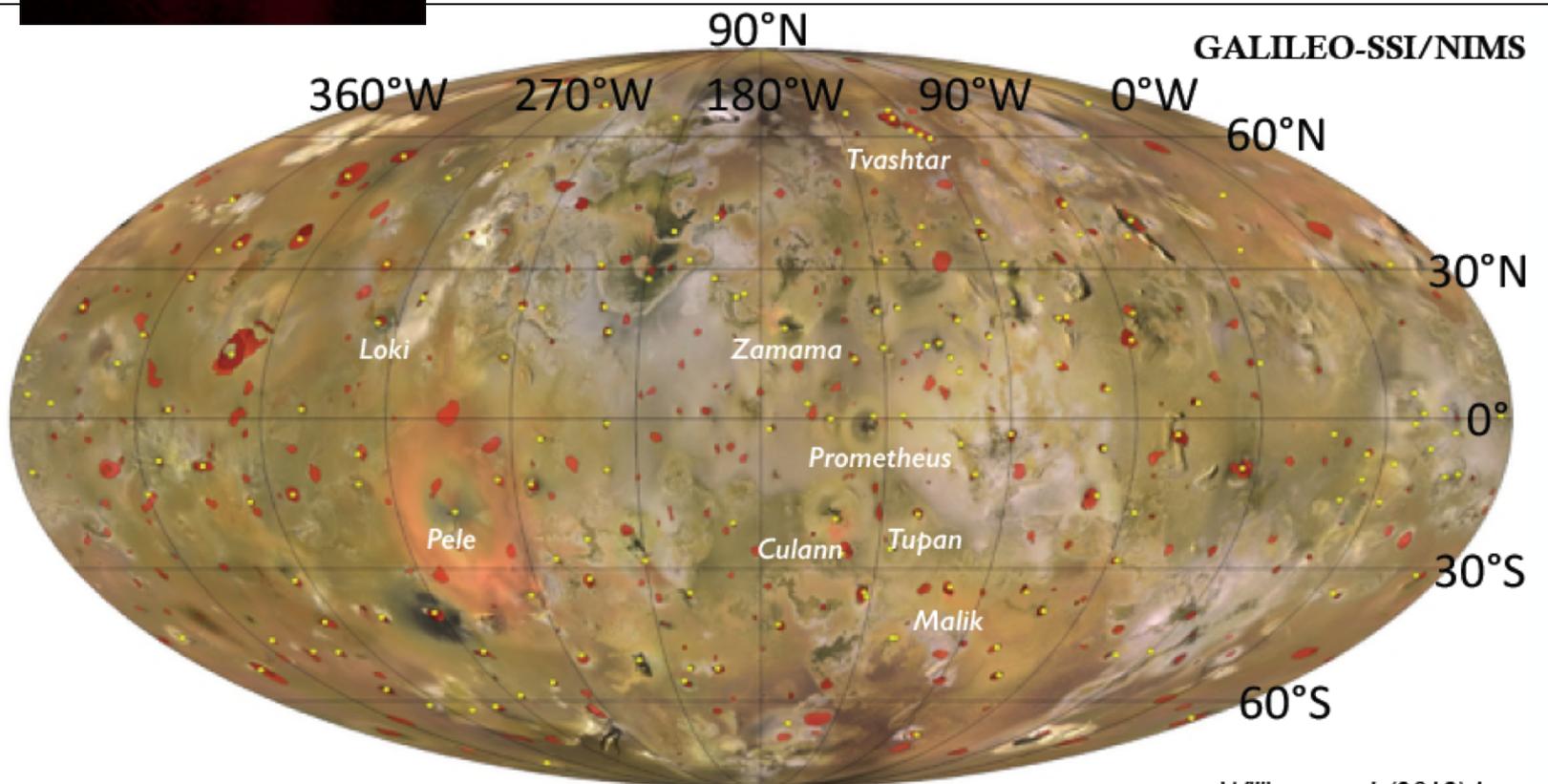
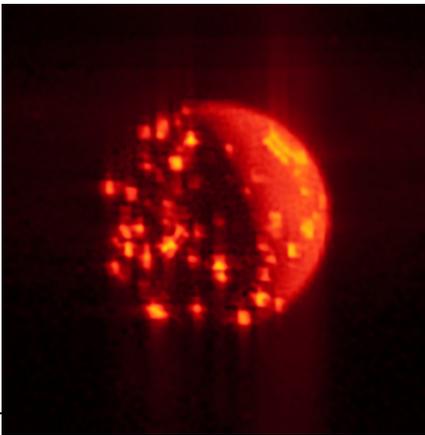


IO ESCAPE PROCESSES

Melissa A. McGrath
SETI Institute

Io has extensive volcanic activity

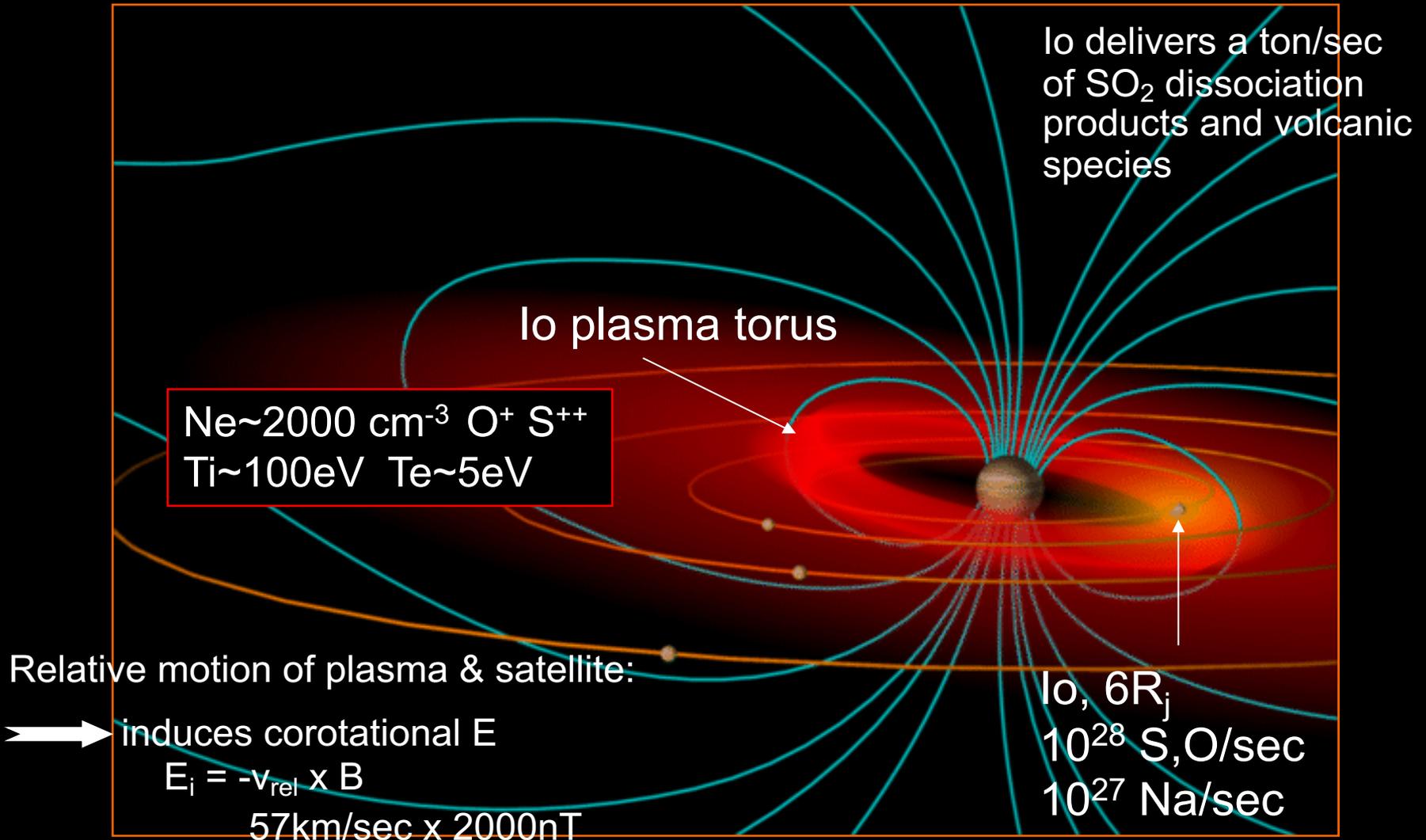


- Paterae floor units ($N = 529$)
- Hotspots ($N = 173$)

