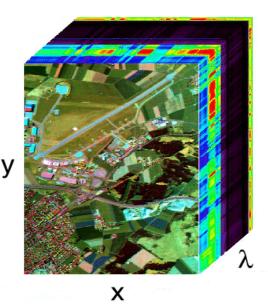
Potential of hyperspectral imagery for geophysical applications

Rodolphe MARION* & Rémi MICHEL <u>rodolphe.marion@cea.fr</u>

*Chairman, French Hyperspectral Permanent Working Group, 2009



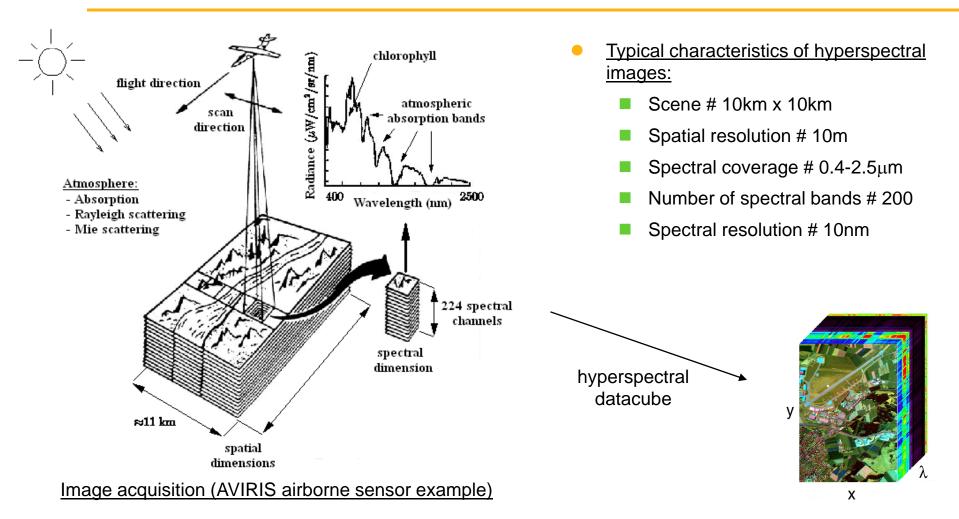
Monitoring Earth Surface Changes from Space

October 28 - 30, 2009



KISS Workshop, Monitoring Earth Surface Changes from Space, Caltech, October 28-30 2009

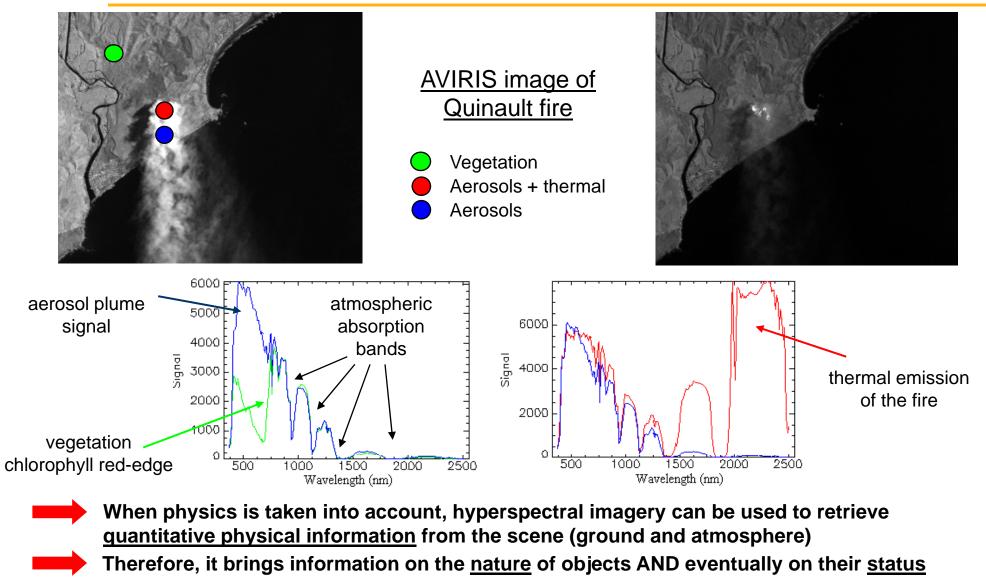
Principles of hyperspectral data acquisition



Each pixel of the image is a « continuous » spectrum containing detailed information on ground and atmosphere

Hyperspectral imagery is generally qualified as imaging spectroscopy

Physics of the image



⁽stressed vegetation, wet soil...)

