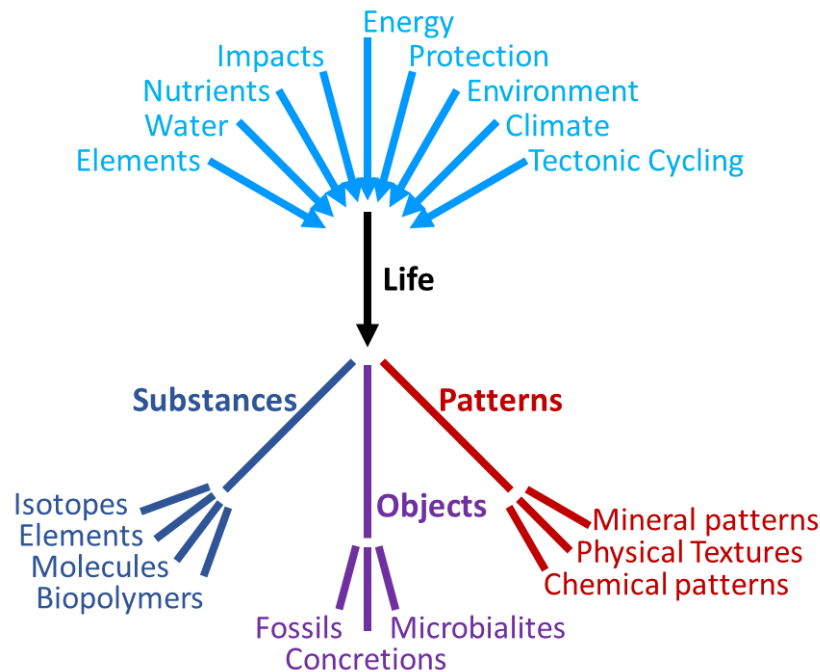


Understanding Astrobiological False Positives from Terrestrial Microbiology

Martin Homann

Biosignatures in terrestrial and planetary contexts



- **Objects, substances, and/or patterns** whose origin specifically require a biological agent (Des Marais et al. 2003).
- Candidate biosignatures are ranked by 3 criteria: **reliability, survivability, and detectability.**
- The usefulness of a biosignature comes not only from the probability of life having produced it, but also from the **improbability of non-biological processes** producing it (Mustard et al, 2013).

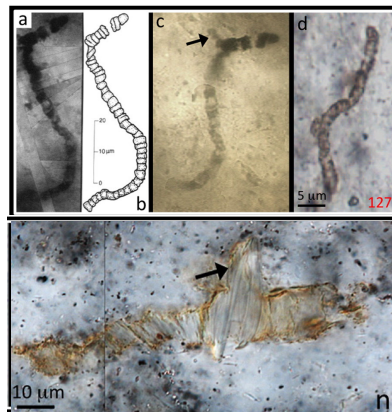
False positives and controversial biosignatures

Allan Hills 8441 (4.09 Ga)



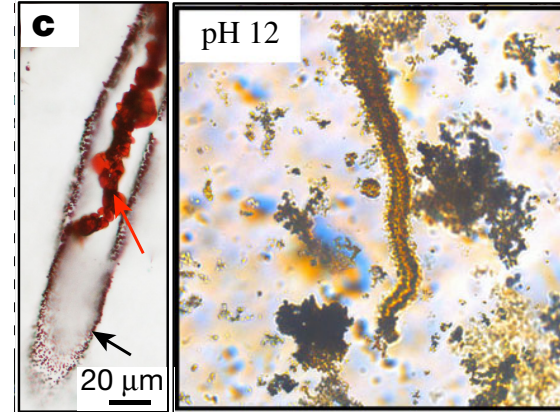
McKay, 1996: NASA

Apex filaments (3.5 Ga)