Williams et al. (2012) *Icarus*

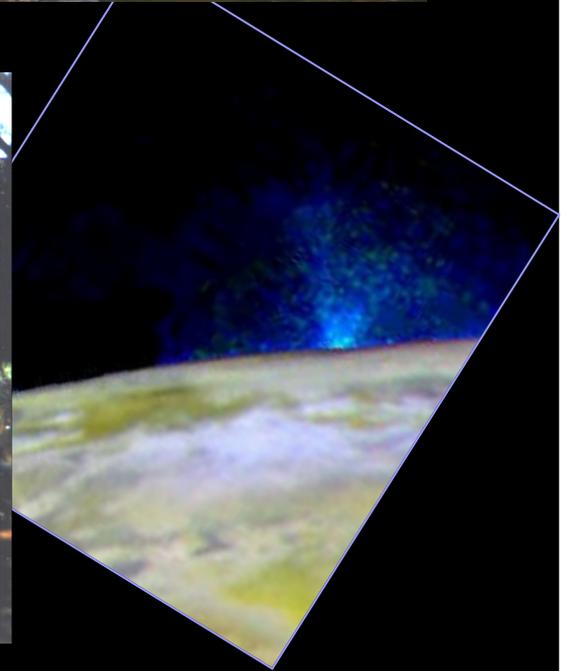


Io & the Jovian magnetosphere



- ~500kV potential across Io
- drives currents of few x 10⁶ Amps

“Small” volcanoes (the majority, plume ht <10s of km) – nothing escapes

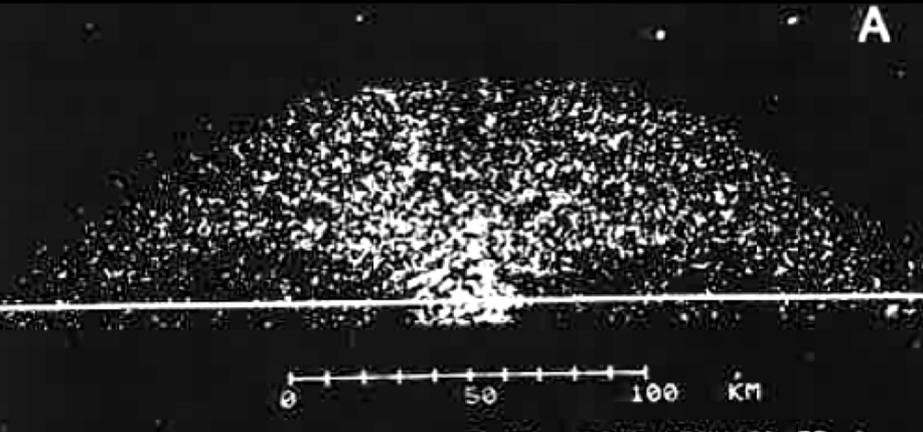


“Large” volcanoes – 100s of km in ht, only a few of these

A



A



Detailed modeling of plumes,
ballistic & shock models

(Strom et al. 1981)

Limiting envelope for a set of
ballistic trajectories:

$$z_M = v_o^2/2g - gr_M^2/2v_o^2$$

z = height

r = radial distance from vent

g = surface gravity

v_o = ejection velocity

at vent, $r_M=0$, $\rightarrow v_o = \sqrt{2gz_M}$

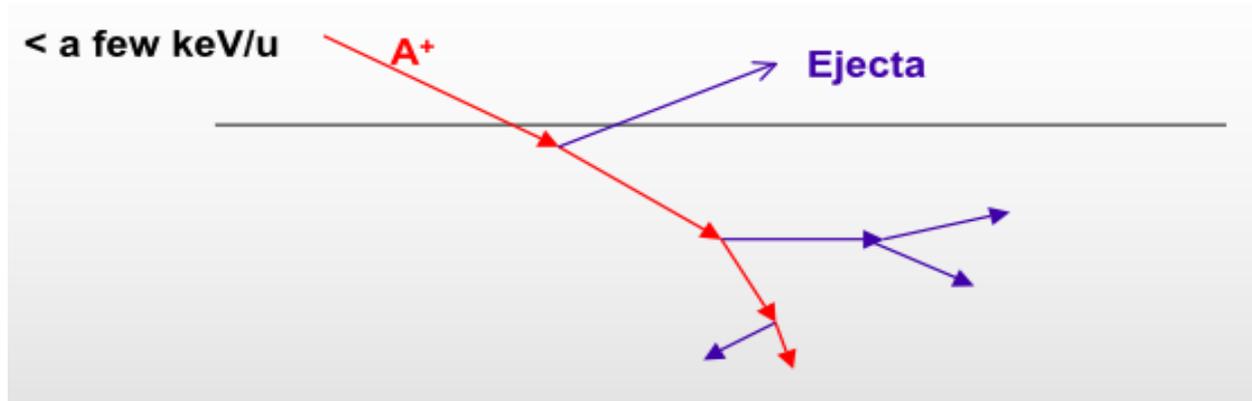
$$v_o \leq 1 \text{ km/sec}$$

$$v_{\text{esc}} = 2.56 \text{ km/sec}$$

➔ nothing escapes

Surface sputtering

An elastic, multiple-collision-cascade process triggered by an ion impacting a solid surface that ejects multiple particles from relatively close to the surface.



Process characterized by:

- (1) Yield = # particles removed per incident ion
 - (2) E spectrum of sputtered particles
- Both can be measured experimentally.

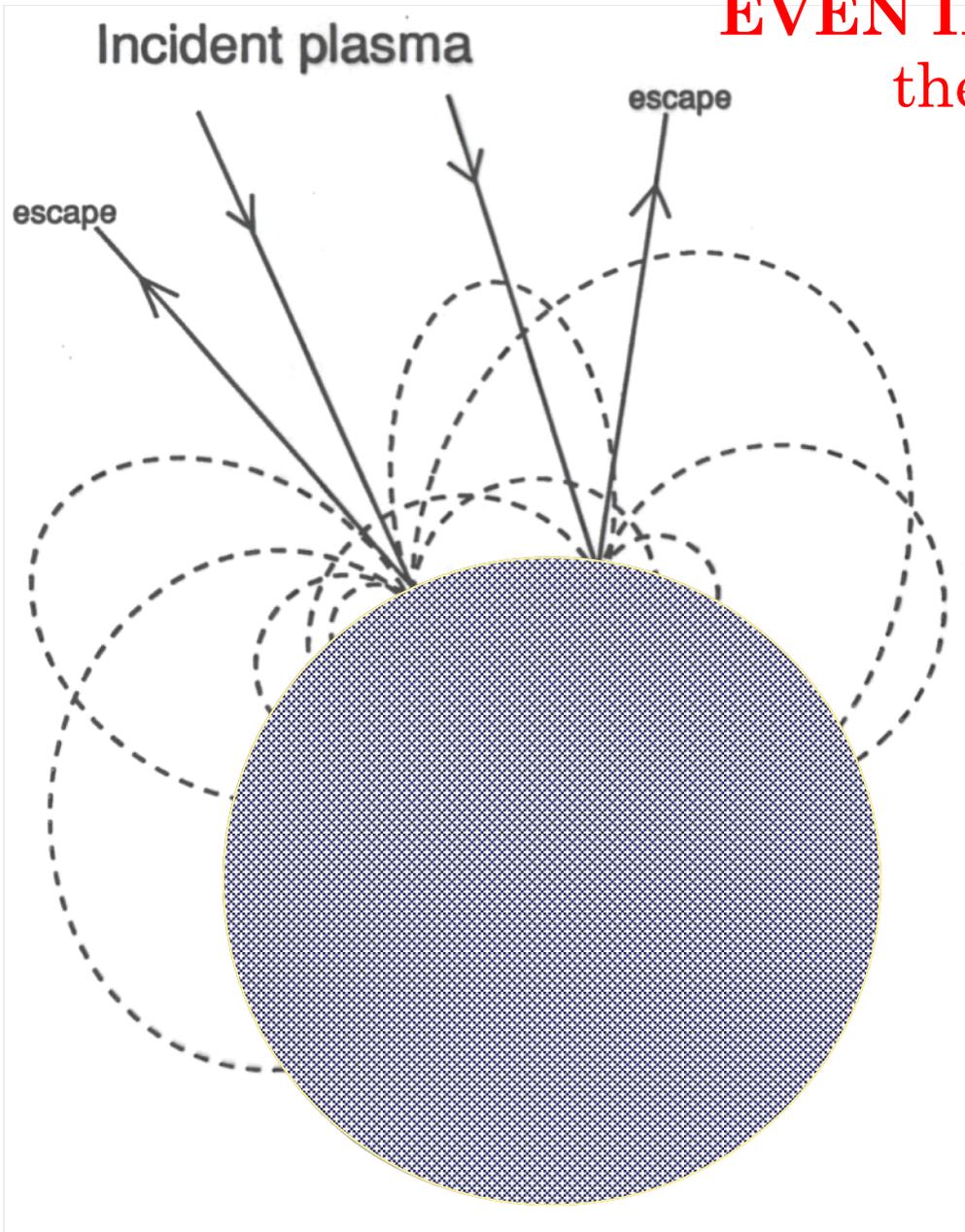
At lo:

