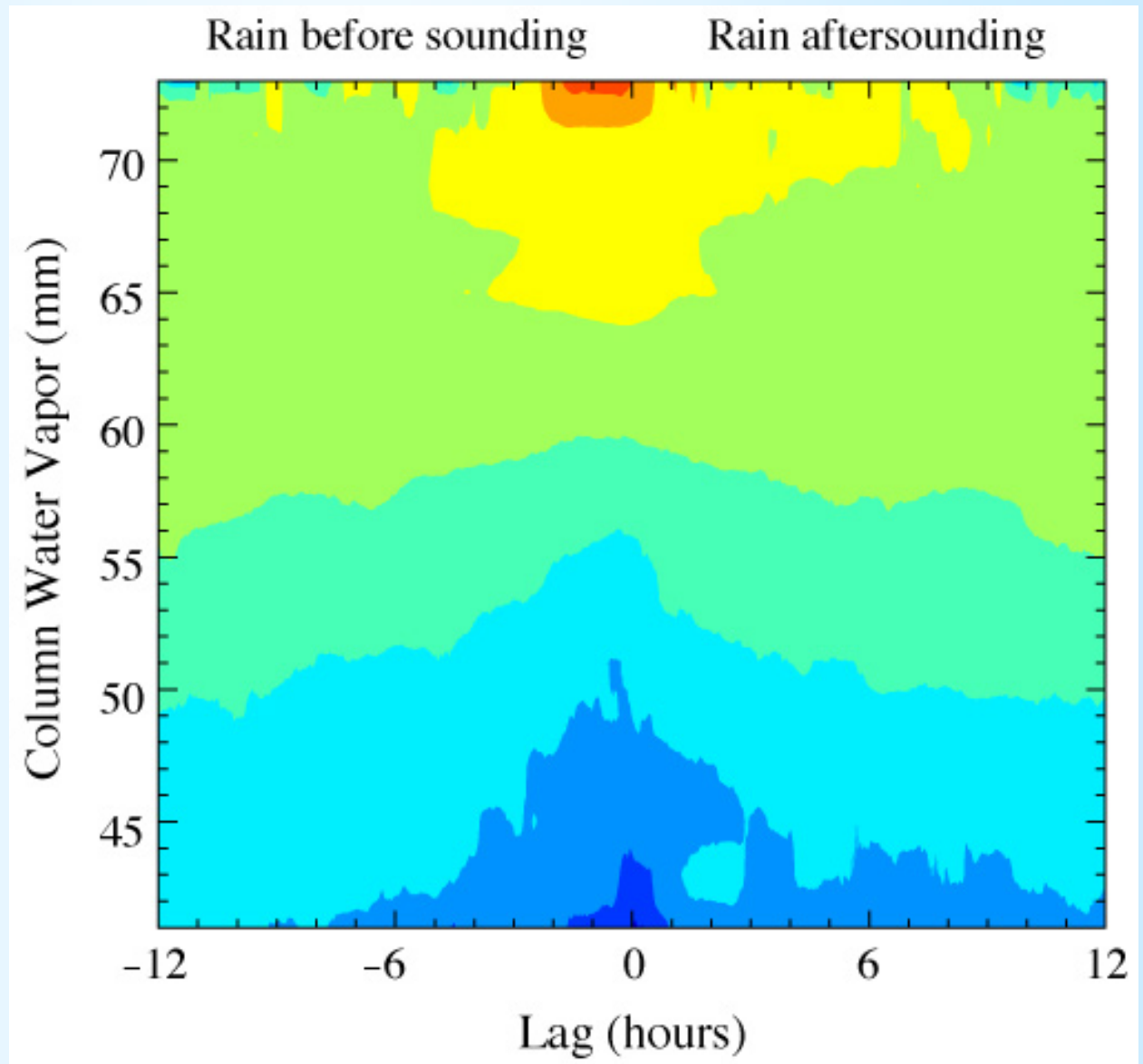
Prec & column water vapor: autocorrelations in time

- Long autocorrelation times for vertically integrated moisture (once lofted, it floats around)
- Nauru ARM site upward looking radiometer + optical gauge



Precip conditioned on lag/lead column water vapor

- High water vapor several hours ahead still useful for pickup in precipitation
- Consistent with high water vapor \Rightarrow favorable environment, but stochastic plume
- Nauru ARM site upward looking radiometer + optical gauge