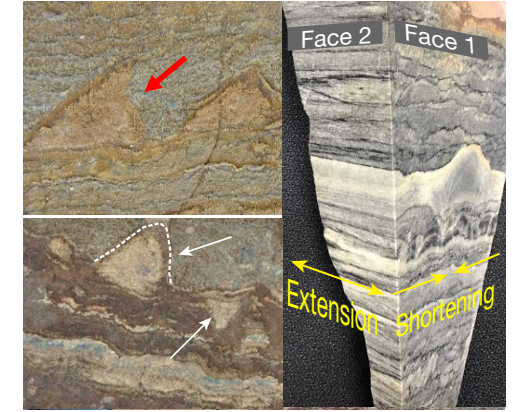
Schopf, 1993
Brasier et al., 2002
Wacey et al., 2015

Hematite filaments (3.77 Ga)



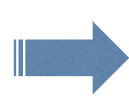
Dodd et al., 2017
McMahon, 2019

Stromatolites? (3.7 Ga)



Nutman et al., 2016
Allwood et al., 2018

- Abiotic nanobacteria-like structures
- Mineral artefacts; sheets of phyllosilicates with carbonaceous coatings
- Possibly abiotic Fe-mineralized chemical gardens
- Probably products of structural deformation of layered rocks



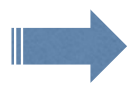
Morphological evidence alone is insufficient to determine the biogenicity of a lifelike structure.

What lines of evidence provide the clearest validation of ancient biology in the rock record?

