

# Geodetic tools for characterizing hillslope processes and landslides:

## 5 goals

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## Goal 1. Interpretation and prediction of hillslope evolution



Landscape evolution reflects:

- tectonic forcing
- rock type
- climate & biology

$$\frac{\partial z}{\partial t} = -\nabla \cdot q_s + U$$

The equation is presented on a yellow background. A blue arrow labeled "Erosion" points to the term  $-\nabla \cdot q_s$ . A red arrow labeled "Tectonics" points to the term  $+U$ .

Dietrich et al., 2003

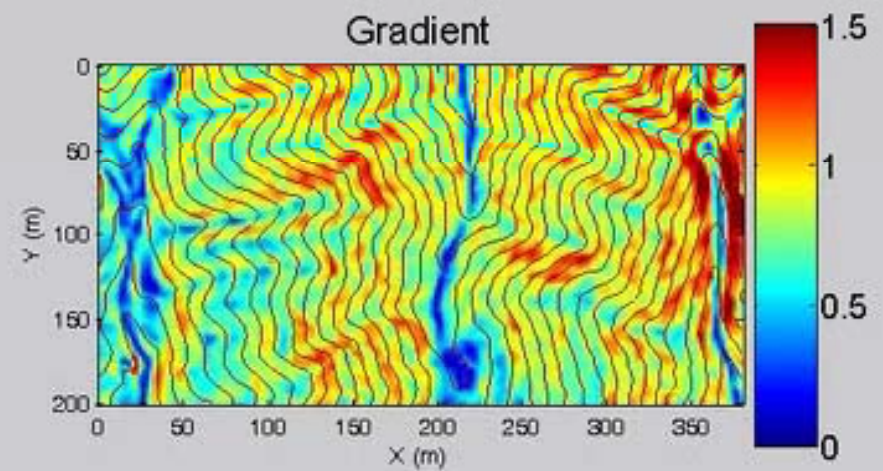
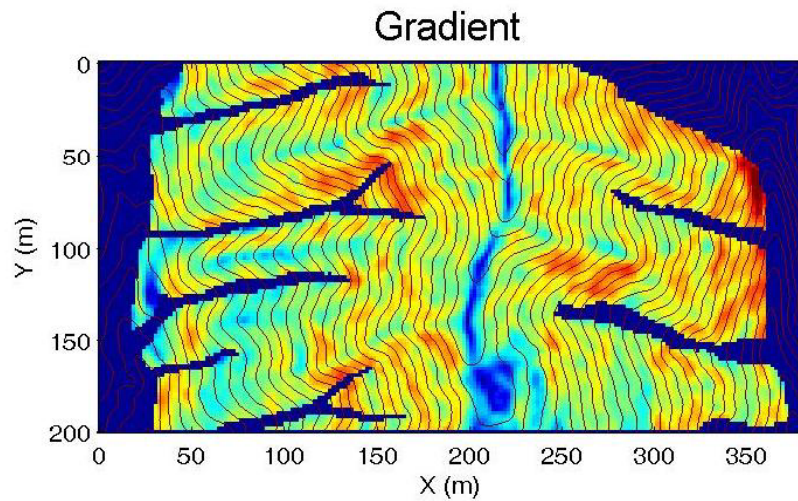
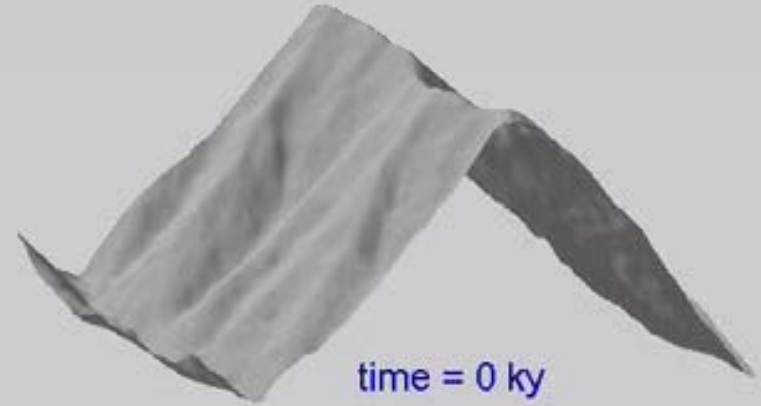
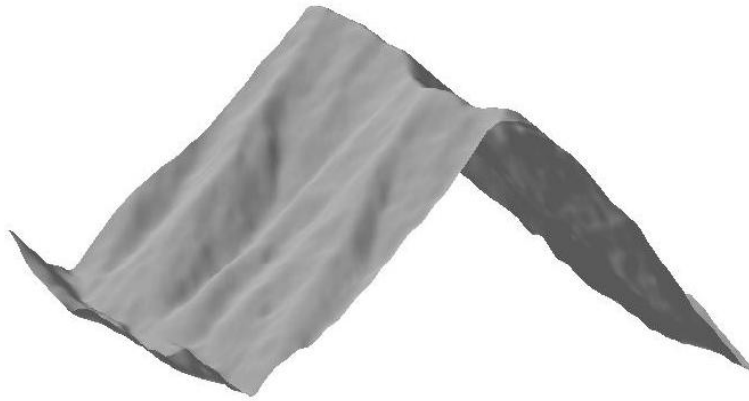
- Most landscapes erode at rates less than  $0.5 \text{ mm yr}^{-1}$
- Most erosion rate estimates are derived from river sediments (i.e., catchment-averaged)
- To test/calibrate erosion models, we often rely on landscape morphology (e.g., lidar)

# Morphologic change after 500,000 yrs...

$$q_s = KS$$

S = local gradient

Initial conditions (current via lidar)



Roering, 2008, *GSA Bulletin*

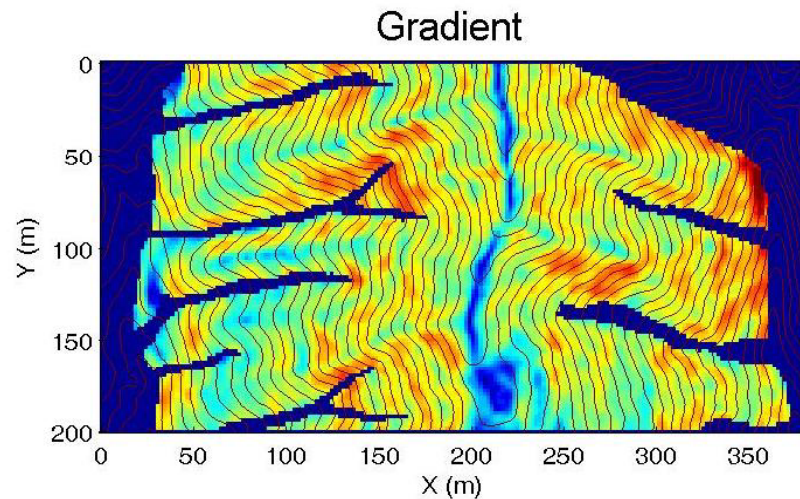
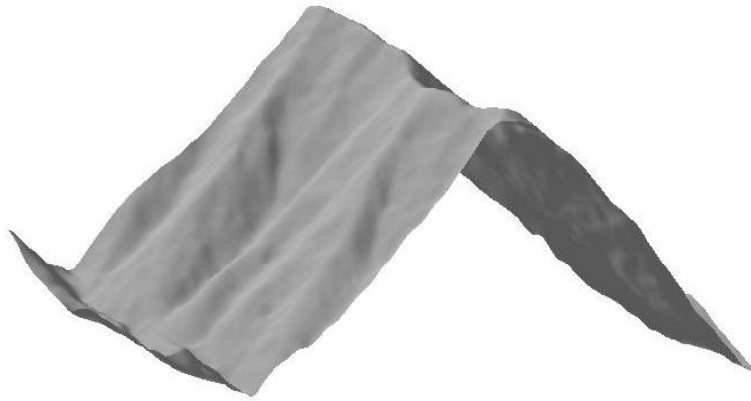
## Morphologic change after 500,000 yrs...

- 1) Steady-state hillslopes?
- 2) Continuum approximation of transport?