Hyperspectral imagery is <u>complementary</u> to other modalities (visible, thermal, radar) KISS Workshop, Monitoring Earth Surface Changes from Space, Caltech, October 28-30 2009

A favorable international context

- œ
- Current sensors (mainly airborne):
 - AVIRIS ([0.4-2.5μm], 224 bands, NASA)
 - CASI ([0.4-0.95µm], 20 bands, Itres)
 - HYDICE ([0.4-2.5μm], 210 bands, NRL)
 - DAIS ([0.4-12.5μm], 211 bands, DLR)
 - HyMap ([0.4-2.5μm], 126 bands, HyVista)
 - MIVIS ([0.4-12.5μm], 102 bands, Sensytech)
 - APEX ([0.4-2.5μm], 300 bands, ESA)
- Satellite projects:
 - Artemis (USA) : launched May 2009 (US military applications only)
 - **EnMap ([0.4-2.5μm], 270 bands, 30m, DLR) : 2012-2013**
 - PRISMA ([0.4-2.5µm], 200 bands, 20m + PAN 5m, ASI) : 2012-2013
 - Others: Hyper-X (Japan), HERO (Canada), MSMI (South Africa) ???
 - France : CNES working group, phase 0...



Increased availability of hyperspectral data in a near future...



CASI airborne sensor



Hyperion sensor onboard EO-1 satellite

Overview of scientific applications

Atmospheric parameters (gas, aerosols...) useful for:

Surface phenomenon detection (fires, volcanoes, methane emissions...)

Pollution estimation in the boundary layer

Geosciences:

Geology, mineralogy, mining and oil prospecting, soils quality, degradation and pollution, volcanism, crisis management...

• Vegetation:

- Biochemical content (pigments, water, dry matter...)
- Canopy structure (Leaf Area Index...)
- Fluorescence, stress detection...

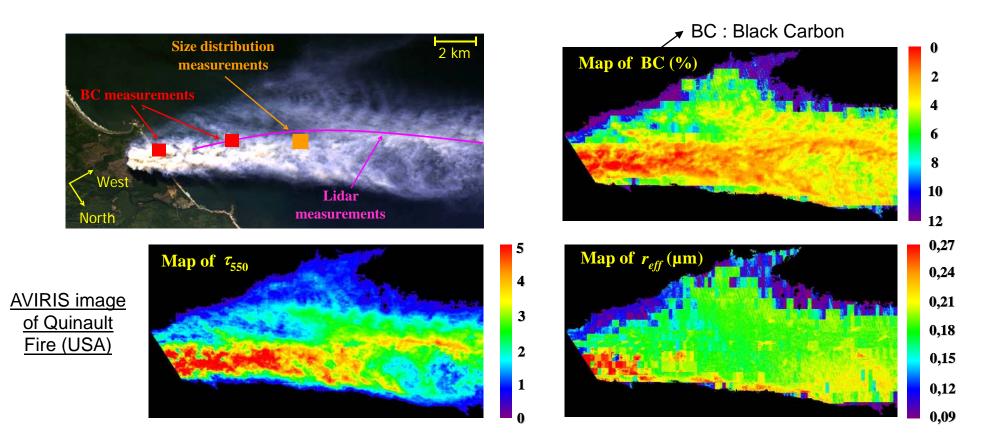
• Urban and natural hazards:

Hyperspectral data usefulness is limited because of its low spatial resolution. A interesting perspective: fusion

• Coastal waters:

Water quality, benthic communities, bathymetry, bottom types, seeground interface...