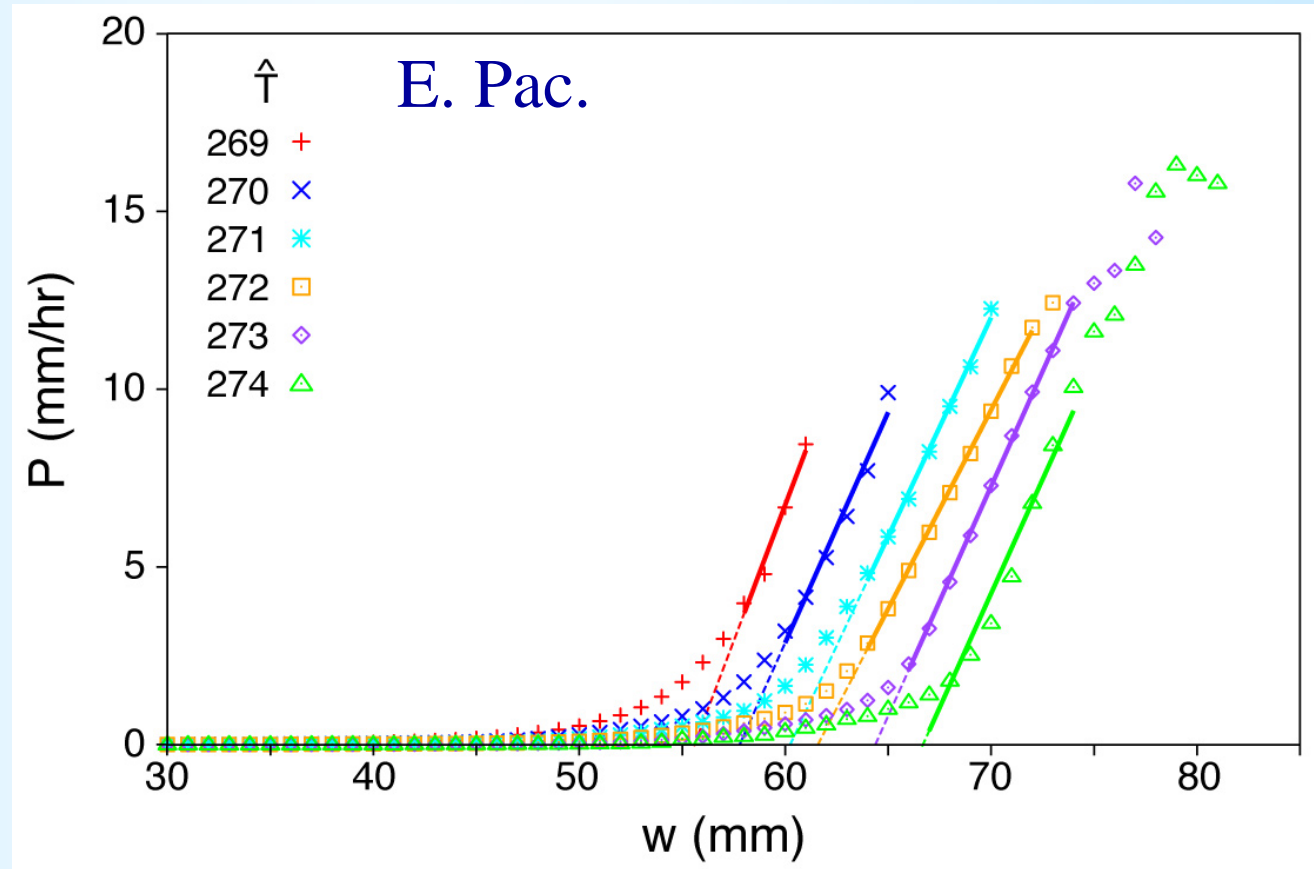
How do models do? CAM3.5 (0.5 degree run) * : Precip. dependence on tropospheric temperature & column water vapor