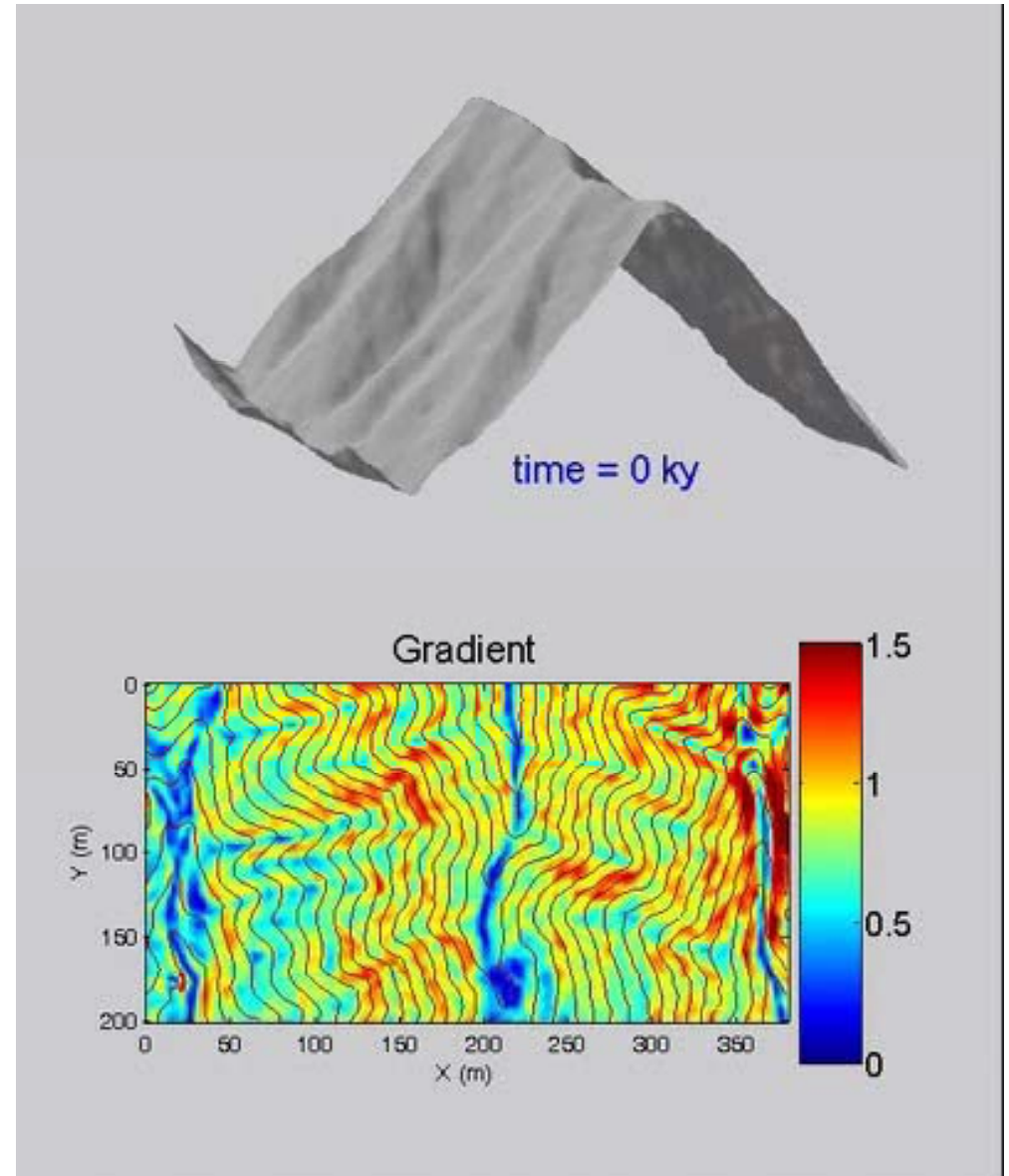
$$q_s = \frac{K(h)S}{1 - (S/S_c)^2}$$

$h$  = soil depth (m)  
 $S$  = local gradient

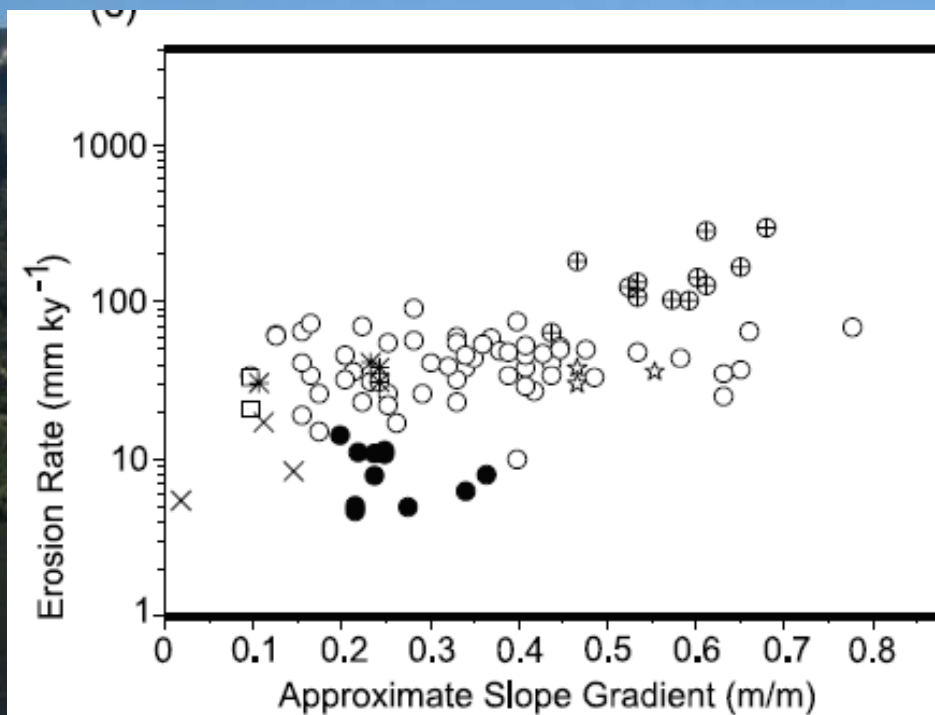
Initial conditions (current via lidar)