$$Y \sim 10-100$$

$$Y(E) \sim E/(E+E_b)^3, E_b \sim 0.5-4$$



EVEN IF all the plasma reached the surface, only a few % of sputtered particles would escape



Io's Local Interaction

- Electrodynamics:
Currents from ion/neutral collisions and pickup
deflect flow

Saur et al. 2002

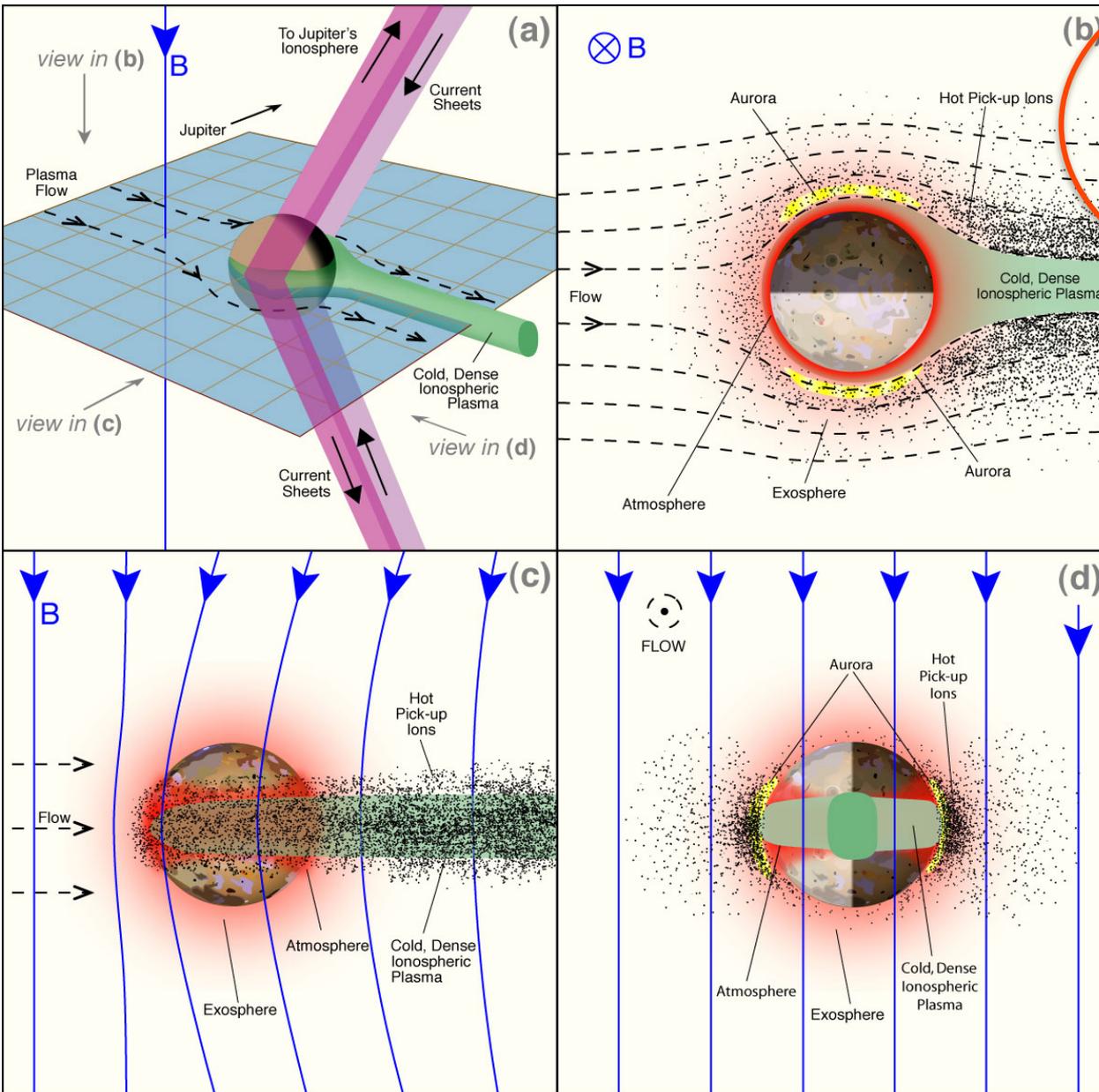
- Heating, ionization and charge-exchange in atmosphere

Smyth 1998

- Cooling, deceleration of upstream plasma
- Acceleration of downstream plasma

Delamere et al. 2003

Slide courtesy of Fran Bagenal



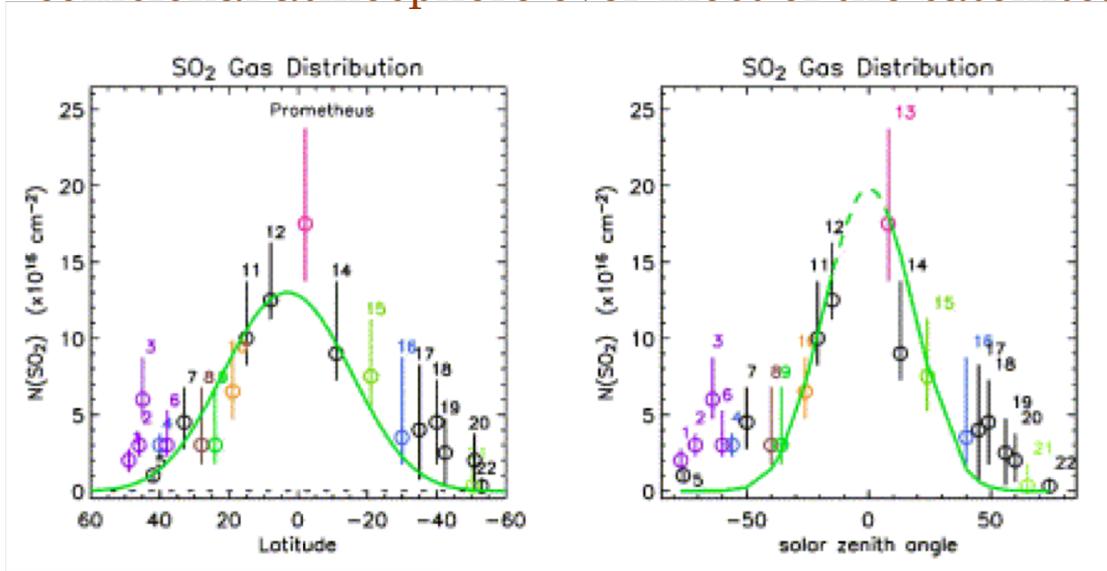
Potential escape processes at Io

From surface:
(1) Volcanoes
(2) Sputtering

Volcanoes and sputtering primarily feed an atmosphere, but do not contribute to escape from Io. Escape occurs almost exclusively from the atmosphere.