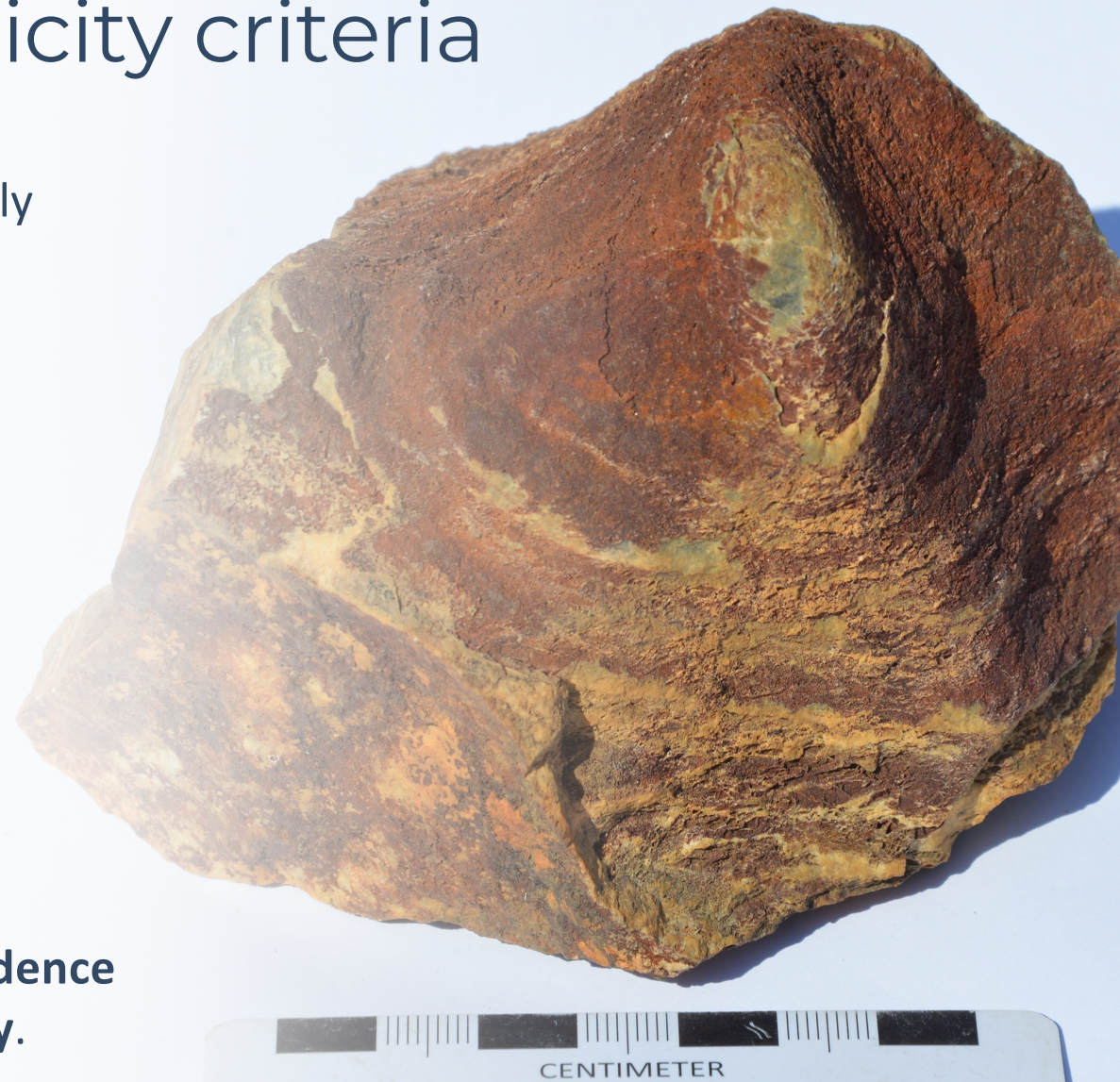


Biogenicity criteria

- Did the object form in a demonstrably habitable paleoenvironment?
- Is it assuredly endogenic and syngenetic to the host rock?
- Are morphology and biofabric consistent with a biotic origin?
- Is the chemical and isotopic composition distinctively life-like?



Multiple, independent lines of evidence are needed to establish biogenicity.



Fossil Biosignatures in Deep Time and Space

Life as we know it

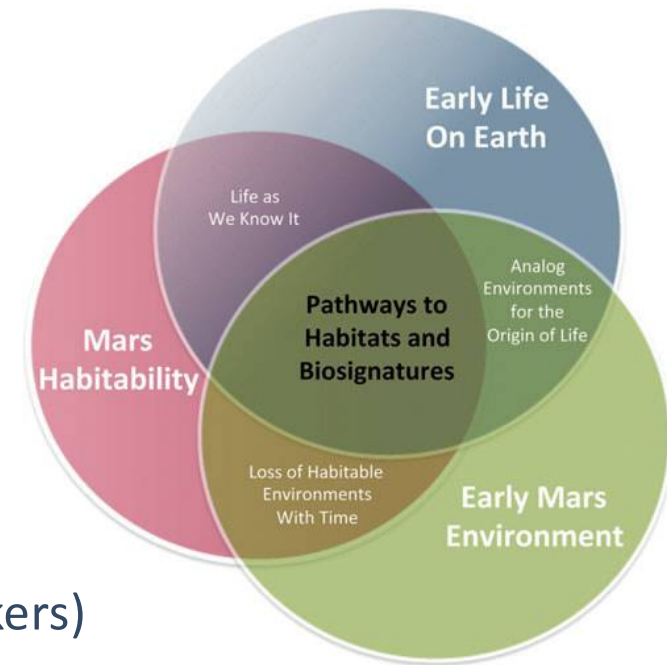
- Carbon-based and requiring a water solvent
- CHNOPS – key biogenic elements of life on Earth

Where to look for?

- Archean sedimentary rock record: 4 – 2.5 Ga
- Lithified aquatic sediments: carbonates, cherts, ... sandstones

What to look for?

- Morphological biosignatures (biofabrics, microfossils)
- Chemical biosignatures (stable isotopes, biominerals, biomarkers)
- Traces of microbial life: microbial mats, biofilms, organic matter



➡ Early diagenetic mineralization is key for preservation and survivability of biosignatures.