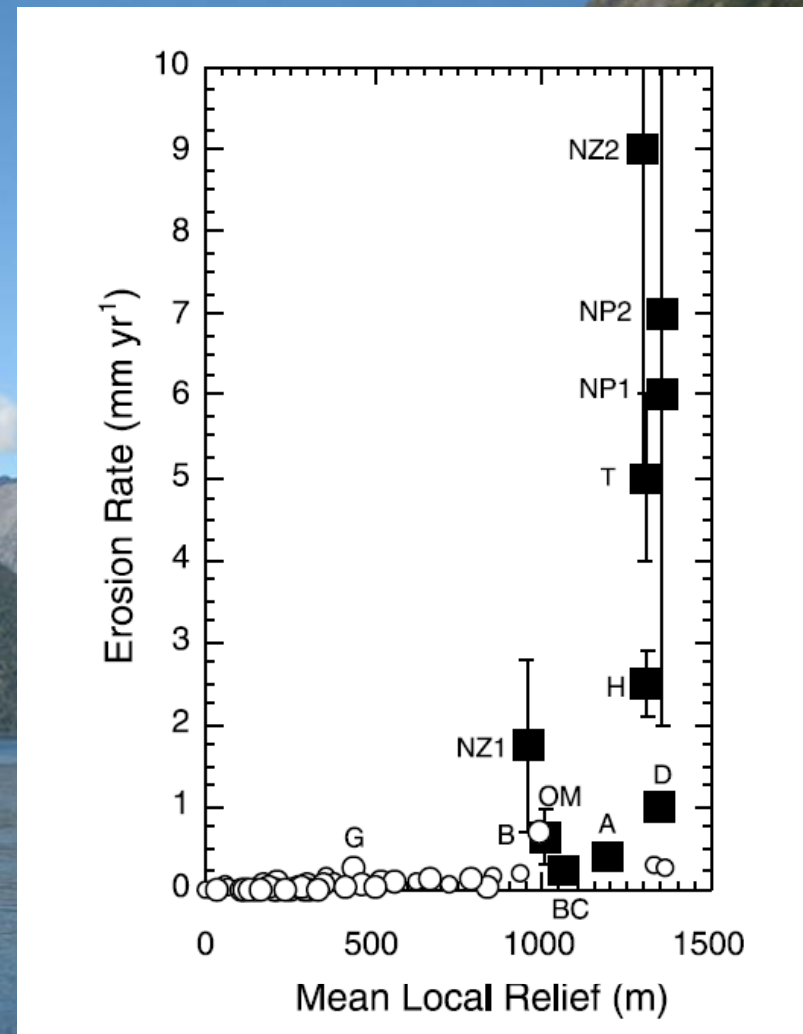
Roering, 2008, *GSA Bulletin*



## Goal 2: Infer rock uplift and erosion from morphology



Von Blanckenburg et al., 2004

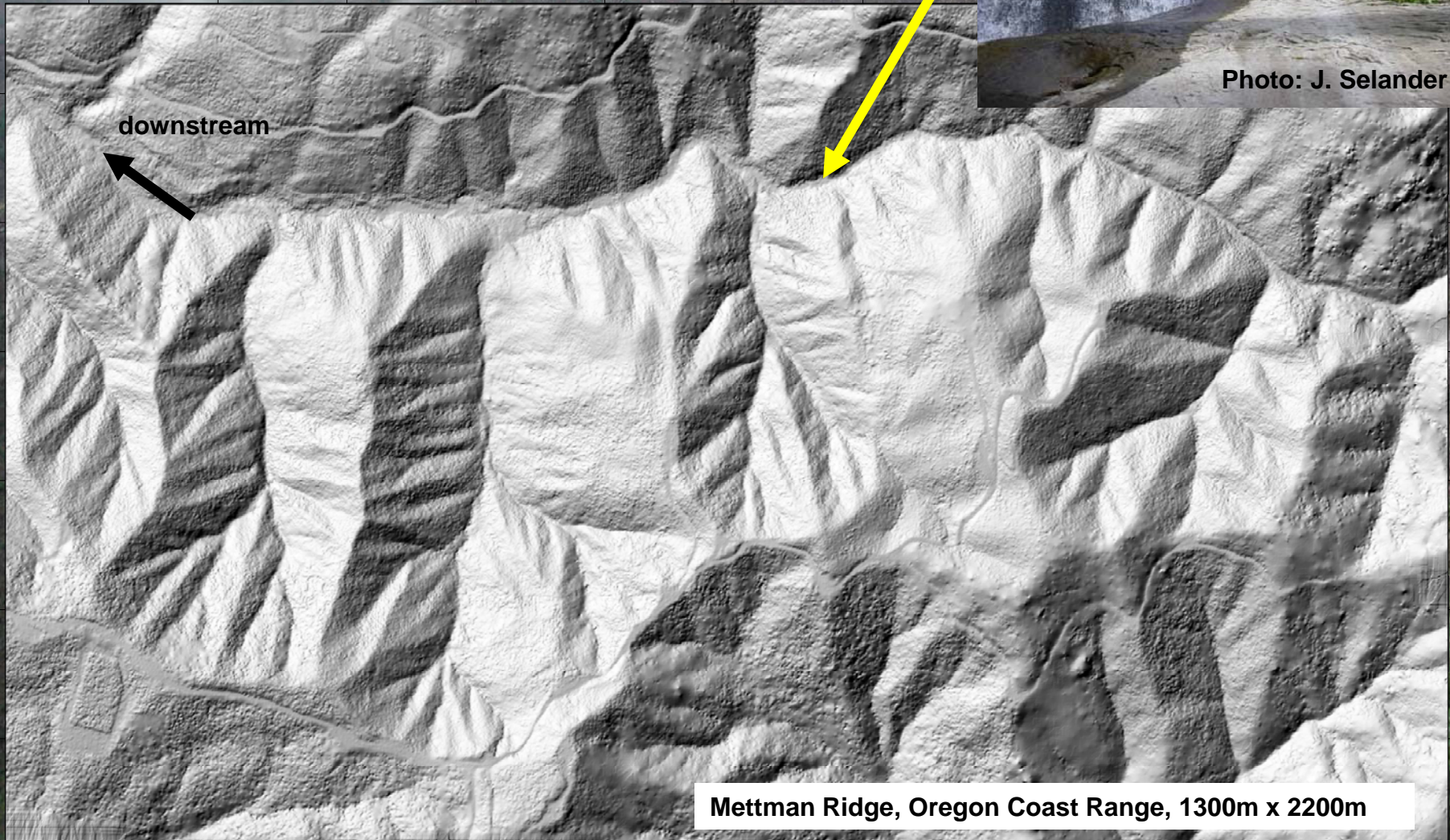


Montgomery & Brandon, 2002

**Process-scale investigations:  
Knickpoint evolution in rivers and landscape  
response**

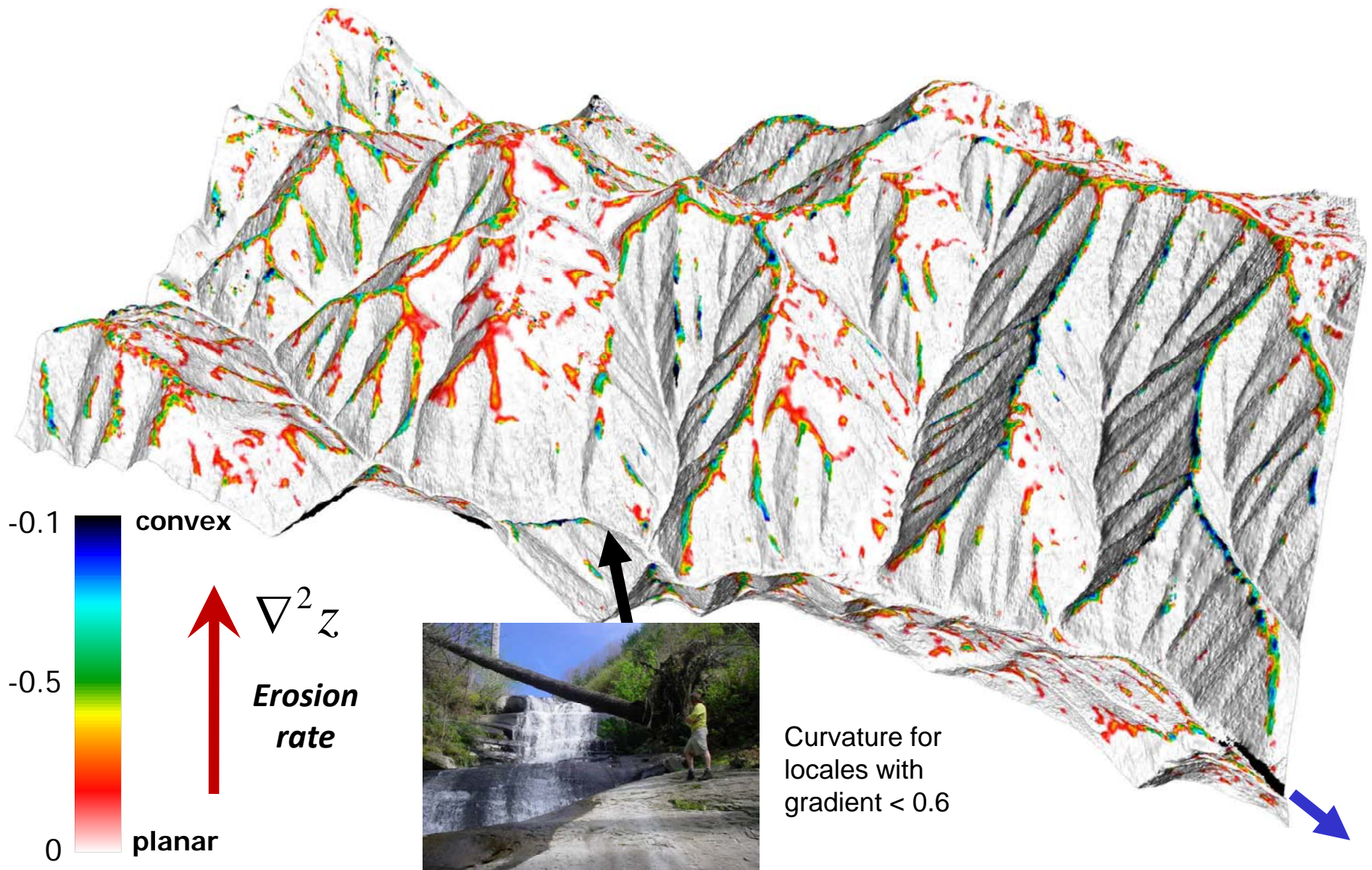


Photo: J. Selander



Mettman Ridge, Oregon Coast Range, 1300m x 2200m

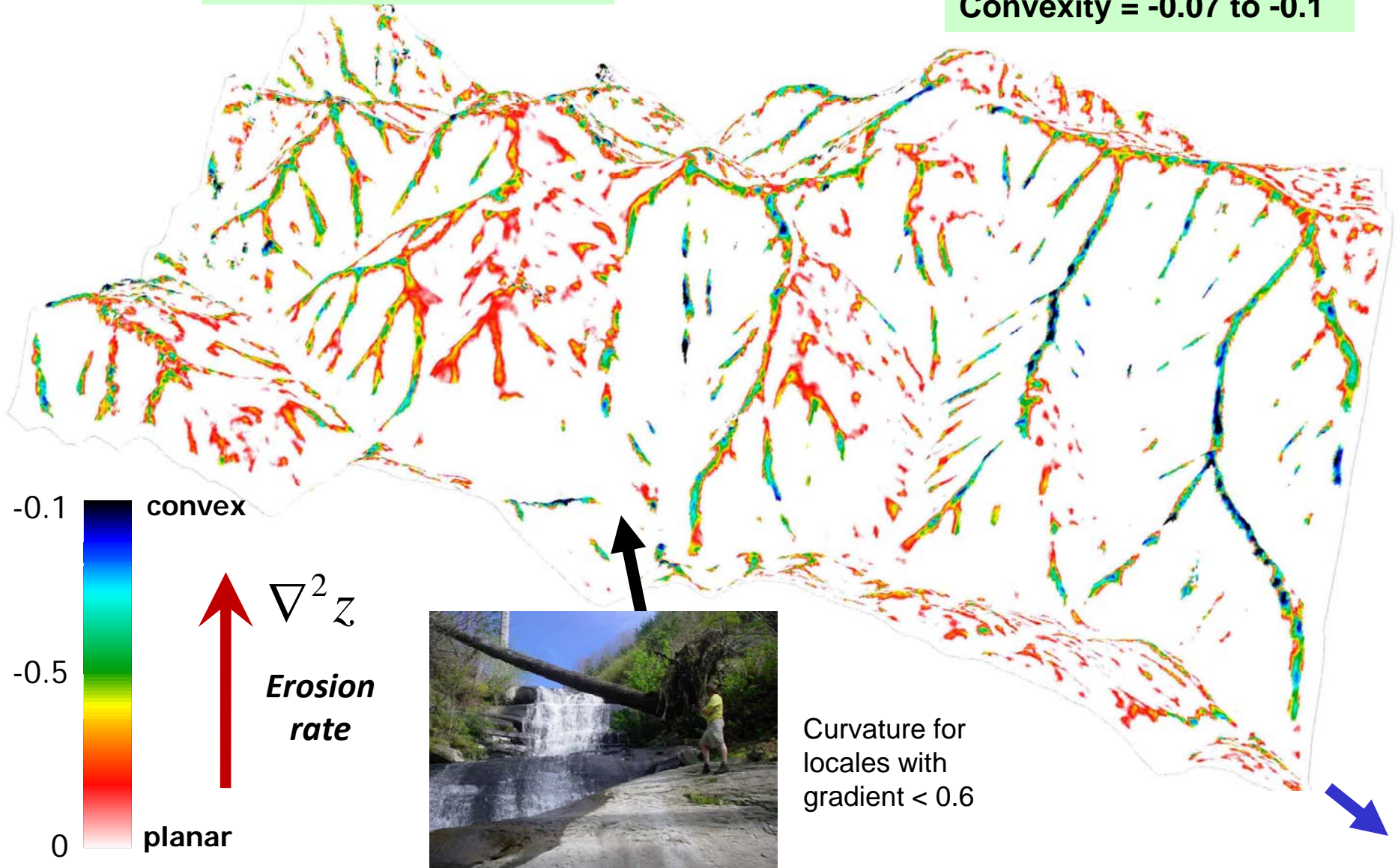
# Spatial variation of hilltop convexity and landscape adjustment



# Spatial variation of hilltop convexity and landscape adjustment

Upstream of Knickpoint:  
**Convexity = -0.02 to -0.06**

Downstream of Knickpoint:  
**Convexity = -0.07 to -0.1**





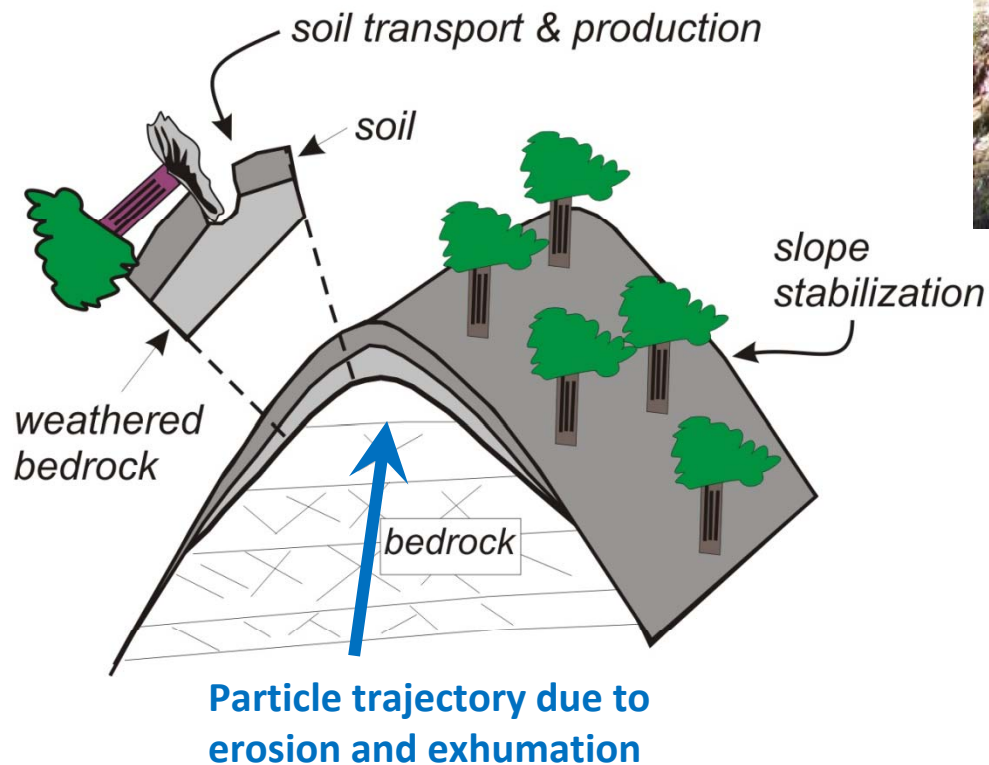
### Goal 3: Measurement of stochastic hillslope processes

What is the magnitude-frequency relationship of disturbances that drive soil production and transport?



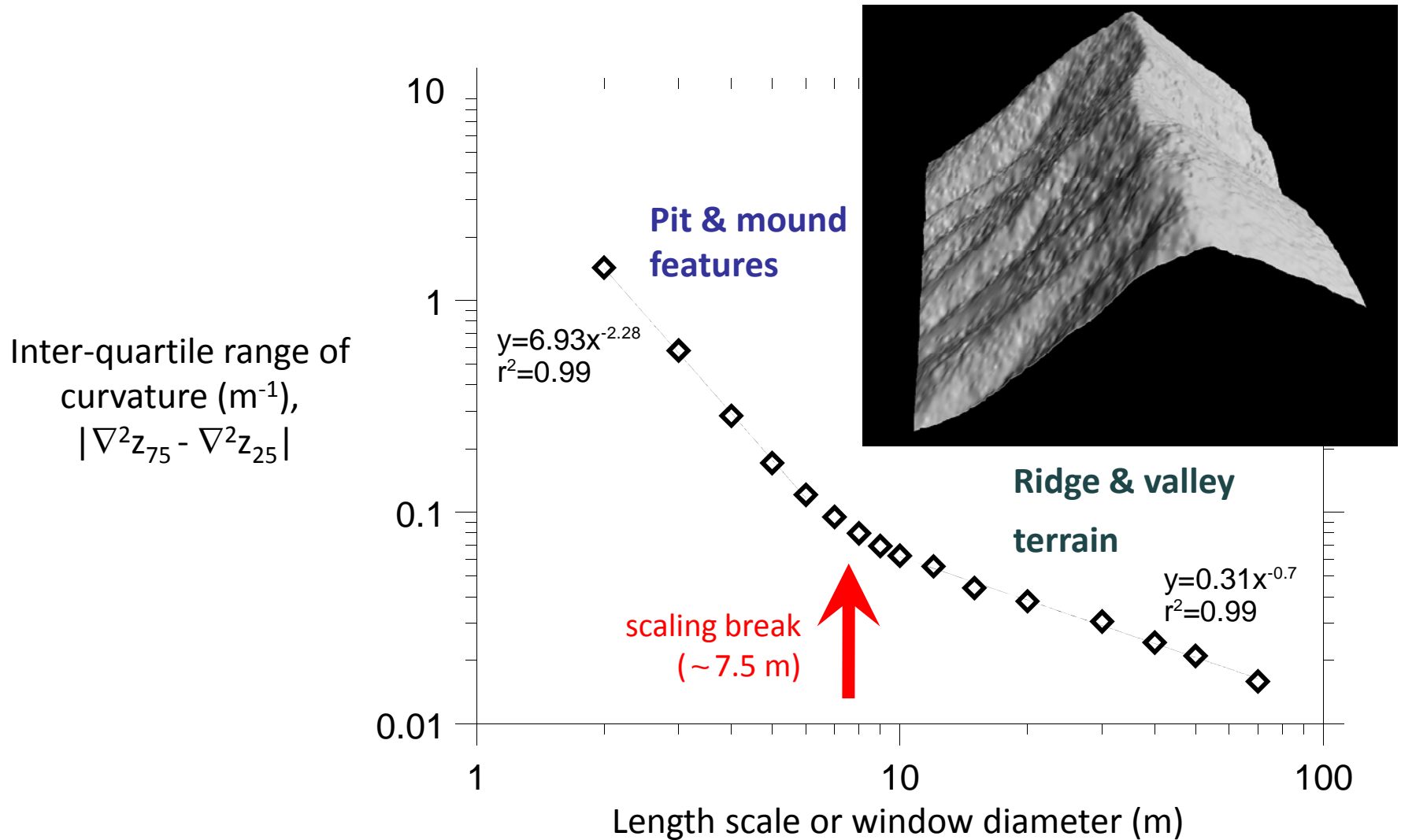
## Goal 4: The evolution of the critical zone (top of canopy to base of weathering front...30m?)