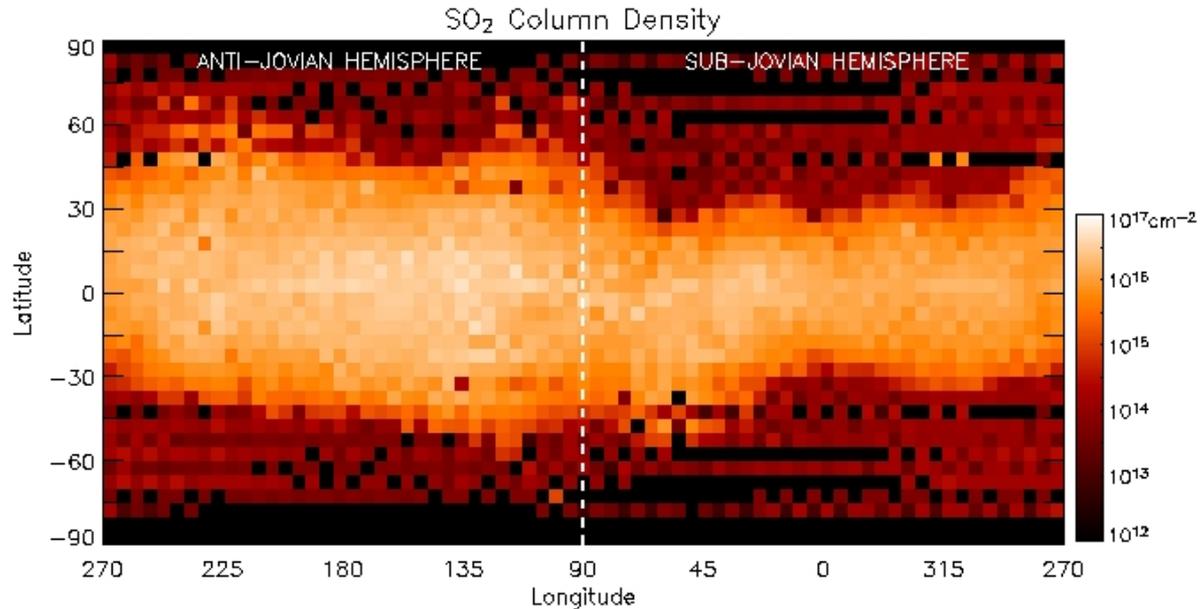


Volcanoes predominant output is SO₂ gas, which at temperatures relevant for Io condenses to cover the surface with SO₂ frost, and also produces a collisional atmosphere over most of the satellite.



Anywhere the temperature is low enough (e.g., poles, terminator, far from hot spots) SO₂ condenses; everywhere else it's part of the collisional atmosphere. Most of the dissociation products of SO₂ do not condense.

Jessup et al. 2004



Feaga et al. 2009

Other Species

Table 19.2. Summary of Io Atmospheric Species

McGrath et al. 2004

Species	Io Abundance*	Reference
SO ₂	~ (1-10) × 10 ¹⁶ in ~ ± (30-45) ^o latitude band ~ (2-10?) × higher in active volcanos	Synthesis of all observations; see §19.2.2, 19.2.3, and 19.2.4 McGrath <i>et al.</i> 2000; Spencer <i>et al.</i> 2000; Spencer <i>et al.</i> 2002, Jessup <i>et al.</i> 2003
S ₂	1 × 10 ¹⁶ , Pele plume (t), SO ₂ /S ₂ ~ (3-12)	Spencer <i>et al.</i> 2000
SO	~ (0.03-0.1) × SO ₂	Lellouch 1996
NaCl	(0.003-0.013) × SO ₂ , active volcanos	Lellouch <i>et al.</i> 2003
O ₂	Inferred (modeling)	Kumar (1982,1985); Summers 1985; Summers & Strobel 1996; Wong & Johnson 1995,1996; Wong & Smyth 2000; Moses <i>et al.</i> 2002a, 2002b
S	3.6 × 10 ¹² < N _S < 1.3 × 10 ¹⁴ (t) ~ 9 × 10 ¹² ; at 2R _{Io} (t) = 0.1 × O	Feaga <i>et al.</i> 2002 (upper limit revised up; see text) Wolven <i>et al.</i> 2001
O	> (4-7) × 10 ¹³ , disk average ~ 1 × 10 ¹⁴ ; at 2R _{Io} (t) = 11 × S	Ballester 1989 Wolven <i>et al.</i> 2001
Na	4 × 10 ¹² , disk average	Bouchez <i>et al.</i> 2000 [see also Burger <i>et al.</i> 2001, Retherford 2002]
K	(1-10) × 10 ⁸ ; Na/K = 10 ± 5 at (10-20) R _{Io}	Brown 2001
Cl	~ 1 × 10 ¹³ , disk average	Feaga and McGrath (2002)
H	~ 2 × 10 ¹²	Strobel & Wolven 2001
CS ₂	< 2 × 10 ¹⁴	McGrath <i>et al.</i> 2000; Spencer <i>et al.</i> 2000; Spencer <i>et al.</i> 2002
CO	< (3.6-6) × 10 ¹⁷	Lellouch <i>et al.</i> 1992
H ₂ S	< (0.7-1.2) × 10 ¹⁶	Lellouch <i>et al.</i> 1992
OCS, S ₂ O, ClO, CS, NaOH	Not detected (mm)	Lellouch <i>et al.</i> 1992
KCl	< 1 × NaCl	Lellouch <i>et al.</i> 2003

* Numbers in vertical column density, cm⁻², unless otherwise noted; (t) = tangential



Potential escape processes at Io - 2

From atmosphere:

(3) **Thermal:** Jeans escape

(4) A host of **non-thermal processes** including:

- Ionization of atmosphere by photons and electrons; ions are then swept up the by Jovian magnetic field.
- Elastic collisions between plasma ions and atmosphere neutrals
- Charge exchange between fast plasma ions & slow atmosphere neutrals

These processes have been considered extensively by numerous authors for many years, and the relative importance of each has been calculated and modeled in some detail.

[A great reference for more detail is Sieveka & Johnson 1984.]



Thermal (“Jeans”) escape = thermal evaporation

$$\Phi_{\text{escape}} = \frac{n(z)v_0}{2\sqrt{\pi}} \left(\frac{v_{\text{esc}}^2}{v_0^2} + 1 \right) \exp\left(-\frac{v_{\text{esc}}^2}{v_0^2} \right) \quad \text{particles m}^{-2} \text{ s}^{-1}$$

where

$n(z)$ is the number density

$$v_0 = \sqrt{\frac{2kT}{m}} \quad \text{is the most probable velocity, as above}$$

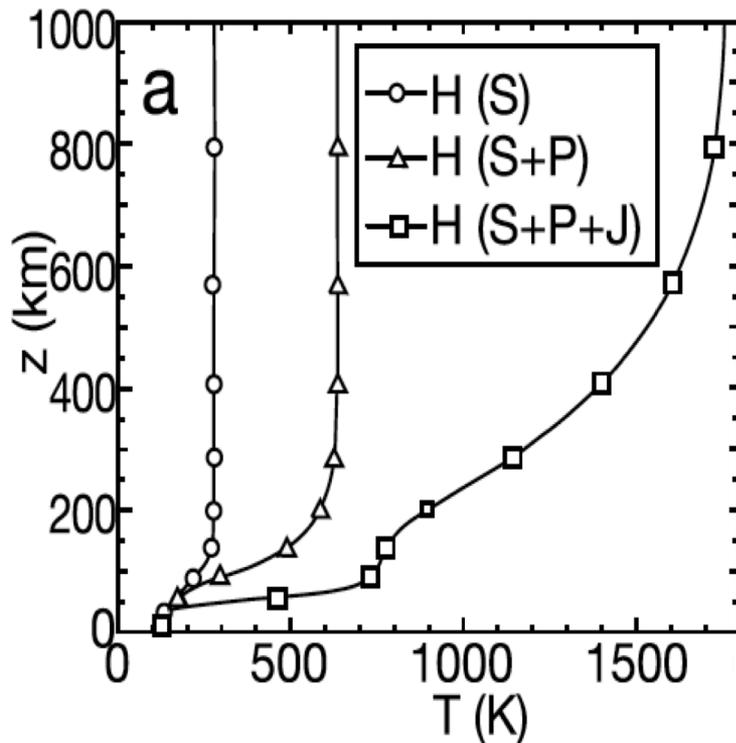
$$v_{\text{esc}} = \sqrt{\frac{2GM_{\text{planet}}}{(R_{\text{planet}} + z)}} \quad \text{is the escape velocity}$$

–Important factor is the final exponential term.

To get a handle on Jeans escape we need to understand the structure of the atmosphere, viz. temperature and composition vs altitude.