Microbial mats

Laminated microbial communities; mainly bacteria and archaea



Microbial mats

Laminated microbial communities; mainly bacteria and archaea

Biochemical
sediments



Microbialites



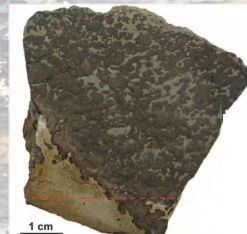
modern



fossil



Stromatolites



Thrombolites

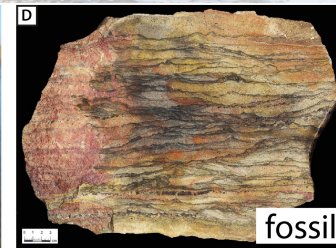
Siliciclastic
sediments



Microbial mats



modern

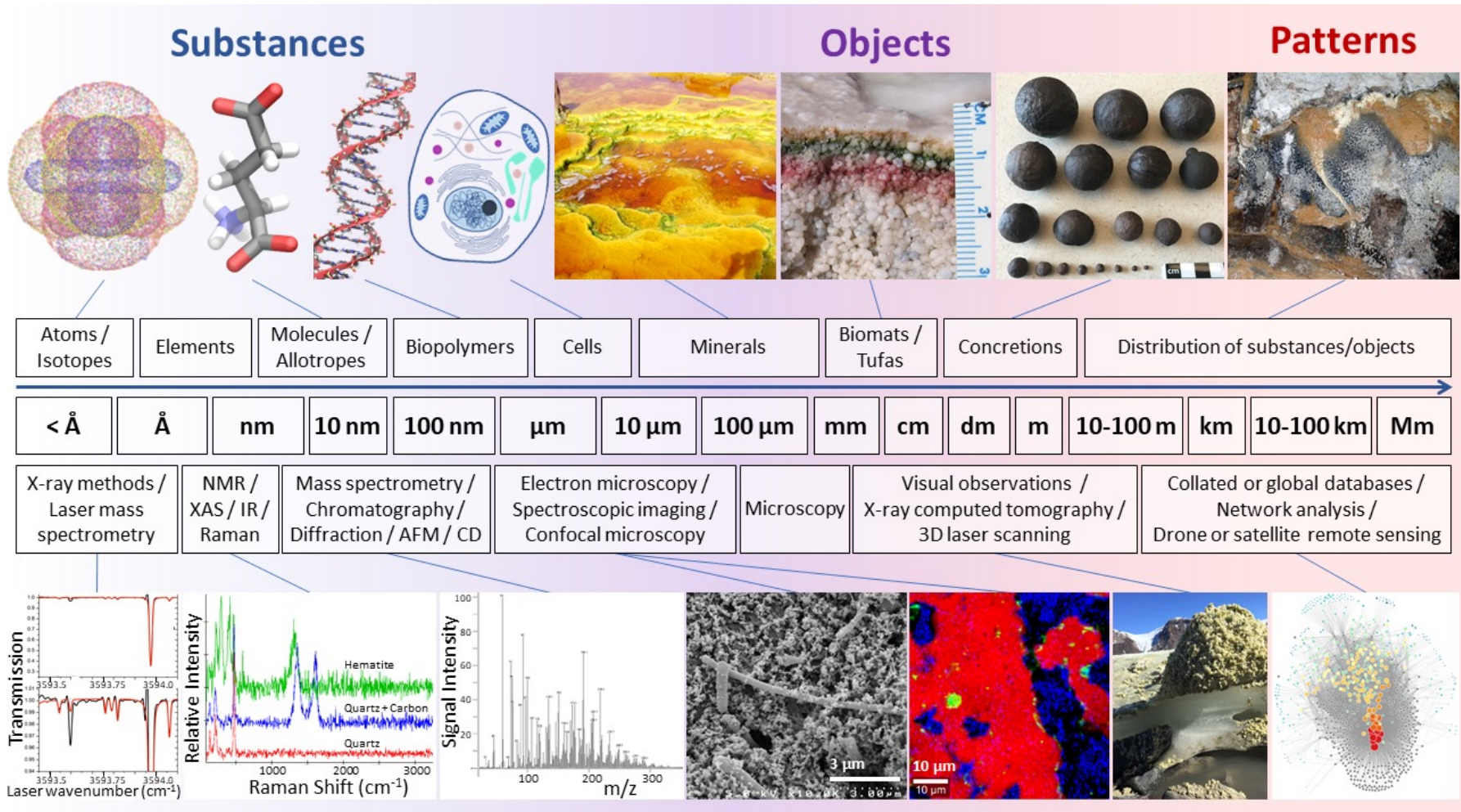


fossil



Bahar Alouane,
Tunisia

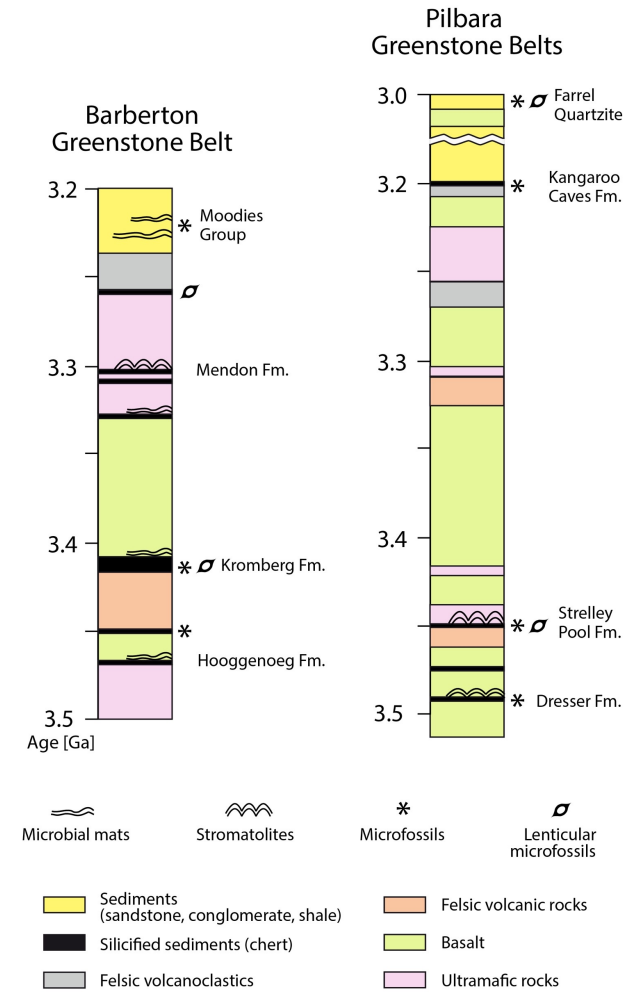
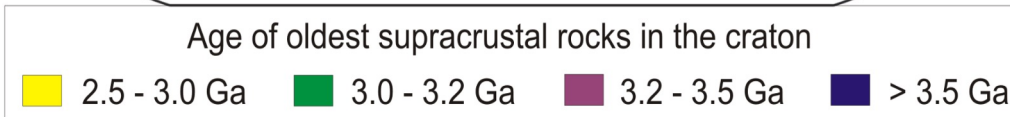
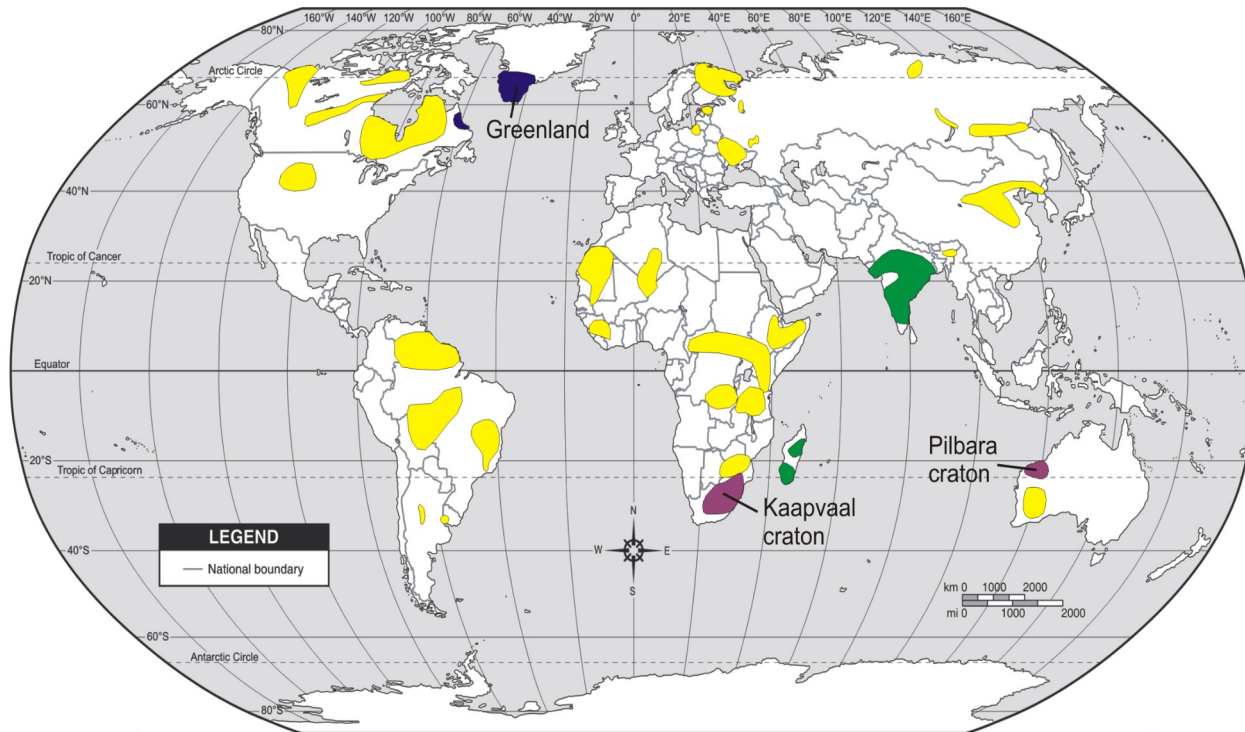
Biosignature and life detection methods