- What is the biotic role?
- How do topography and the critical zone co-evolve?

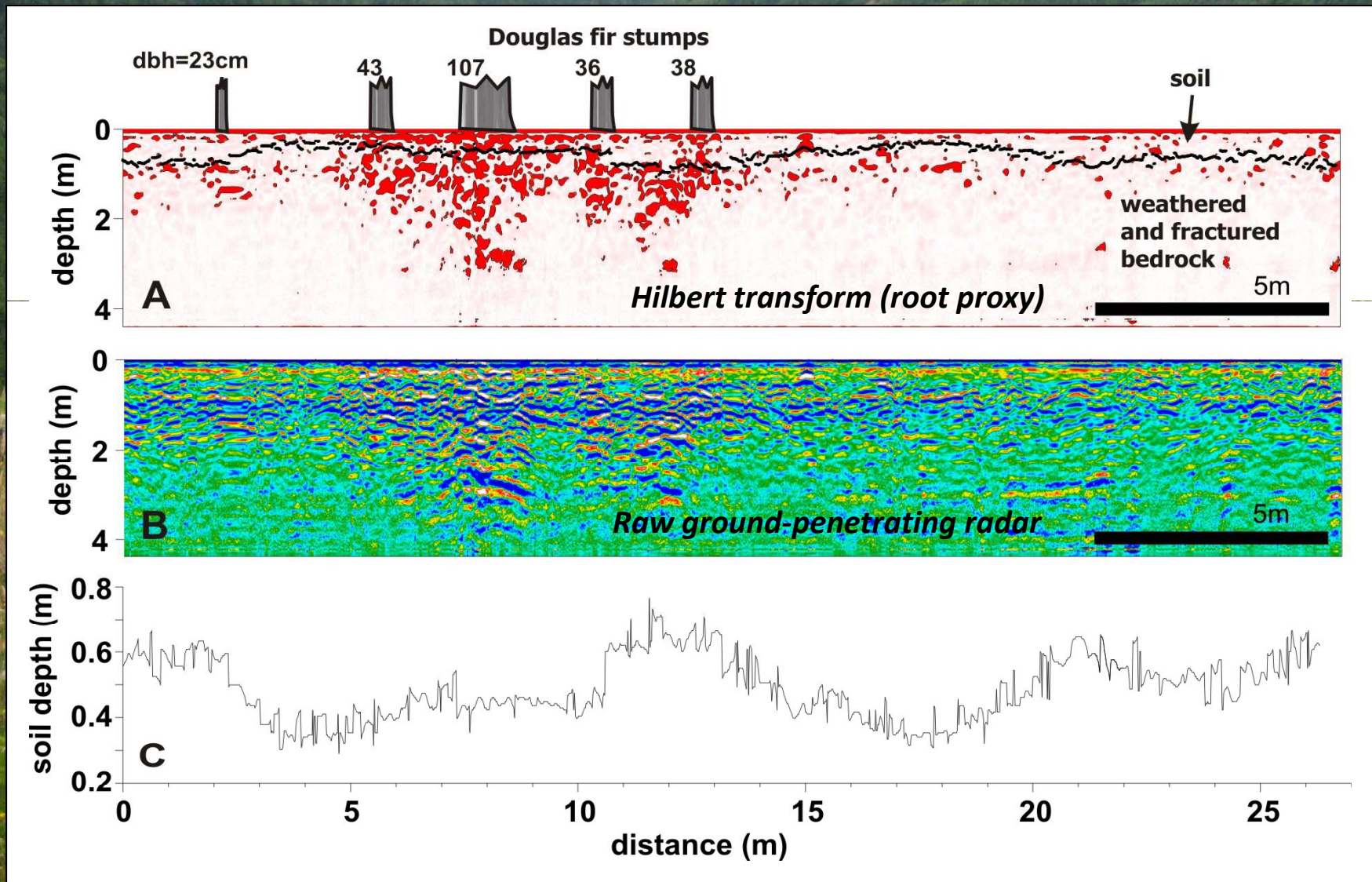


## Biotic signature from airborne lidar

- The land surface becomes increasingly rough at short length scales
- **Pit and mound features** generated by tree turnover dominate small length scales

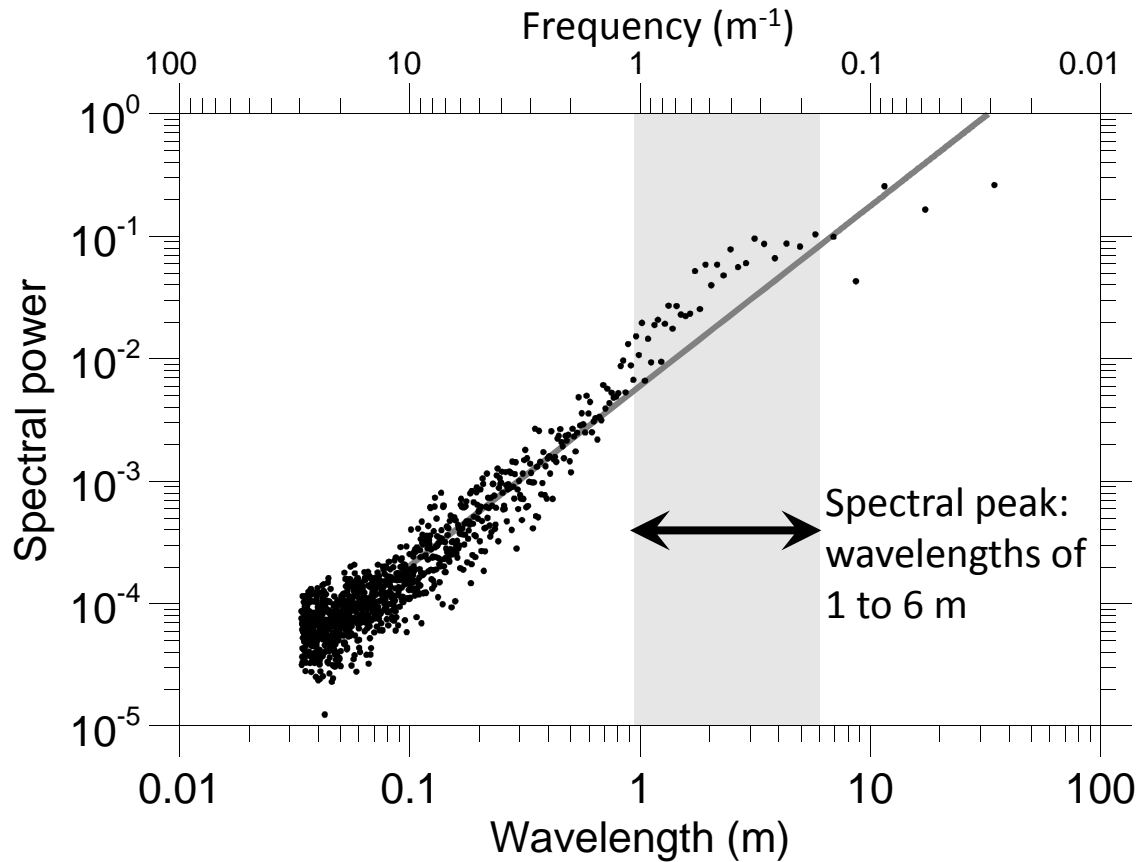
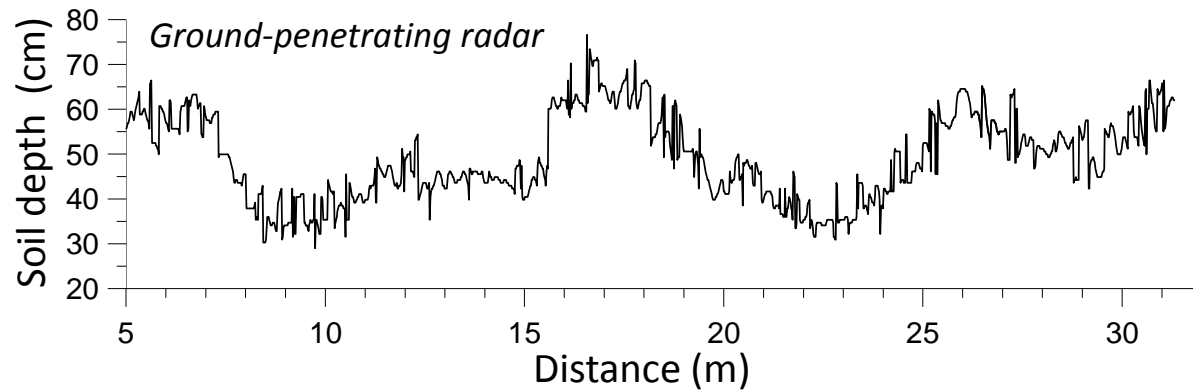


# Ground-penetrating radar for mapping soil depth and root penetration into bedrock



**Is there a characteristic scale of soil depth variations?**

Fourier transformation and spectra averaging of 6 transects



- Characteristic scale = 1 to 6m
- Subsurface reflection of biotic processes (tree root action)