The thermal structure is dominated by plasma energy deposition and the current system (which produces so-called Joule heating just another form of ion-neutral collisions which is parameterized by the electron current).



S = solar heating (photons)
P = plasma heating
J = Joule heating

Exobase, boundary between collisional and non-collisional regimes of the atmosphere, is ~ 500 km

Strobel et al. 1994

Jeans escape at Io is most often calculated using a temperature of $T=1000$ K.



Numerous models for the composition vs altitude exist:

- Kumar 1982 – sublimation atmosphere
- Summers & Strobel 1996 – sublimation atmosphere
- Wong & Johnson 1996
- Moses et al. 2002 – volcano atmosphere

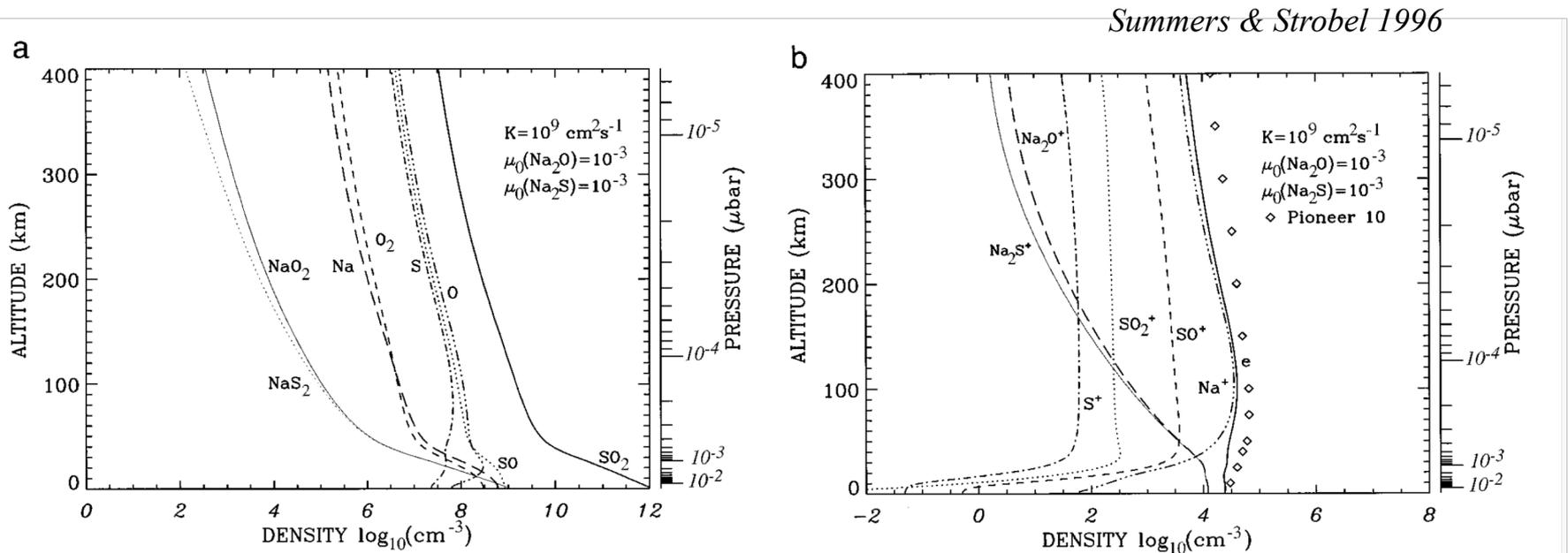


FIG. 7. (a) Major constituents of the neutral atmosphere from the surface to 400-km altitude for the high-density SO_2 case B2 with high values of the eddy mixing coefficient $K_{zz} = 10^9$, and (b) major ions and n_e for B2.

Ionization

This is simple because you mostly just look up or calculate the cross sections/rates using the known plasma characteristics. Once something is ionized it's fate is pretty simple – it is captured and swept away by the magnetic field into the torus. One caveat is the ionosphere, which is part of the “steady state” atmosphere (I'll mention some caveats later).



Elastic collisions

Often approximated using the hard sphere formulation where the cross section is determined by the physical size of the colliding objects (familiar example is billiard balls).

Using a Thomas-Fermi Coulomb potential is much more accurate for ion-neutral collisions, and provides simple formulas for energy distribution and direction of impacting and target species after collision:

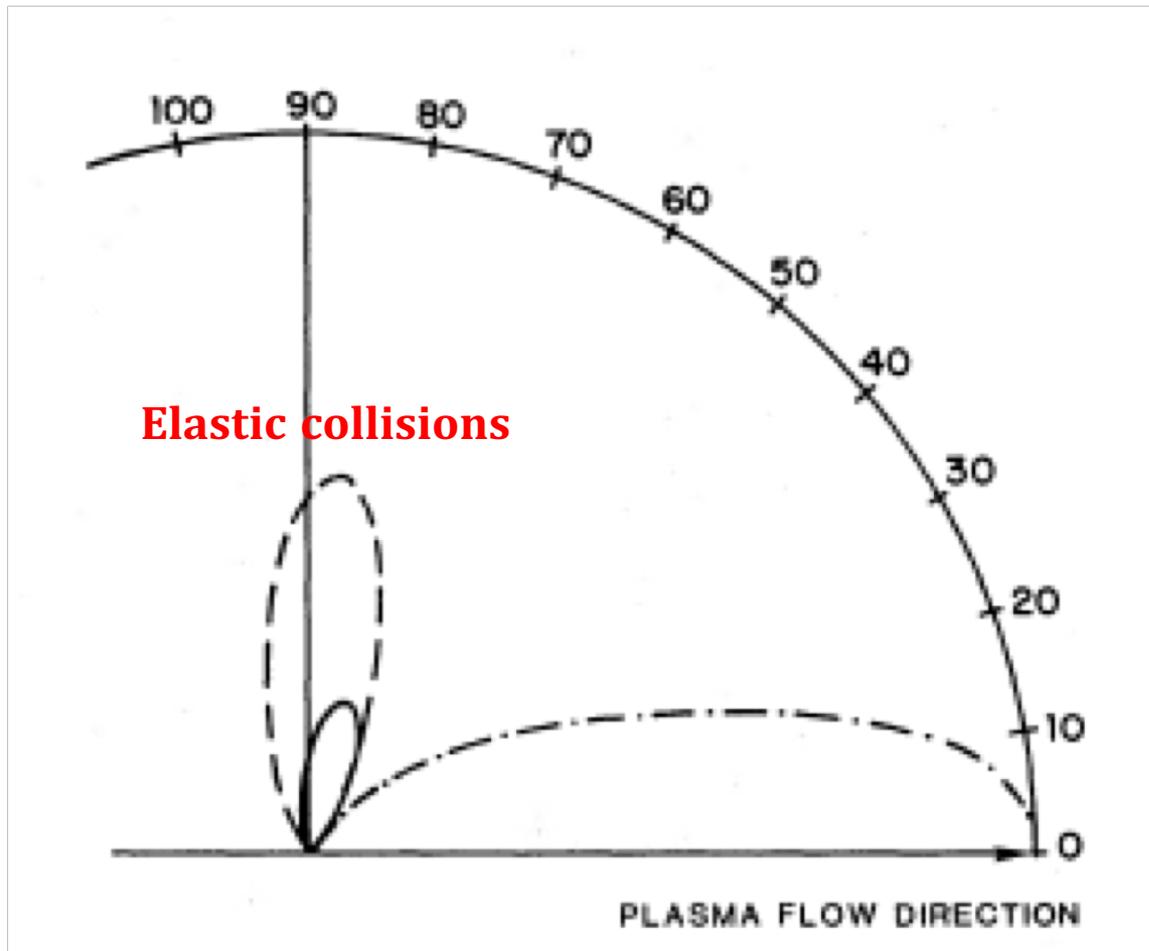
$$\frac{d\sigma}{dt} \approx \frac{\pi a_{12}^2}{2t^{3/2}} \left\{ \frac{\lambda t^{1/6}}{[1 + (2\lambda t^{2/3})^{2/3}]^{3/2}} \right\}$$