Examples of terrestrial biosignatures in deep time



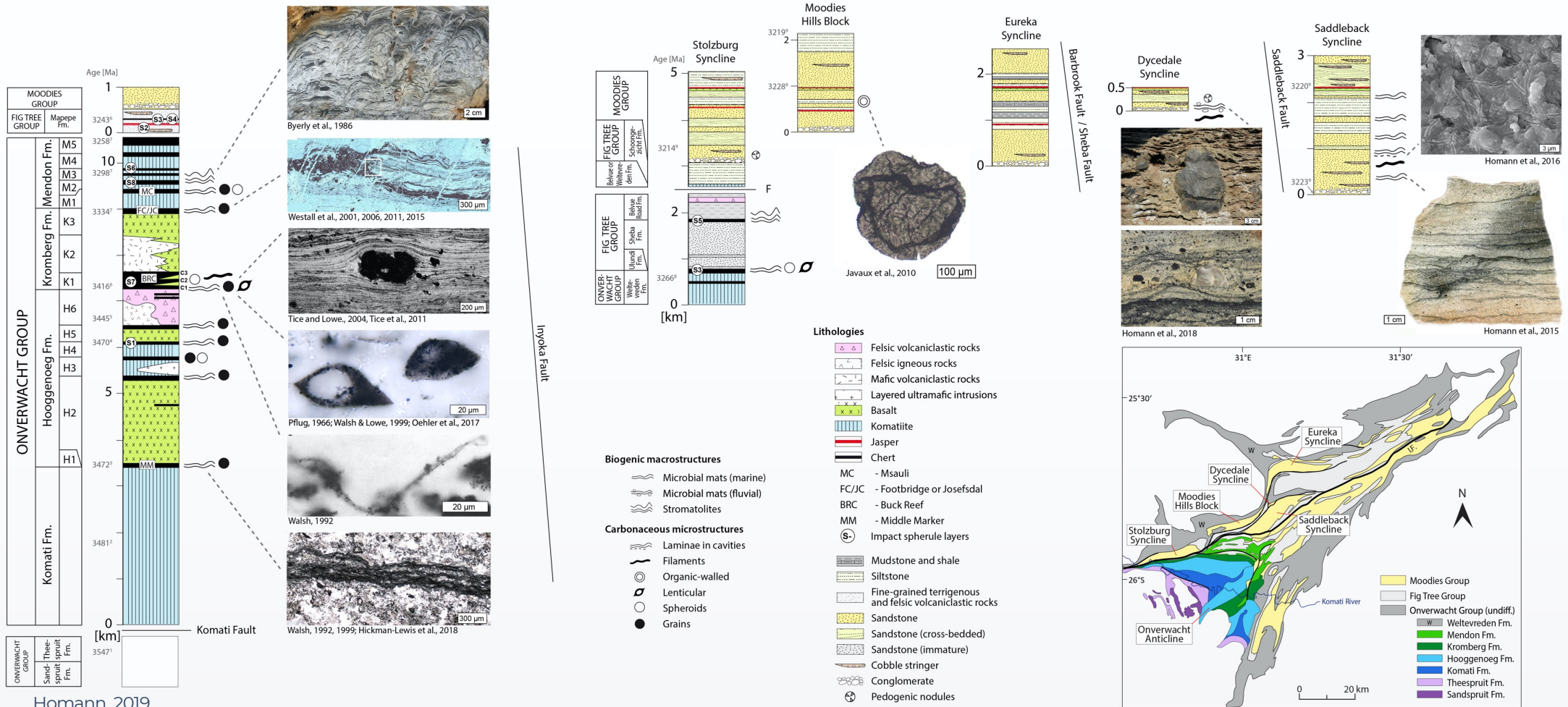
Traces of early life in the Archean rock record