## Goal 5: Landslides and landscape evolution

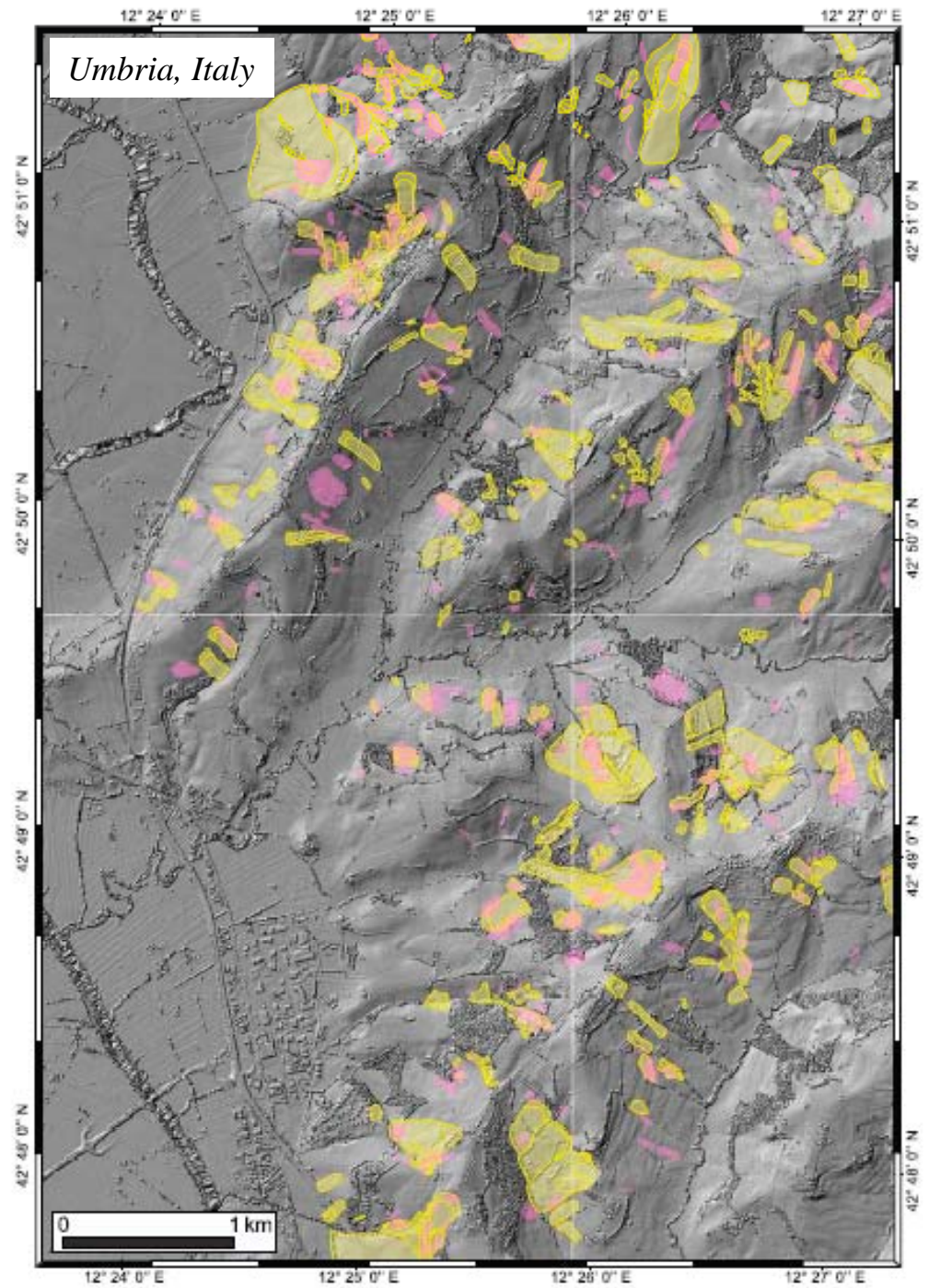
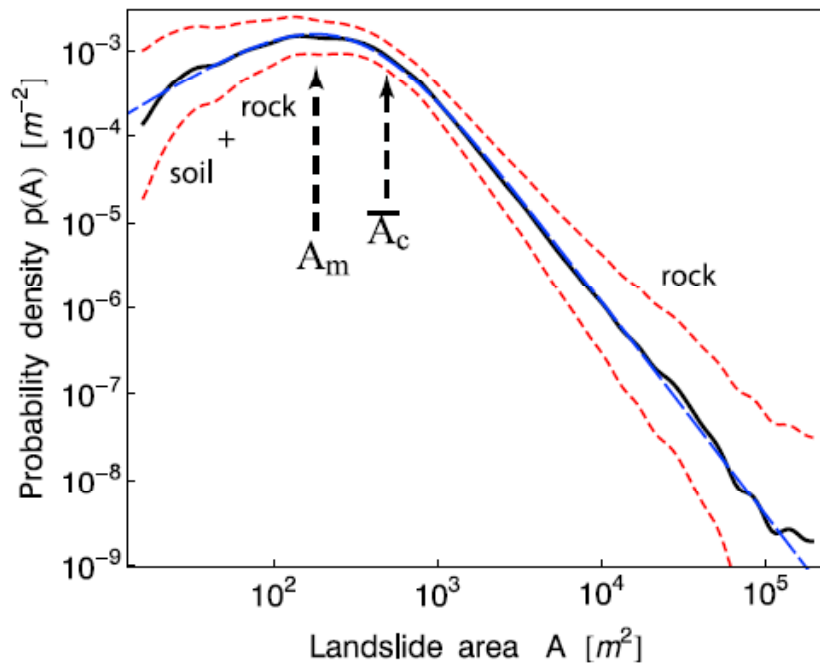
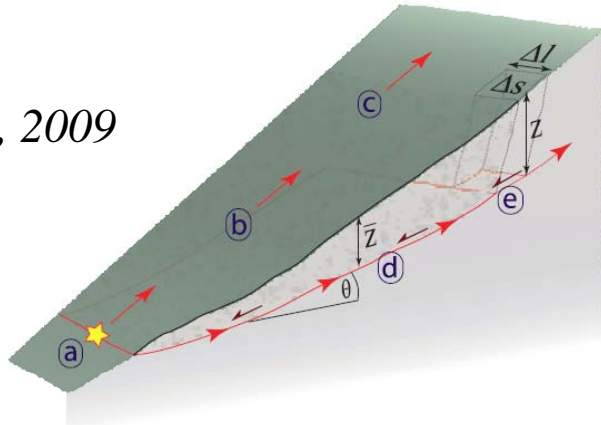
1) What controls the size of landslides and their contribution to landscape evolution?

2) Why do some slides fail catastrophically and others deform slowly with seasonality?



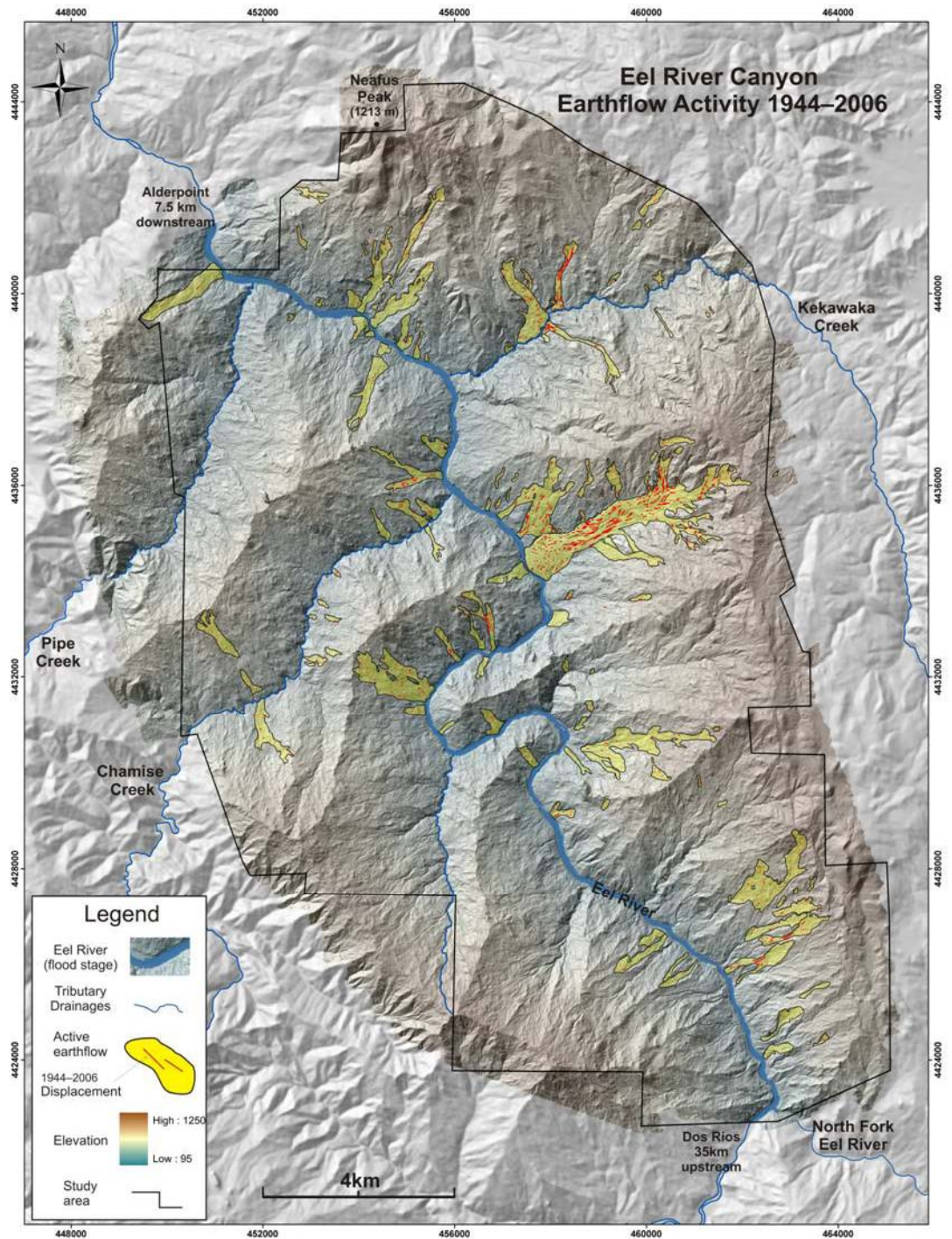
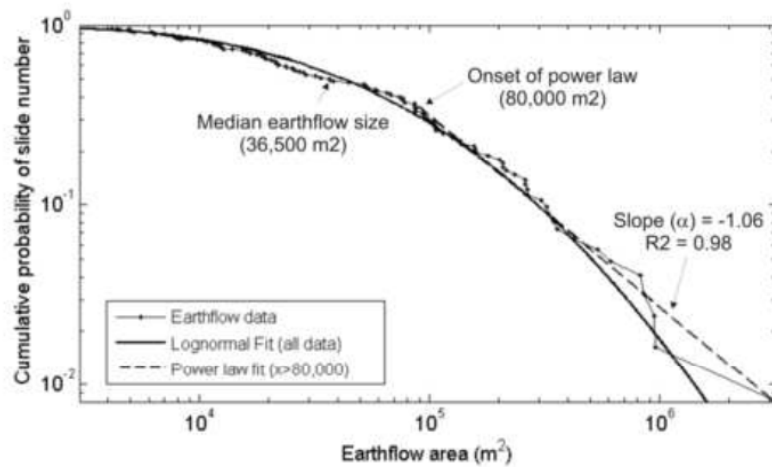
# Landslide mapping inventories: Do landslide statistics yield mechanistic information?