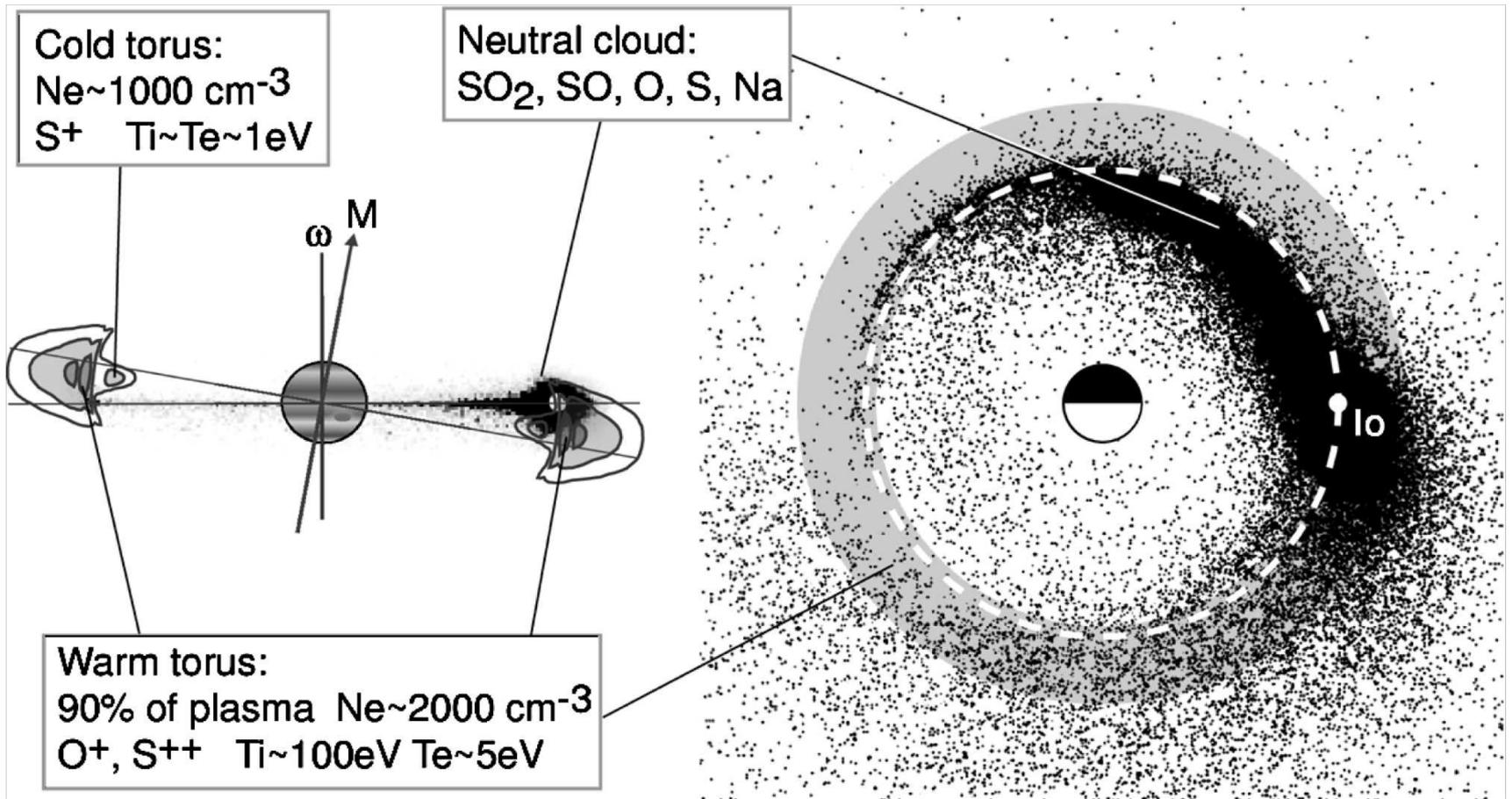
$$\cos \theta_1 = [1 - (E/E_1)(M_1 + M_2)/2M_1]/(1 - E/E_1)^{1/2}$$

$$\cos \theta_2 = (E/E_{\max})^{1/2}$$



Elastic collisions result primarily in low energy target particles directed $\sim 90^\circ$ away from the impactors. Most of these escape Io but not Jupiter so they go into Jupiter orbit near Io and form neutral clouds.





Thomas et al. 2004



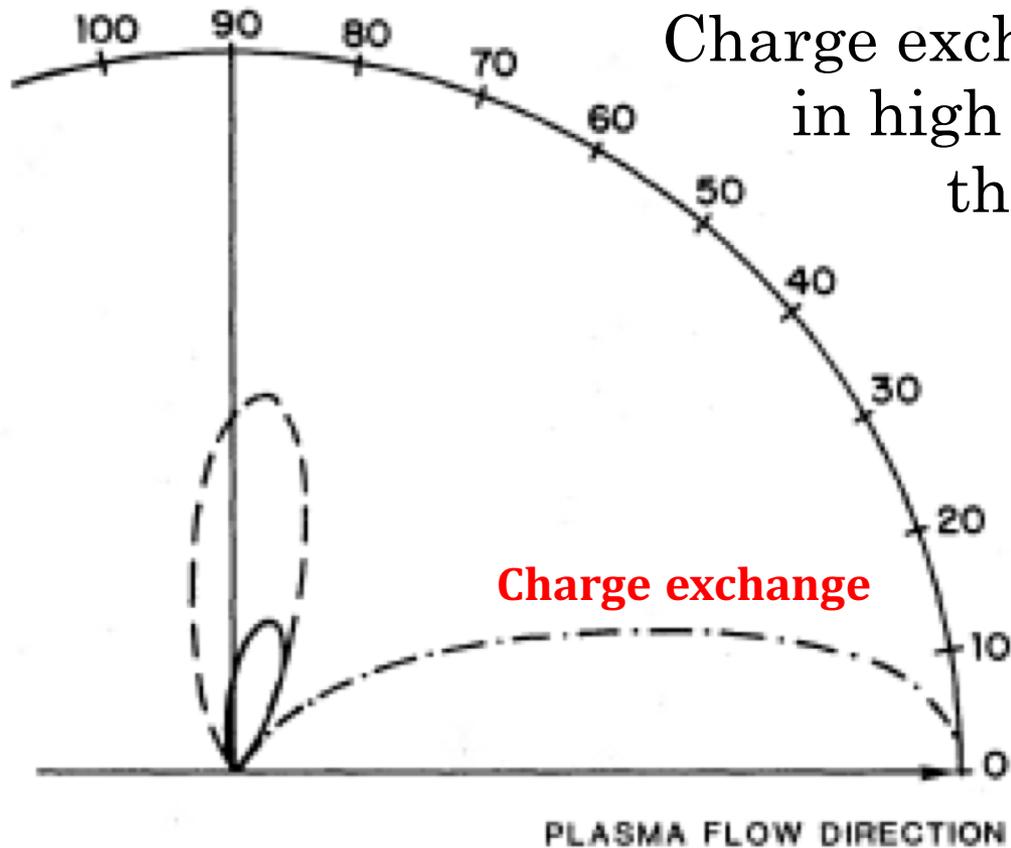
Charge exchange

An ion encounters a neutral and they exchange charge:



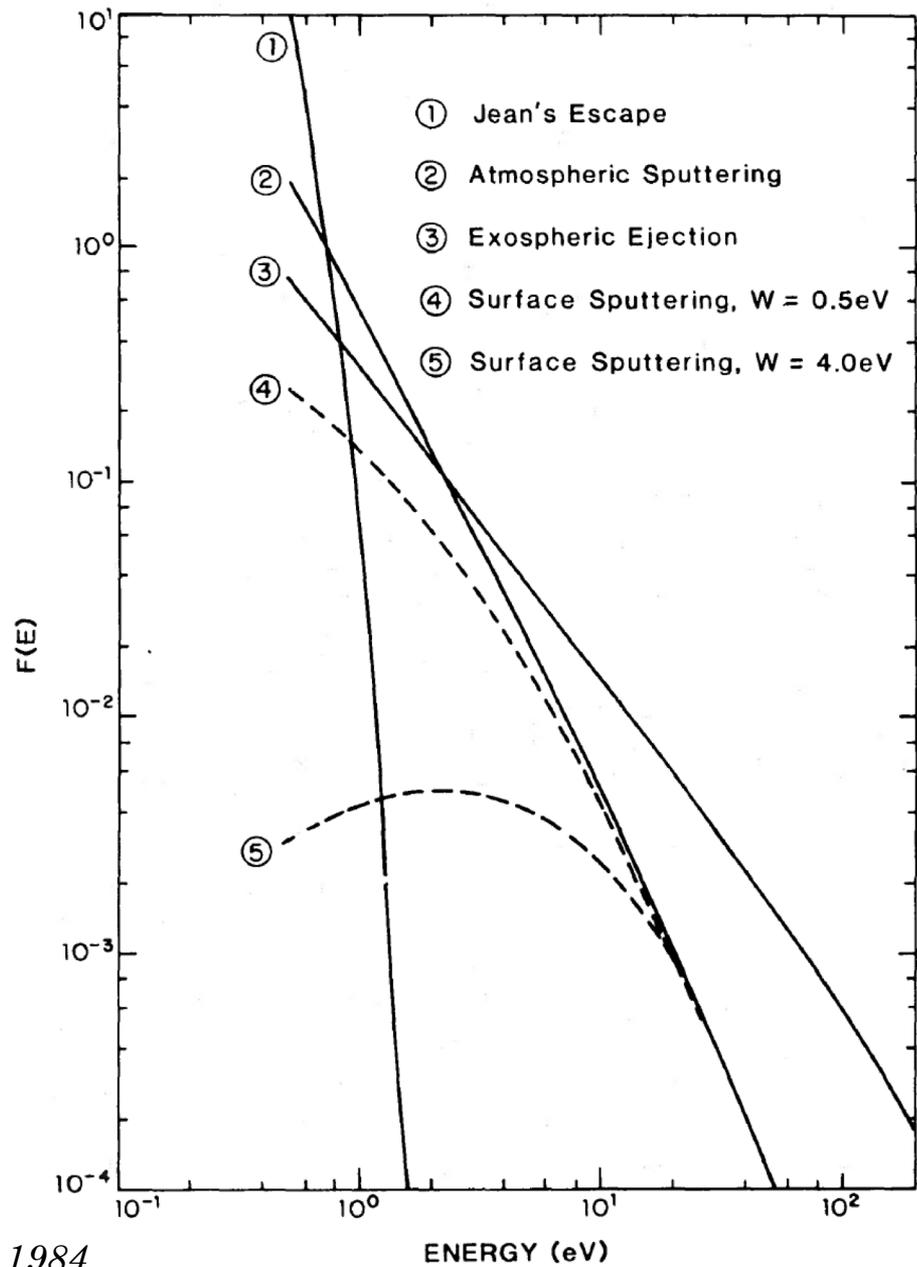
A ends up with the same velocity as A^+ . This is important at I_0 because the ions have much higher energies than the neutrals, so CHEX provides a viable process for energizing neutrals to escape speed.





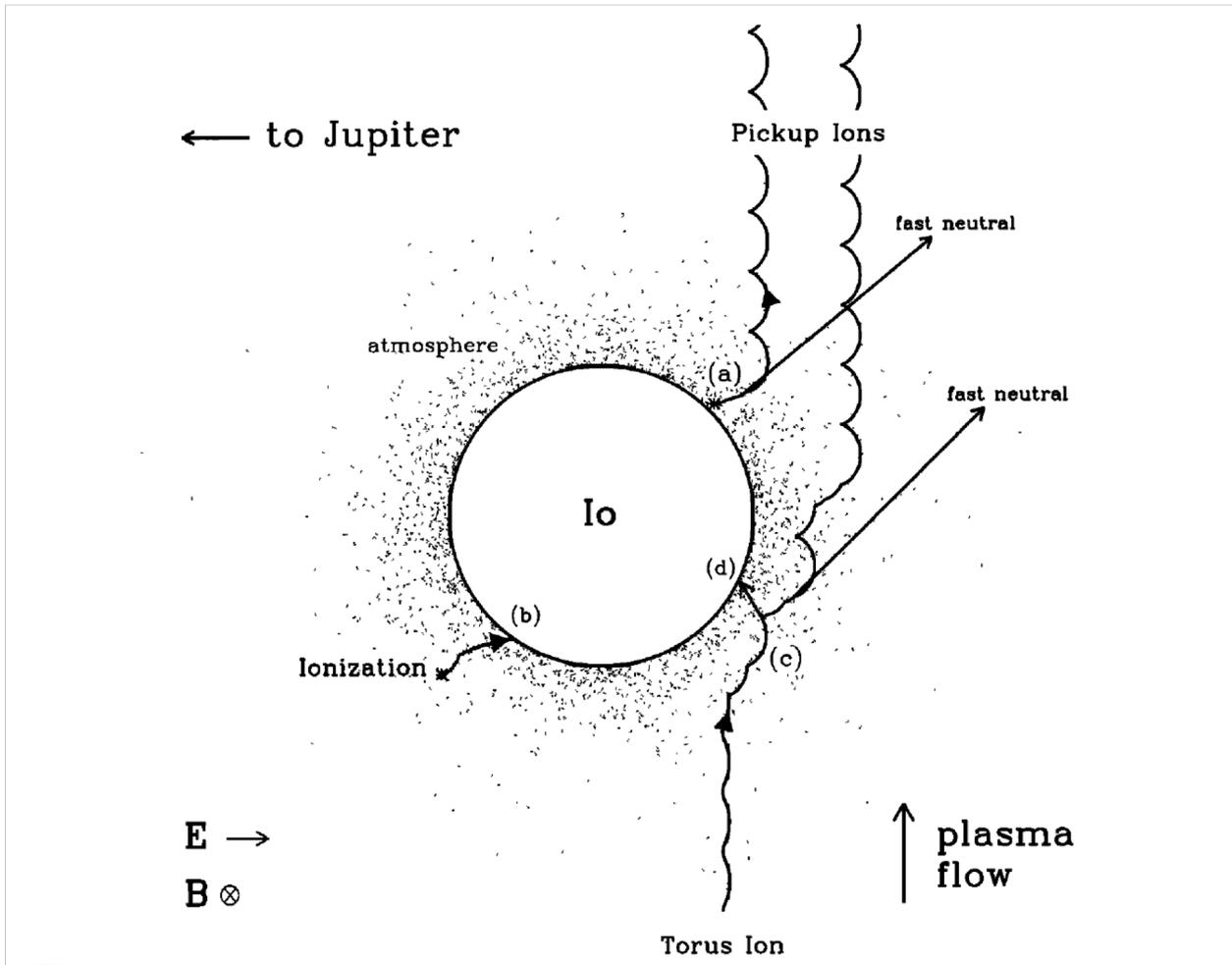
Charge exchange results primarily in high energy target particles that are forward directed and escape Jupiter as well as Io.





Sieveka & Johnson 1984





Wilson & Schneider 1999



Relative importance of these processes

PROCESS	Na IN ATMOSPHERE AS Na		Na IN ATMOSPHERE AS Na ₂ S ^d	
Exobase	$1R_{Io}$	$2R_{Io}$	$1R_{Io}$	$2R_{Io}$
Jeans escape 1000 K	10^{24}	10^{27}	10^{15}	10^{22}
Atmospheric sputtering	1×10^{26}	8×10^{26}	6×10^{25}	5×10^{26}
“Exospheric” ejection ^e	3×10^{25}	5×10^{26}	2×10^{25}	3×10^{26}