- Averages conditioned on vert. avg. temp. T , as well as column water vapor w

- Linear fits above critical (motivated by parameterizn)

$$P(w)=a(w-w_c)^\beta$$

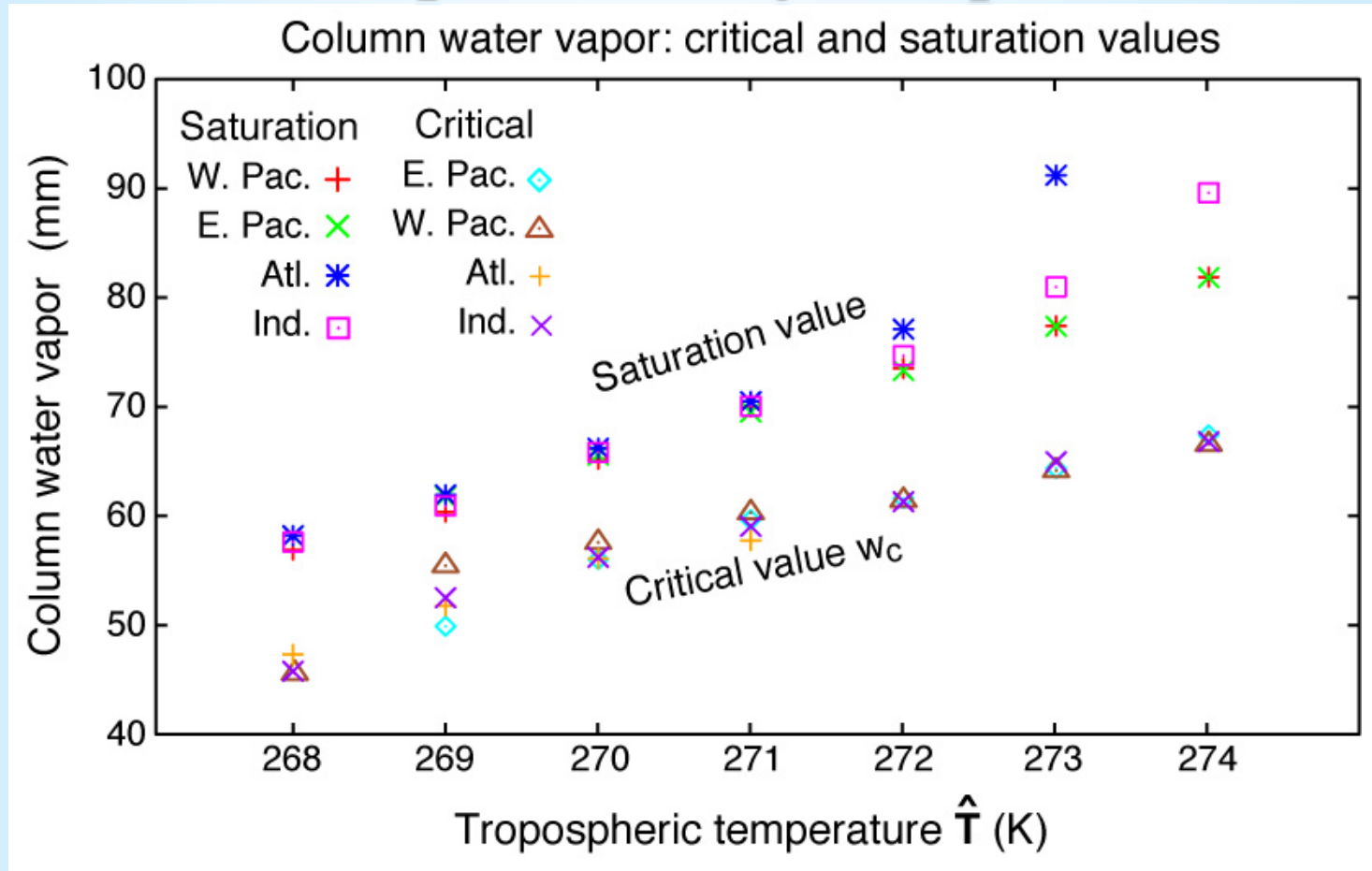
as obs. but $\beta=1$:
to estimate w_c



*Runs, data R. Neale, analysis K. Hales

Critical point dependence on temperature

CAM3.5 preliminary comparison

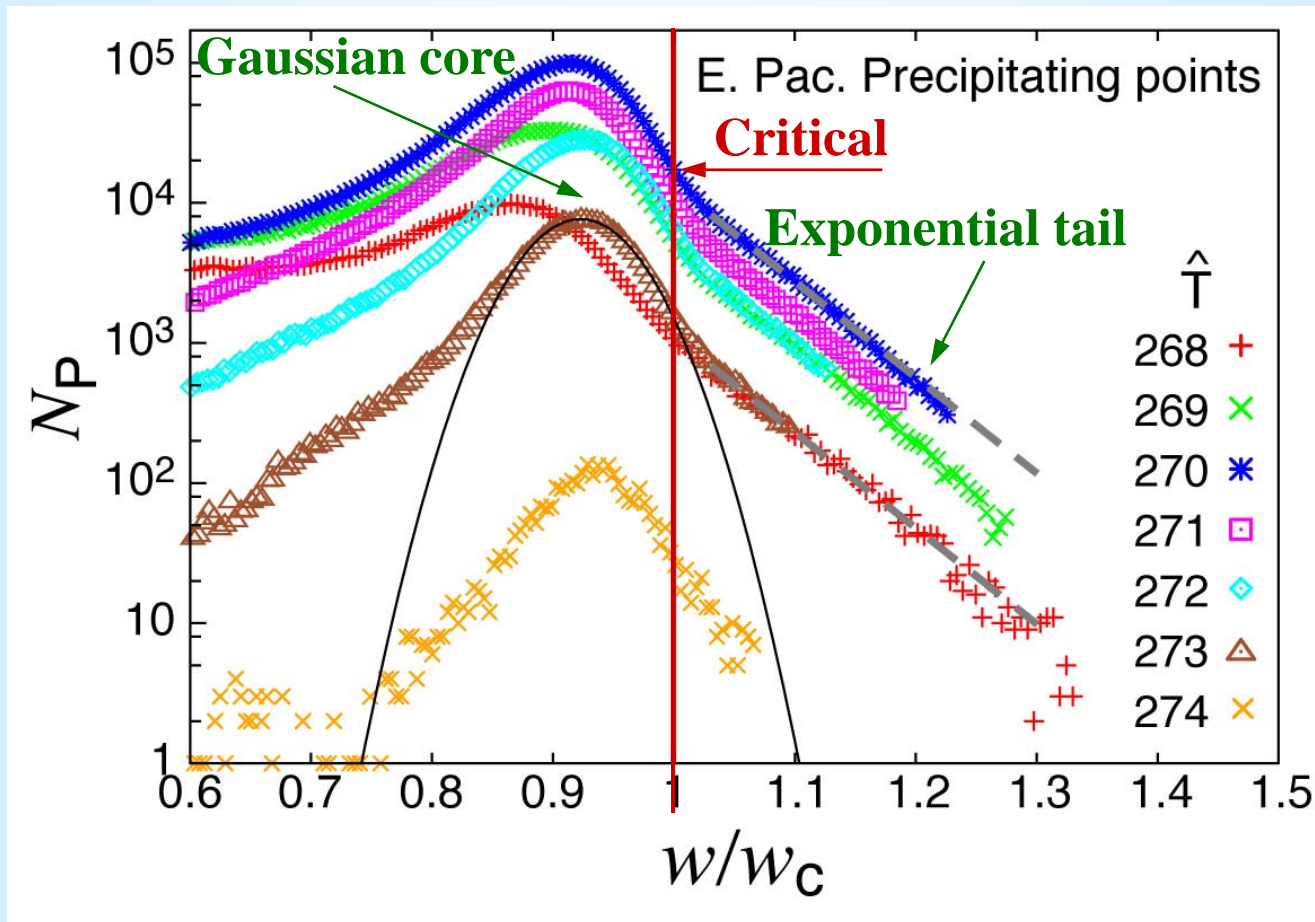


- critical water vapor w_c for each vert. avg. temp. \hat{T}
- Compare to vert. int. saturation vapor value binned by same \hat{T}
- Suggests suitable entraining plumes can capture T dependence

Obs. Freq. of occurrence of w/w_c (precipitating pts)

Eastern Pacific for various tropospheric temperatures

- Peak just below critical pt. \Rightarrow self-organization toward w_c
- But exponential tail above critical pt. \Rightarrow more large events
- with Gaussian core, akin to forced tracer advection-diffusion problems (e.g. Shraiman & Siggia 1994, Pierrehumbert 2000, Bourlioux & Majda 2002)