*Stark &  
Guzzetti, 2009*



## Eel River (*Mackey, PhD. 2009*)

- LiDAR and historical air photos
- 122 active earthflow features (7.3% of study area)
- Earthflow sediment yield to channels:  $0.53 \text{ mm yr}^{-1}$
- Denudation from suspended sediment records:  $\sim 0.9 \text{ mm yr}^{-1}$





InSAR using ALOS in Northern California (Roering et al., 2009, GRL)

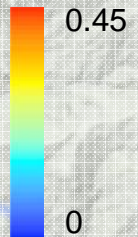


**North Fork Eel River**

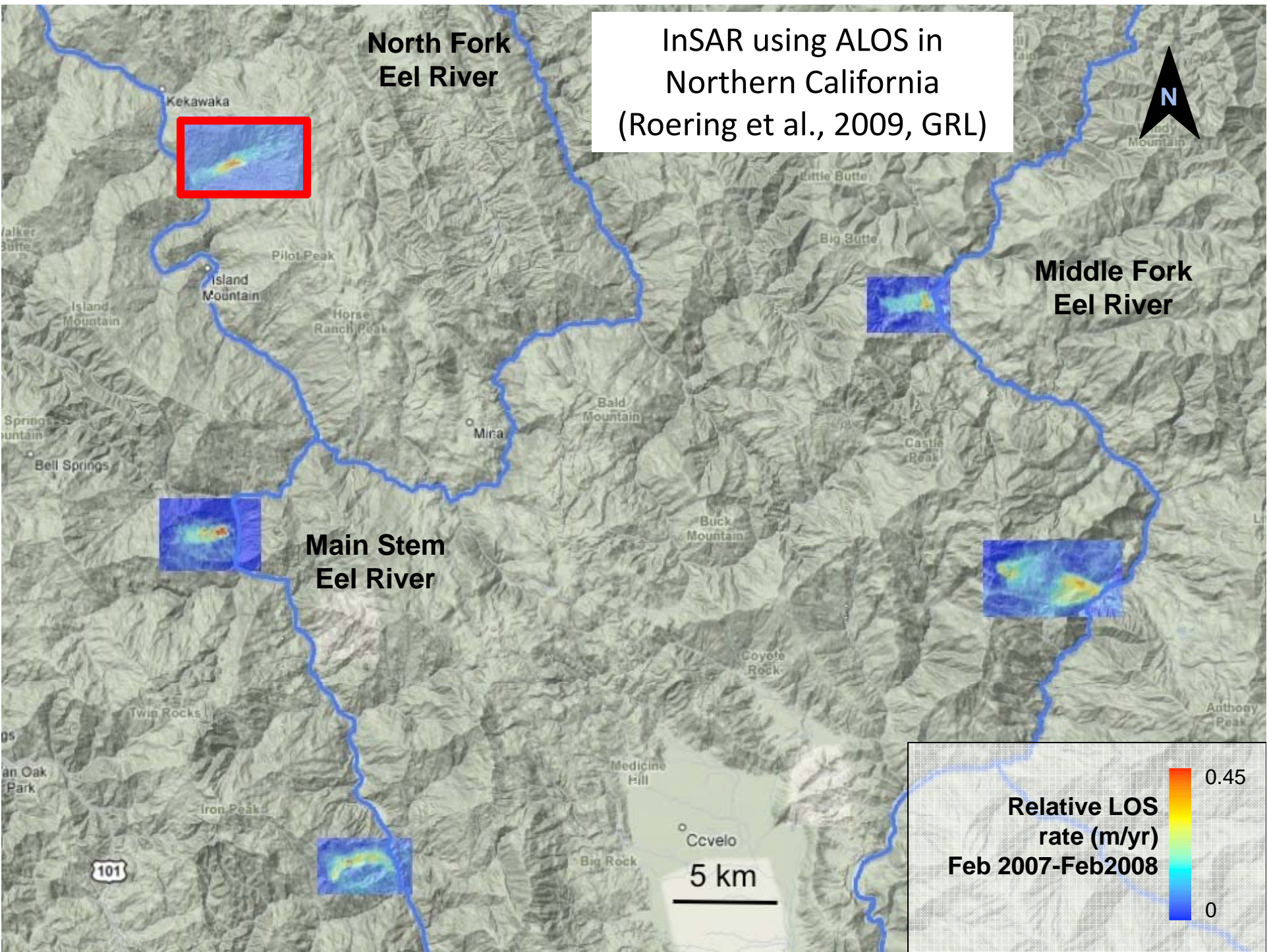
**Middle Fork Eel River**

**Main Stem Eel River**

Relative LOS rate (m/yr)  
Feb 2007-Feb2008

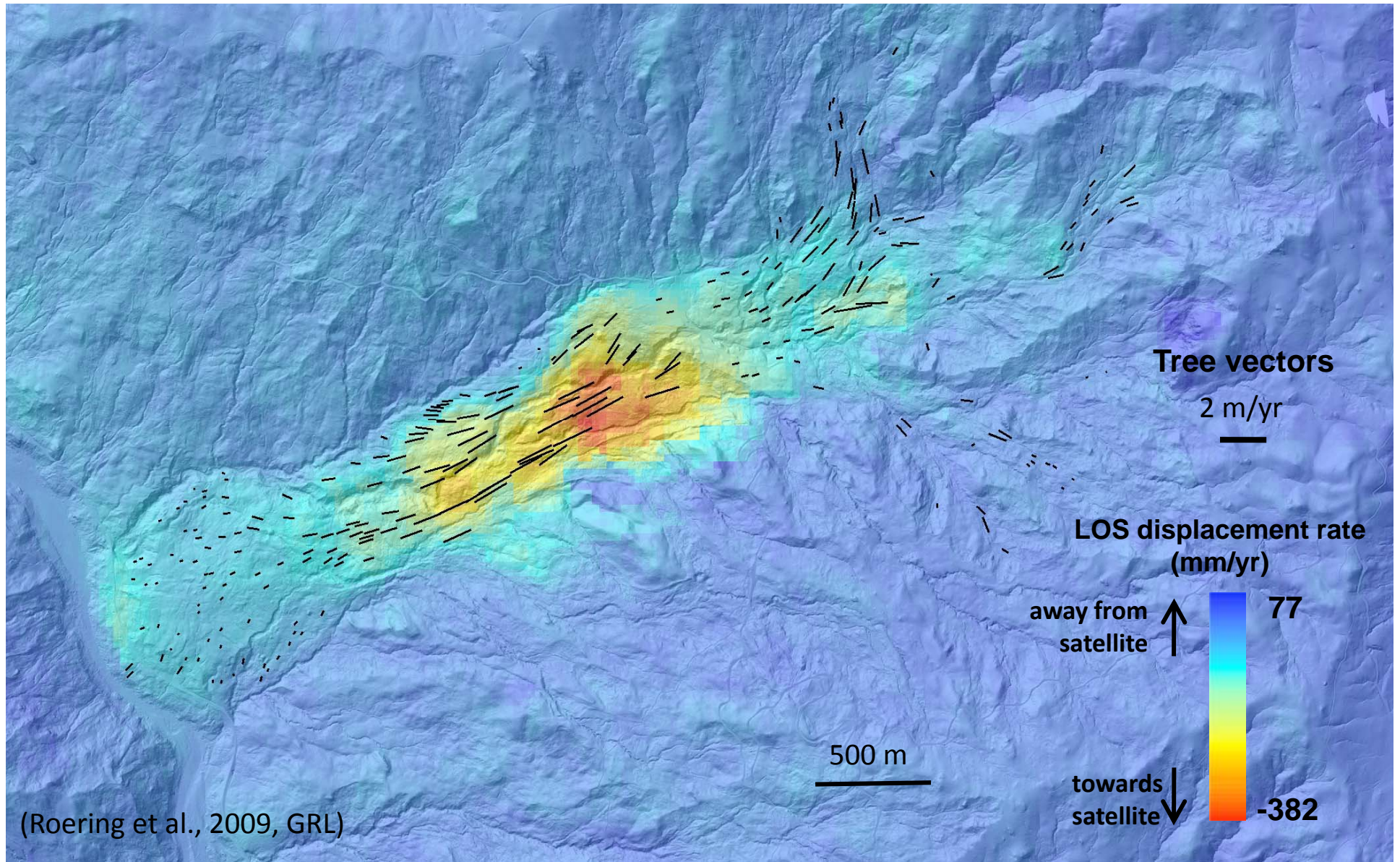


5 km



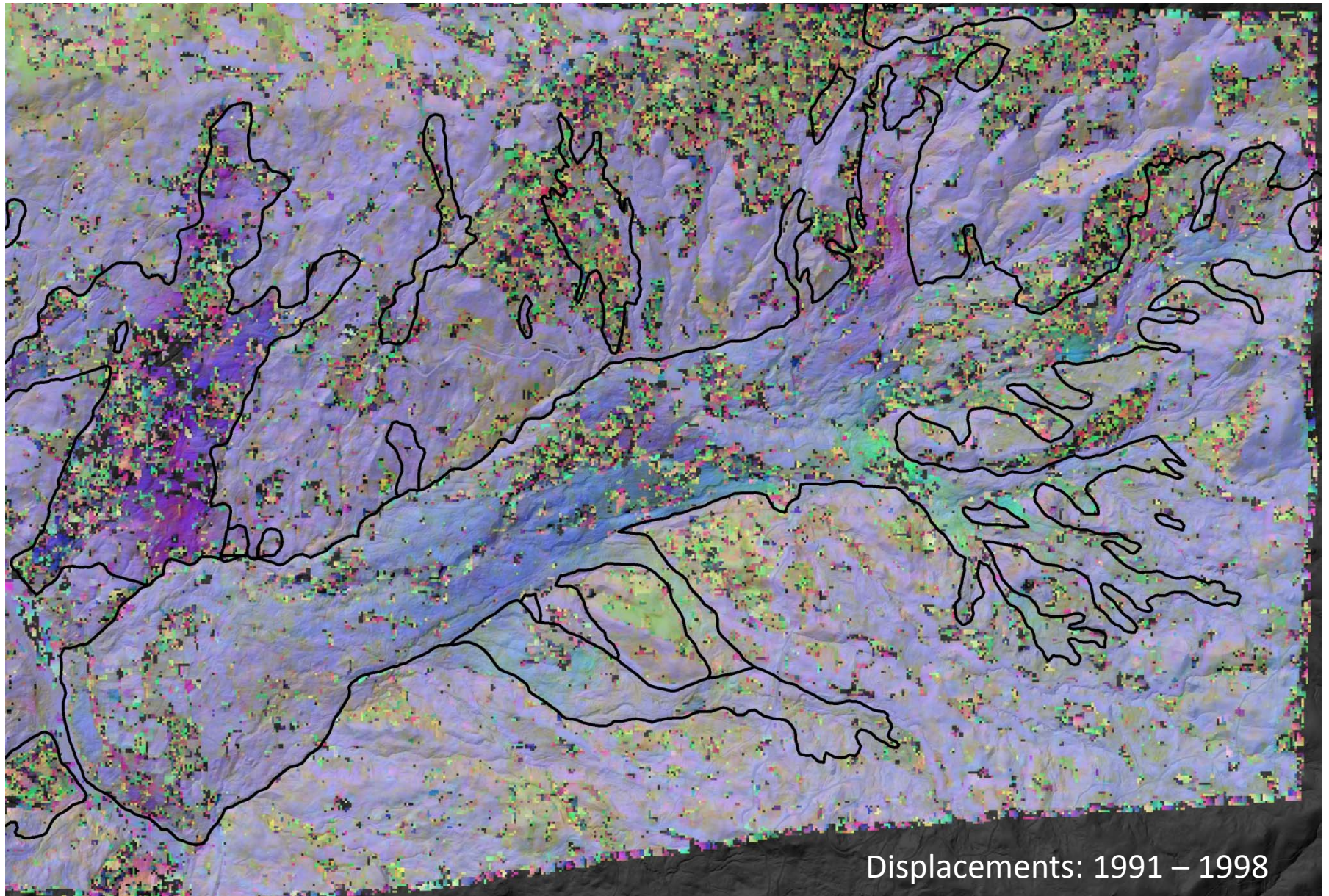
## InSAR and photo-derived displacements