Sieveka & Johnson 1984

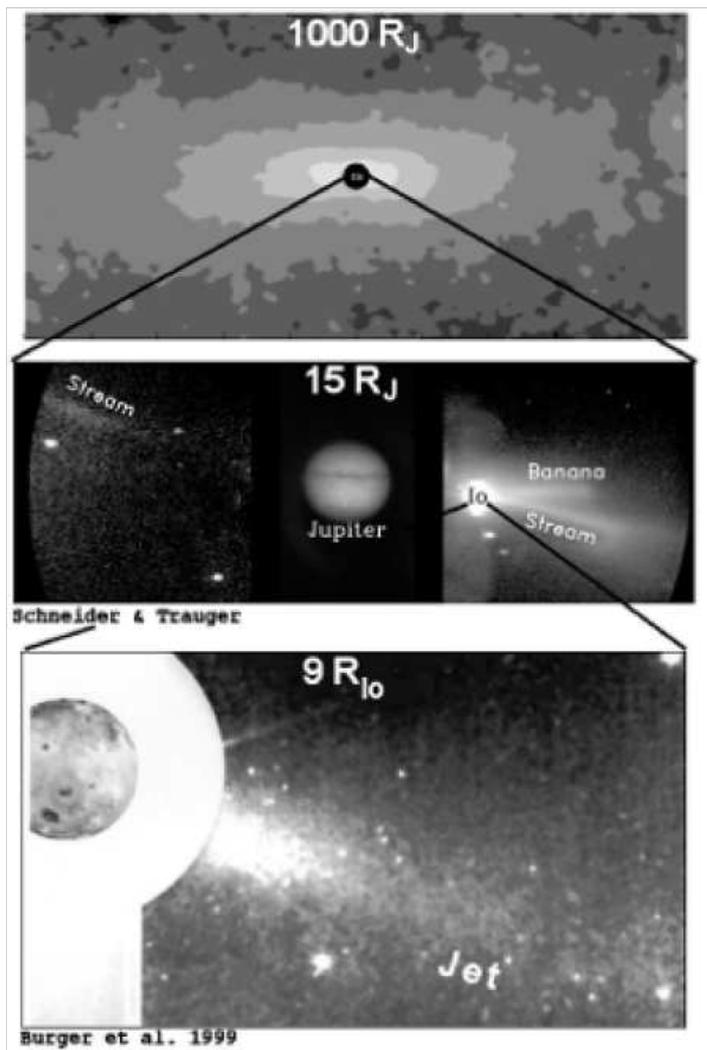
CORONAL SOURCES ^a (#SEC ⁻¹)								
Target	Y_{es}	Sputter from exobase	Charge transfer		Single collision		Electron ionization	
			In	Out ^b	O ⁺	S ⁺	In	Out ^b
SO ₂	4.1	1.3×10^{28}	3.4×10^{27}	4.8×10^{27}	6.4×10^{27}	7.0×10^{27}	8.8×10^{27}	1.9×10^{28}
SO	6.2	3.1×10^{28}	8.6×10^{27}	1.5×10^{28}	1.4×10^{28}	2.4×10^{28}	3.2×10^{27}	7.8×10^{27}
O ₂	11.3	3.5×10^{28}	6.3×10^{27}	1.4×10^{28}	1.6×10^{28}	2.6×10^{28}	2.5×10^{27}	7.7×10^{27}
S	13.5	4.2×10^{28}	1.7×10^{28}	2.8×10^{28}	2.3×10^{28}	3.6×10^{28}	3.1×10^{28}	6.4×10^{28}
O	17.5	5.4×10^{28}	9.1×10^{27}	2.1×10^{28}	2.2×10^{28}	3.5×10^{28}	3.4×10^{27}	1.0×10^{28}

McGrath & Johnson 1987

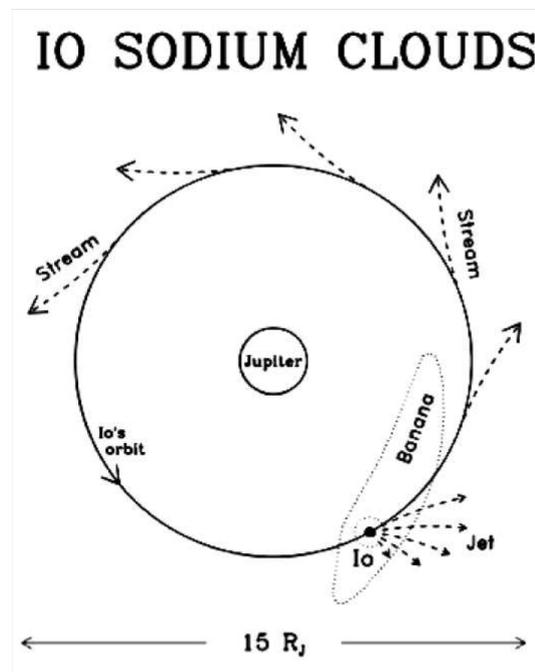


Na is highly visible at Io via resonant scattering of sunlight so provides an excellent tracer of some of these processes.

View of Na emission from Earth



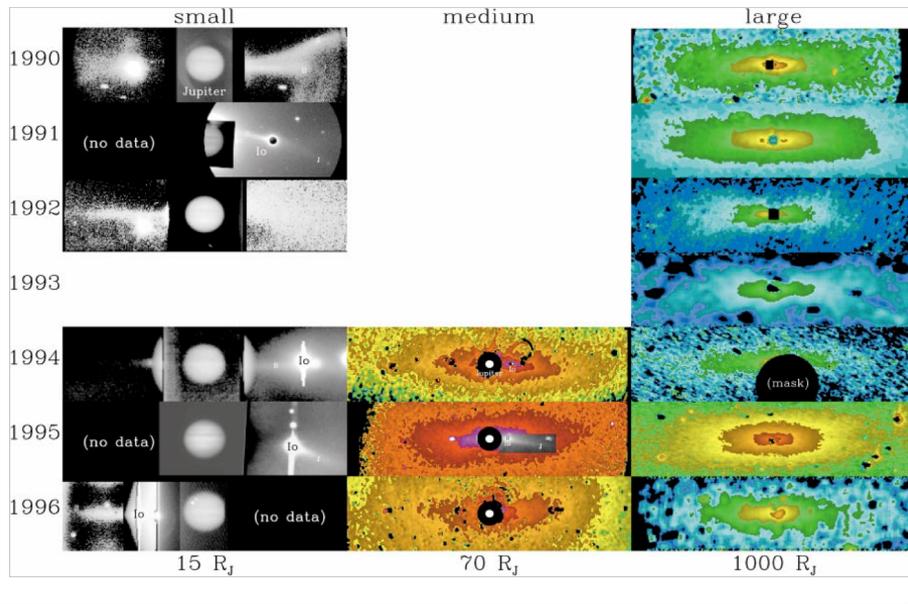
View from above Jupiter N pole



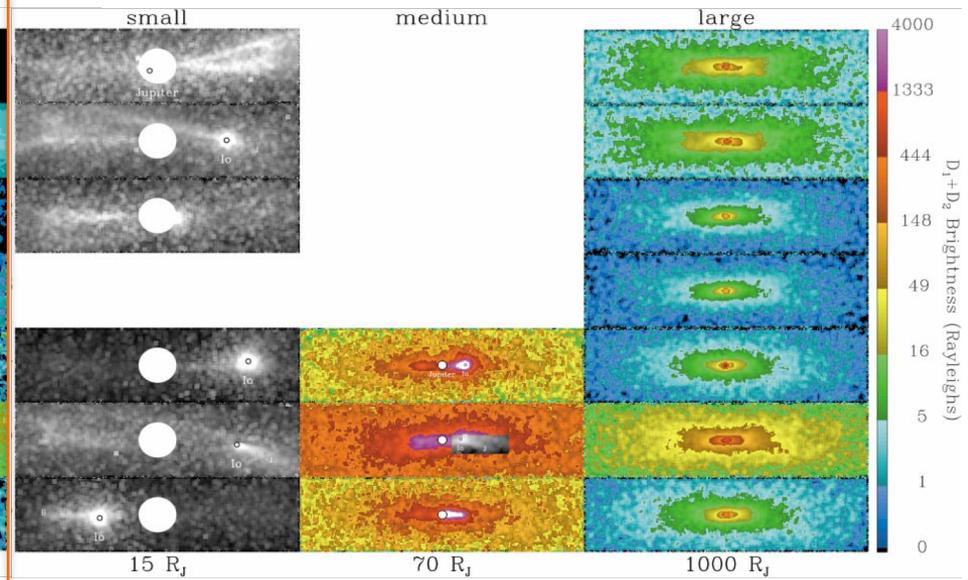
Thomas et al. 2004



Data



Best model



The most important things to keep in mind for escape processes at Io:

- Io is immersed in a dense, heavy-ion plasma
- Io is subjected to a strong electric field that drives substantial currents at the satellite
- Io has a stable, collisional atmosphere



Summary

- Escape at Io is dominated by non-thermal processes involving Jovian magnetospheric plasma and neutrals in Io's atmosphere and neutral clouds.
- Independent of the indisputable observational evidence for a collisional atmosphere at Io, the supply rates for S,O,Na require such an atmosphere.
- Observed Na clouds and jets confirm that non-thermal processes dominate the escape at Io.
- To what extent do volcanoes affect the atmosphere at the exobase? Is the atmosphere well mixed due to active plumes?



The End