Barberton Greenstone Belt (BGB), South Africa



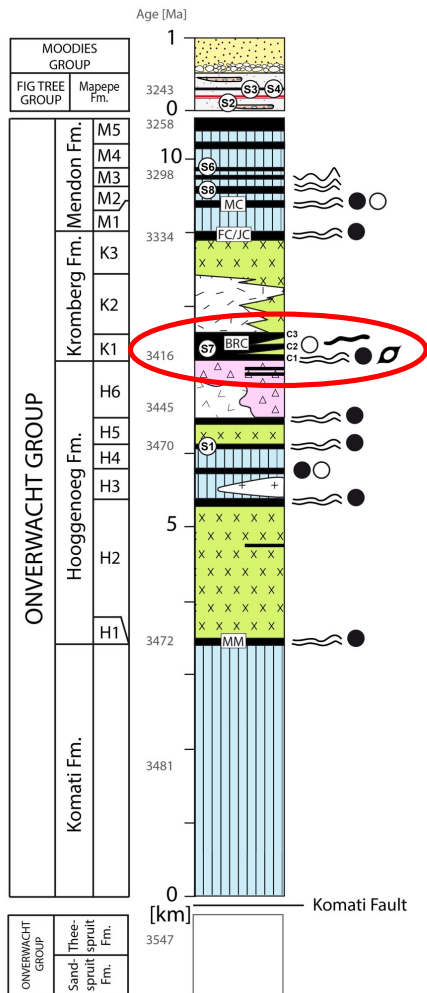
Evidence of early life in the BGB (3.5 – 3.2 Ga)