- InSAR (stack of 17 infs): Feb 13, 2007 – Feb 16, 2008
- Tree vectors (Mackey, 2009): air photos 1964 – 2006
- InSAR velocities (horizontally projected) are 20% slower
- Satellite orientation relative to terrain and slope deformation



Dynamics using automated photo rectification,  
coregistration, and subpixel correlation

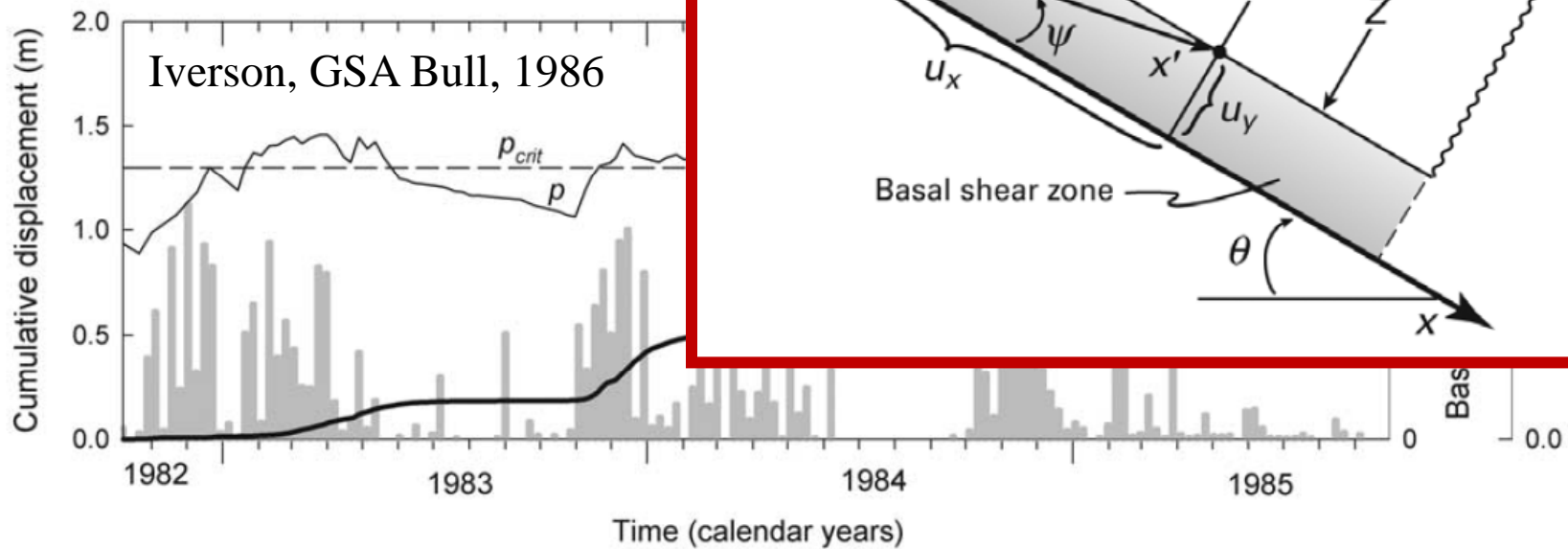
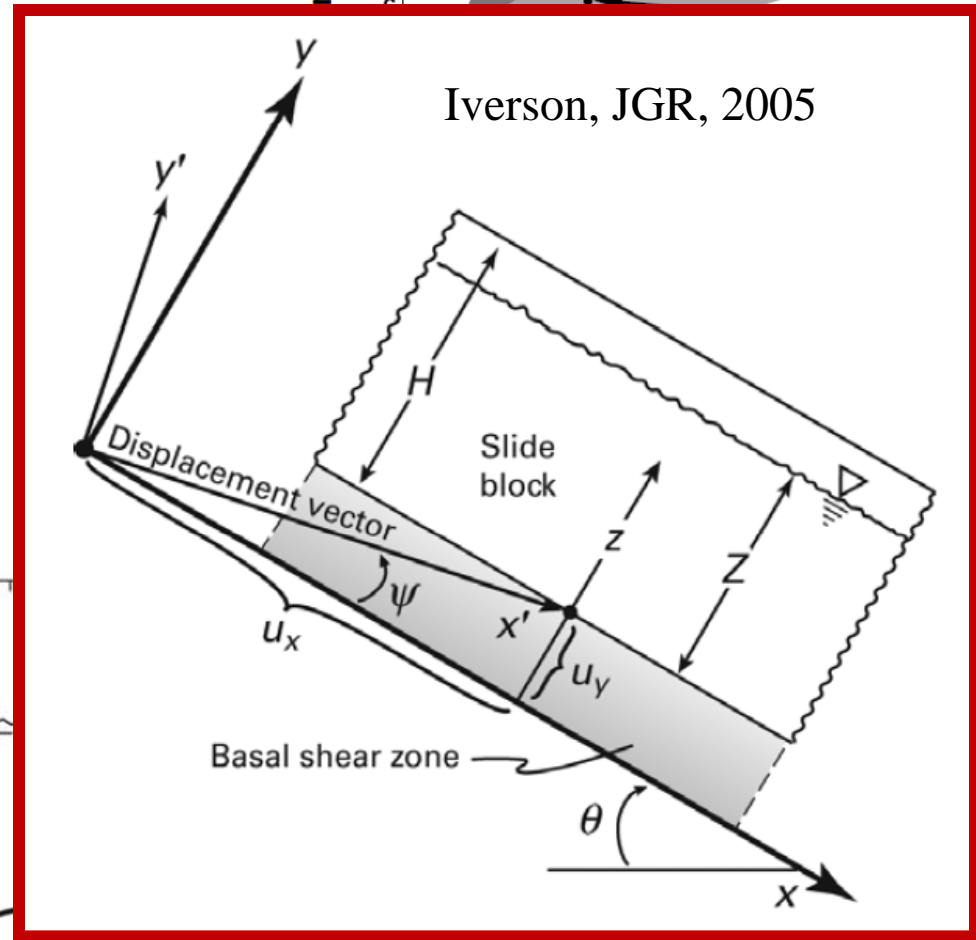
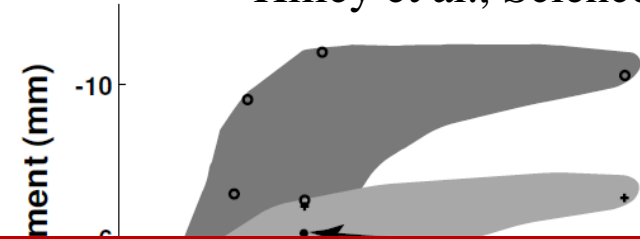
*COSI-Corr (Leprince et al., Eos, 2008)*



**Pervasive slow-moving slides self-regulate and do not fail catastrophically**

Shear-zone dilatancy may permit negative pore pressure-shear feedbacks and thus allow for slow, steady motion

- Can we image this feedback?
- Is there a limit to shear zone dilatancy?

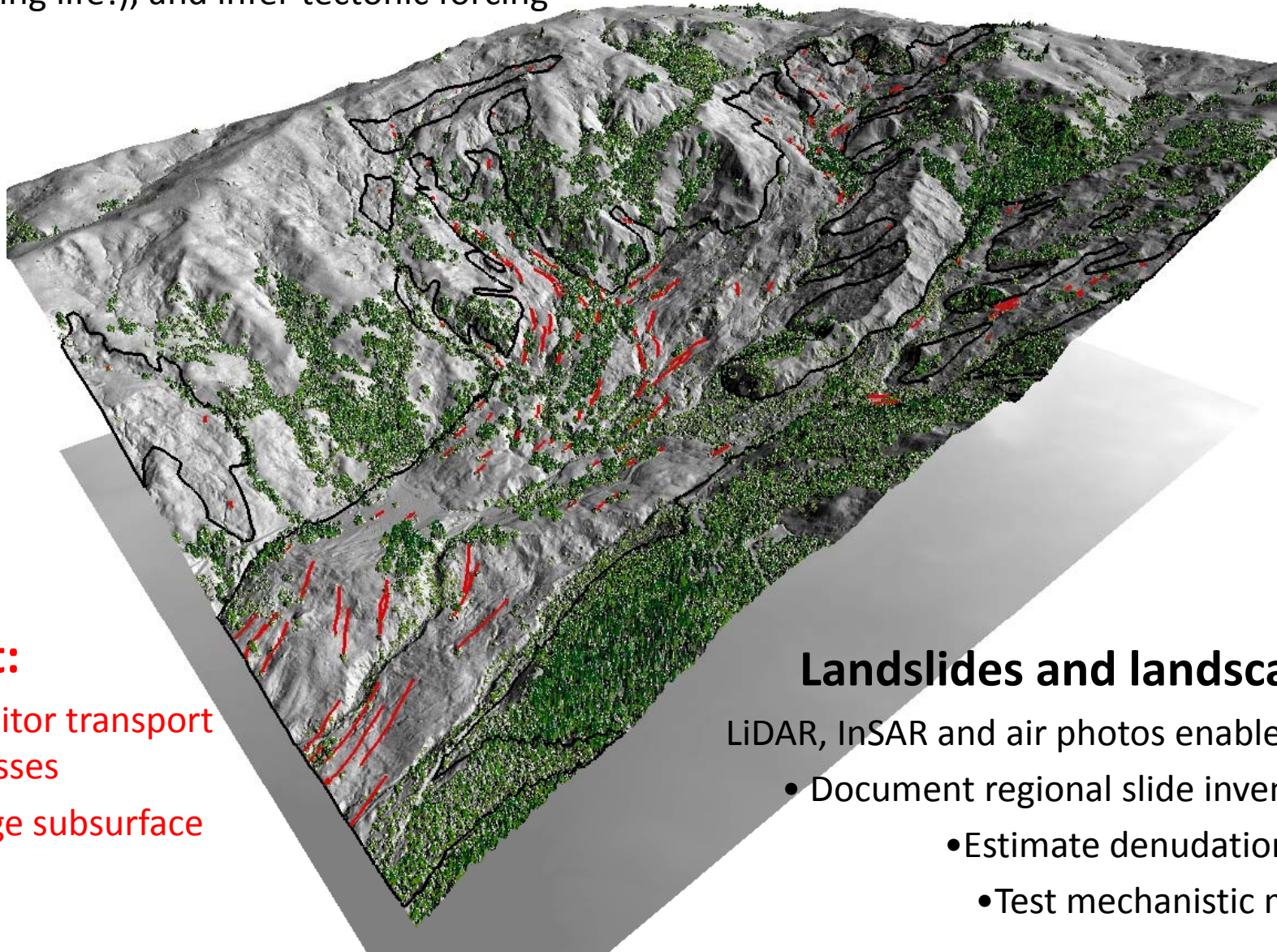


## Hillslopes:

Airborne lidar and radar enables us to:

Test models and make predictions, quantify process signatures (including life!), and infer tectonic forcing