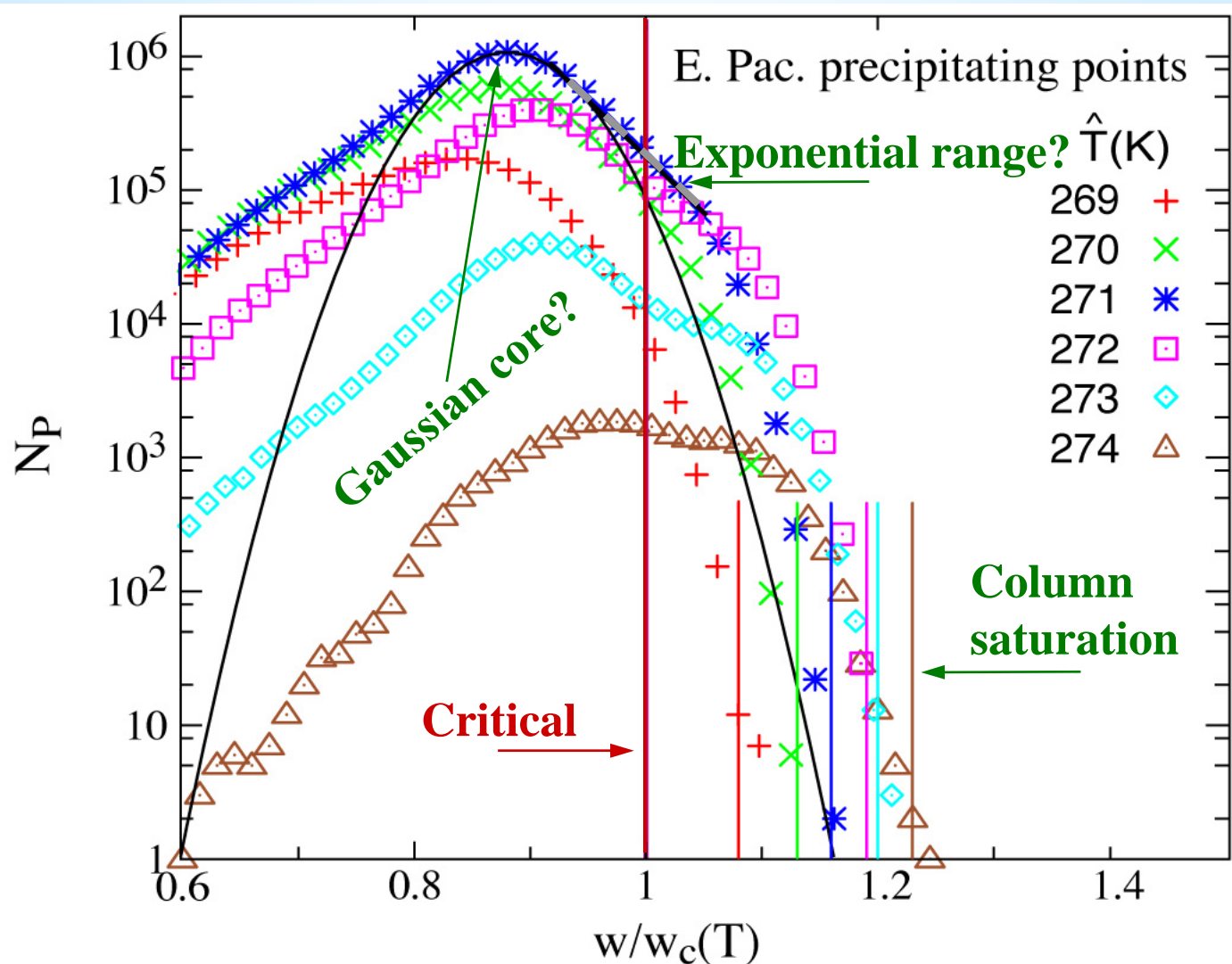


Precipitating freq. of occurrence vs. w/w_c

Eastern Pacific for various tropospheric temperatures

•CAM3.5 preliminary comparison

•Includes super-Gaussian ~exponential range above critical pt.



Summary

- **These statistics for precipitation and buoyancy related variables at short time scales provide new ways to quantify the transition to tropical deep convection as needed for models**
- **Tracer distributions consistent with simple prototypes; core with stretched exponential tails ubiquitous**

Current retrievals are great but could sure use

- **vertical dependence on temperature and water vapor in deep convective regions, land,...**
- **Coordinated observations of condensate loading, freezing**
- **huge number of observations allow statistics to be computed consistently through range with large events**
- **Multiple tracers promising**