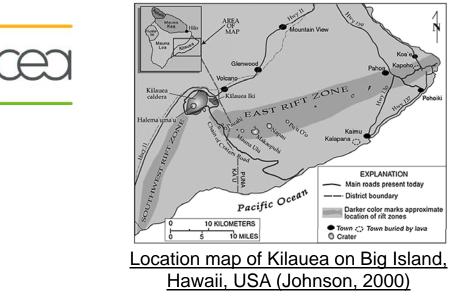
Quinault biomass burning aerosol plume analysis by hyperspectral imagery*

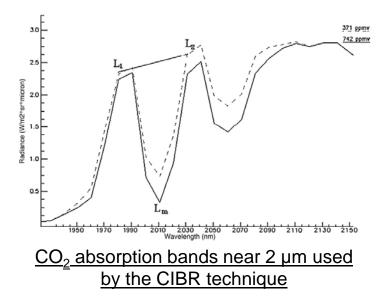


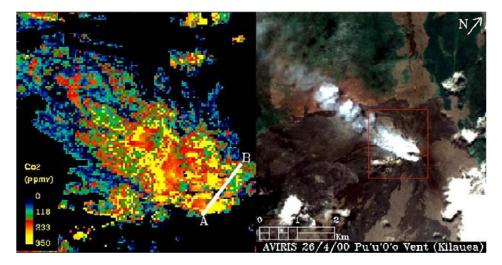
In situ measured parameters (samples , lidar) and estimation from the image (concentration τ_{550} , composition BC, size distribution r_{eff}) are in adequacy

* A. Alakian, R. Marion, and X. Briottet, "Remote sensing of aerosol plumes: a semianalytical model," *Applied Optics*, Vol. 47, No. 11, pp. 1851-1866, 10 April 2008

Carbon dioxide of Pu`u`O`o volcanic plume at Kilauea retrieved by AVIRIS hyperspectral data*



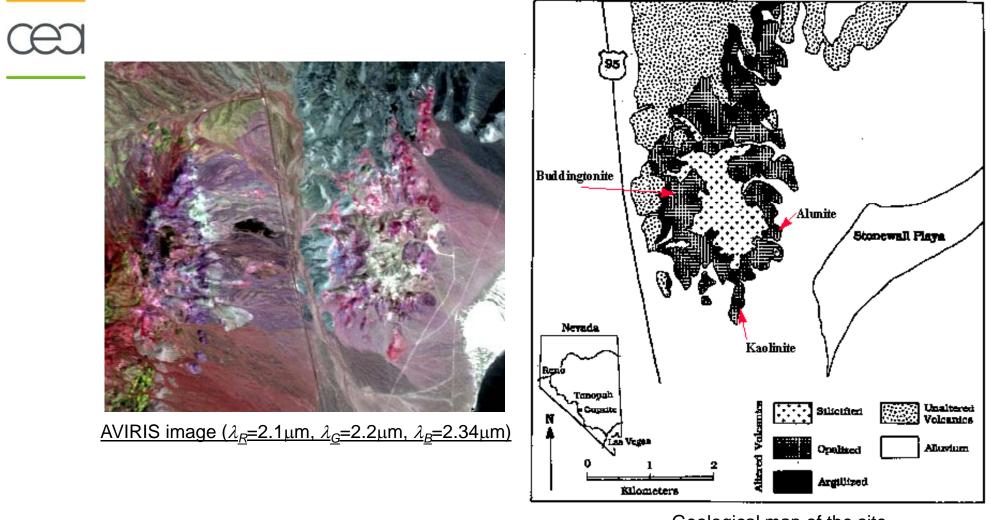




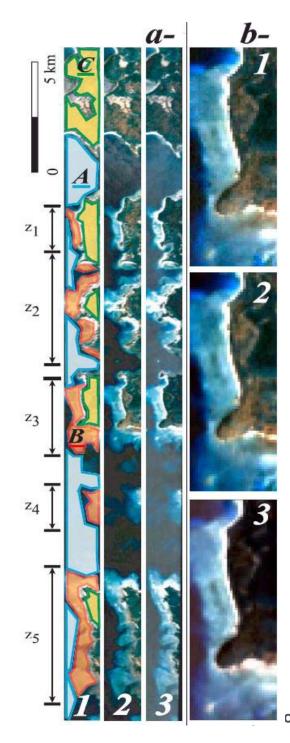
Map of volcanic plume carbon dioxide (left) and AVIRIS image of Pu`u `O`o Vent plume (right)

* C. Spinetti, V. Carrère, M. F. Buongiorno, A. J. Sutton, T. Elias, "Carbon dioxide of Pu`u`O`o volcanic plume at Kilauea retrieved by Style in the special data," Freih and Sensing of the Environmente, Volland, pp. 319213499,020089