## Summary



## Next:

- Monitor transport processes
- Image subsurface

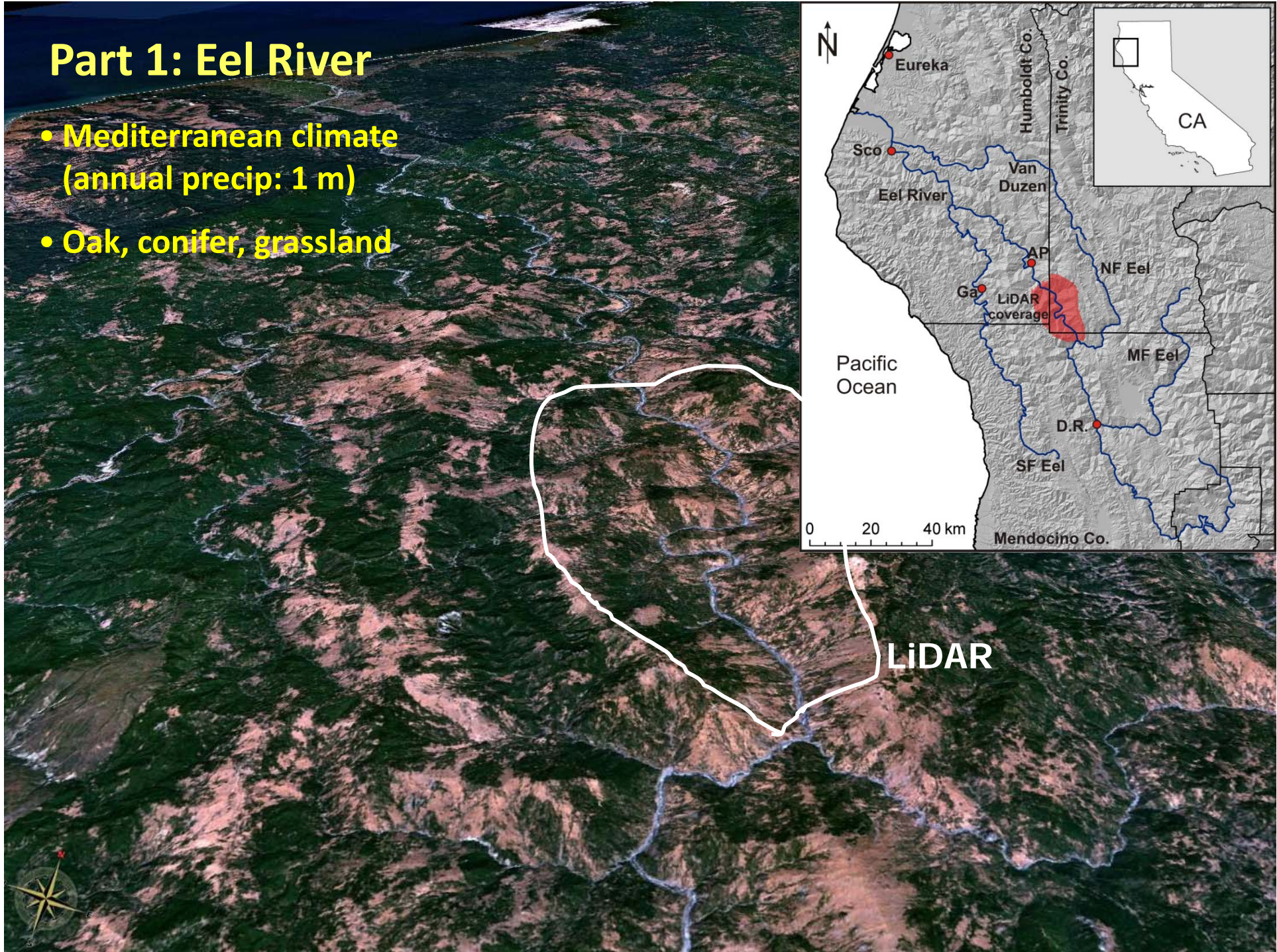
## Landslides and landscapes:

LiDAR, InSAR and air photos enable us to:

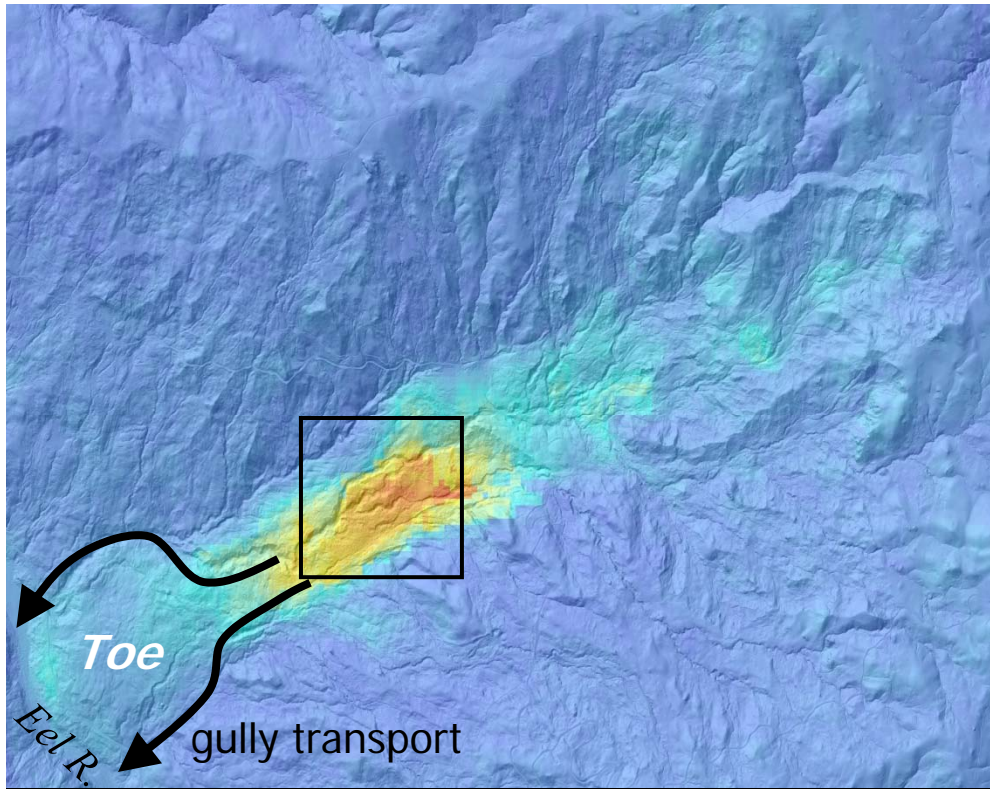
- Document regional slide inventories
- Estimate denudation rates
- Test mechanistic models

# Part 1: Eel River

- Mediterranean climate (annual precip: 1 m)
- Oak, conifer, grassland

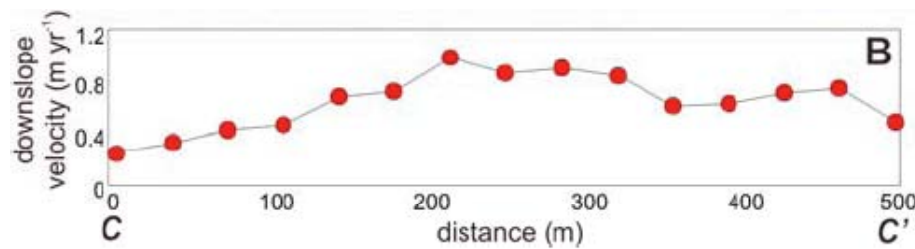


LiDAR

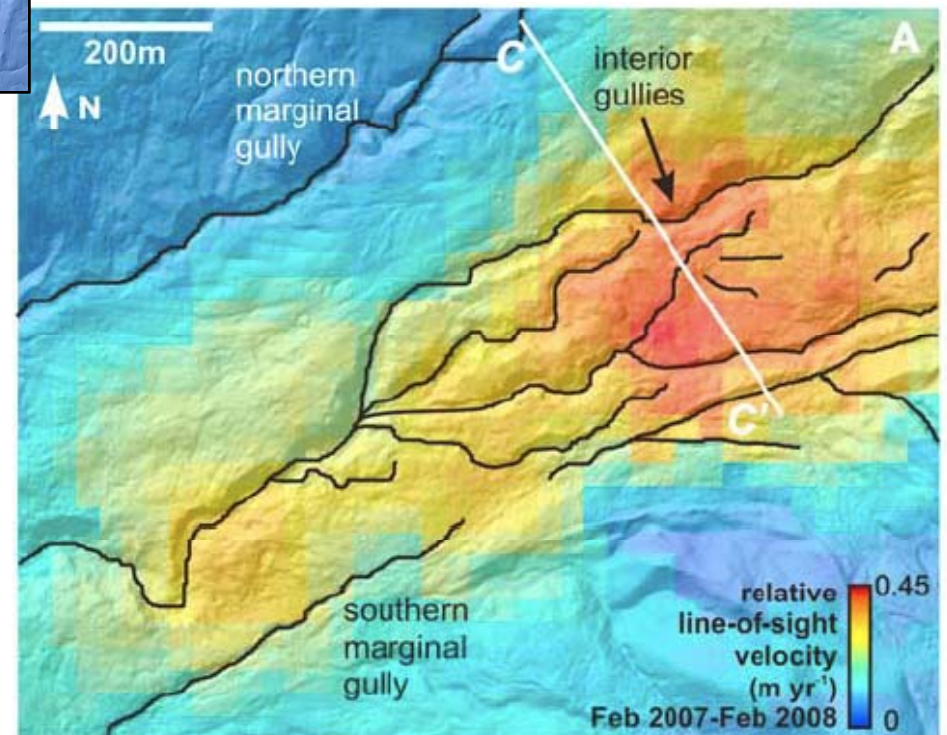


## Geomorphic implication #1: Sediment production at Boulder Crk

- Sediment flux through transport zone exceeds  $4100 \text{ m}^3 \text{ yr}^{-1}$
- Basin lowering rate  $\geq 1.6 \text{ mm yr}^{-1}$
- Gullies appear to facilitate delivery of earthflow-mobilized sediment to the channel network



*Roering et al. (GRL, 2009)*



# Hillslope evolution and nonlinear slope-dependent transport

$$q_s = \frac{-K(h)S}{1 - (S/S_c)^2}$$

$S$  = hillslope gradient,  $\nabla z$

$h$  = soil depth (m)

$K$  = transport coefficient ( $L^2$ )

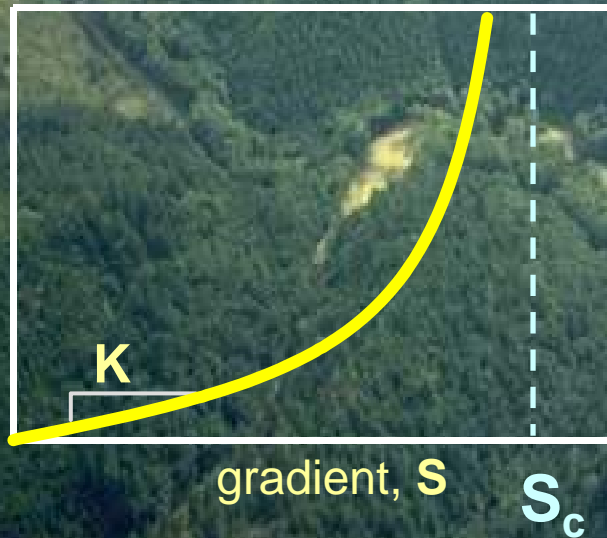
$S_c$  = critical gradient

*Physically-based formulation:*

*Roering et al., 1999*



sediment flux,  $q_s$



- $K$  varies with energy expended by disturbances in the soil mantle...biological connection?