Geology (mineral mapping at Cuprite) [1/3]



Geological map of the site



Technique

Table 1. Hyperion Hyperspectral Images and Sea Level Used in the Present Study^a

200 		Image Reference	Date	Site Longitude	Site Latitude	Sea Level (m)
1	PRE	EO11340512004049110PY	18 February 2004	92.8842	13.3422	1.42
2	PRE	EO11340512004065110PY	5 March 2004	92.8842	13.3422	1.64
3	POST	EO11340512005042110KZ	11 February 2005	92.8000	12.5000	1.90

^aImages acquired to minimize effects of seasonal variations on radiative transfer. Theoretical sea level from tide gauge at Port Blair, South Andaman Islands (http://www.shom.fr/).

$$R = R_g + R_w + (R_b - R_w)e^{-2Kz}$$

R= surface reflectance

Rg =sea surface contribution

 Rw = deep water subsurface scattering contribution of photons that did not reach the sea bottom

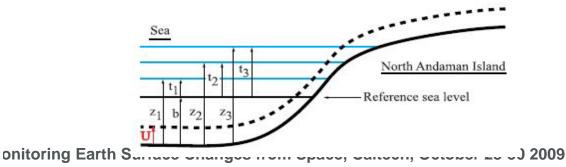
z = depth of the water column

K =effective attenuation coefficient (upwelling and downwelling)

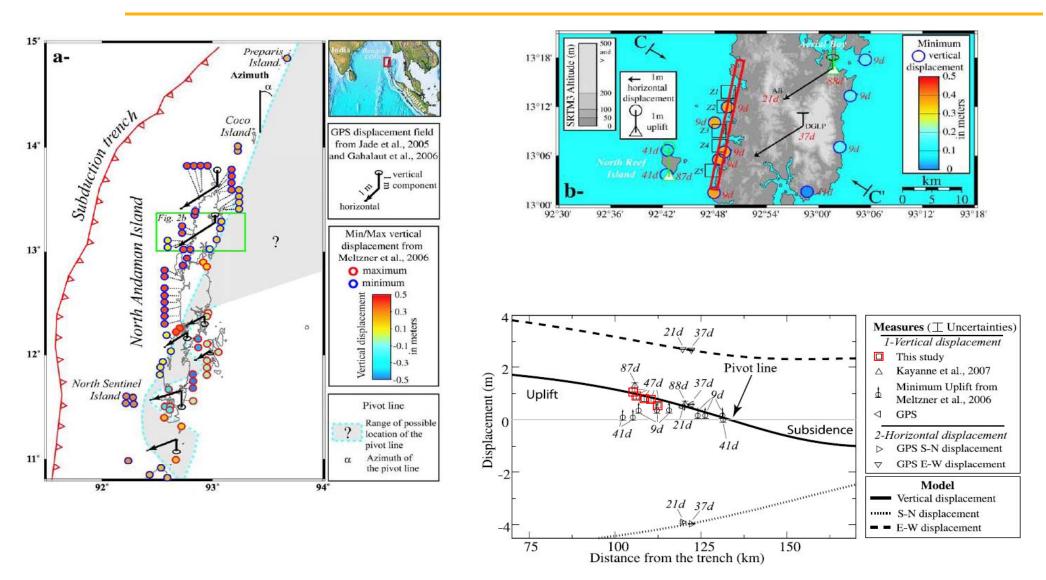
$$K = \frac{1}{2(z_1 - z_2)} \log \left[\frac{R_2 - (R_g + R_w)_2}{R_1 - (R_g + R_w)_1} \right]$$
$$= \frac{1}{2(t_1 - t_2)} \log \left[\frac{R_2 - (R_g + R_w)_2}{R_1 - (R_g + R_w)_1} \right]$$

$$U = t_3 - t_1 + \frac{1}{2K} \log \left[\frac{R_3 - (R_g + R_w)_3}{R_1 - (R_g + R_w)_1} \right].$$

U= tectonic uplift estimated for each wavelength and then averaged over the spectral range (570–690 nm)



Comparison with other techniques



Smet, S., R. Michel, and L. Bollinger (2008), "Uplift of the 2004 Sumatra-Andaman earthquake measured from differential hyperspective in the second states of the second states o