Homann, 2019

Traces of early life in the Buck Reef Chert

Stratigraphic thickness: 250 – 400m; exposed for nearly 50 km



Black and white banded chert

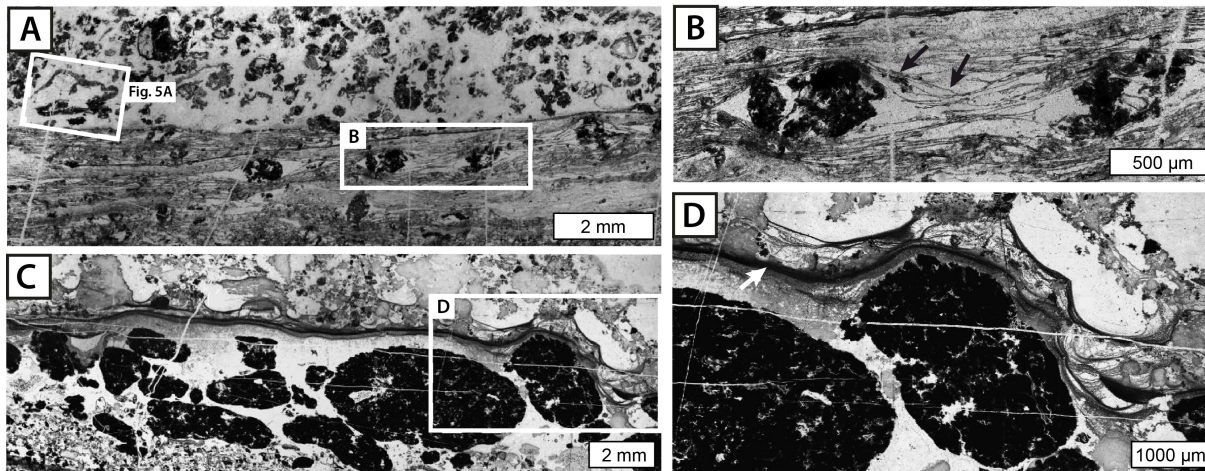


Silicified evaporites

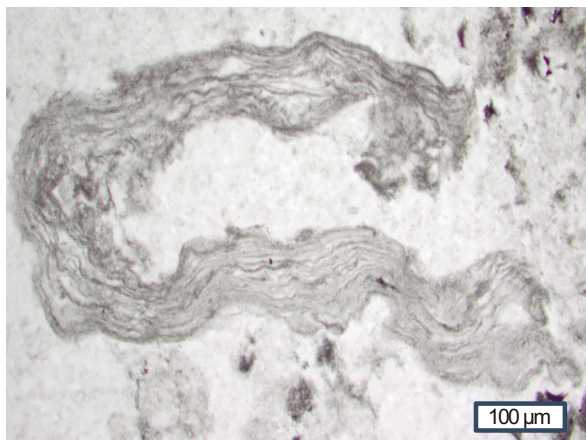


Microbial mats in the 3.4 Ga Buck Reef Chert

Silicified mats draping carbonaceous grains

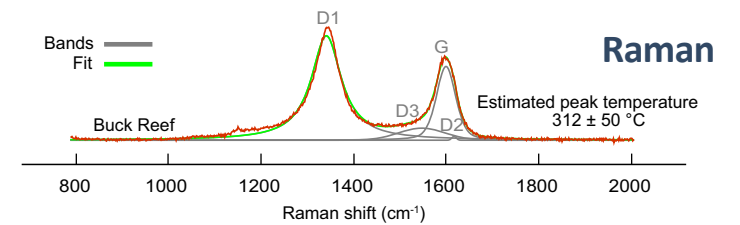
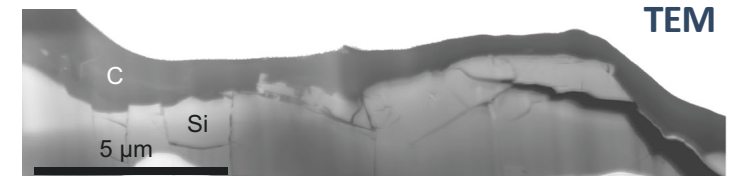
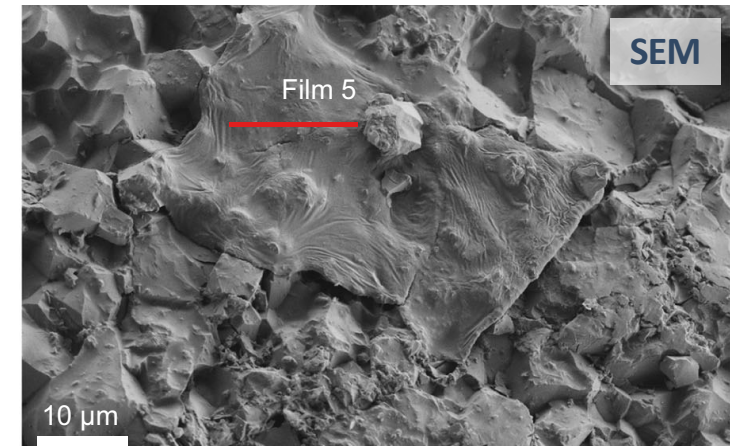


Tice and Lowe et al., 2004; Tice, 2009



Eroded, rolled-up mat fragments indicate former cohesive consistency.

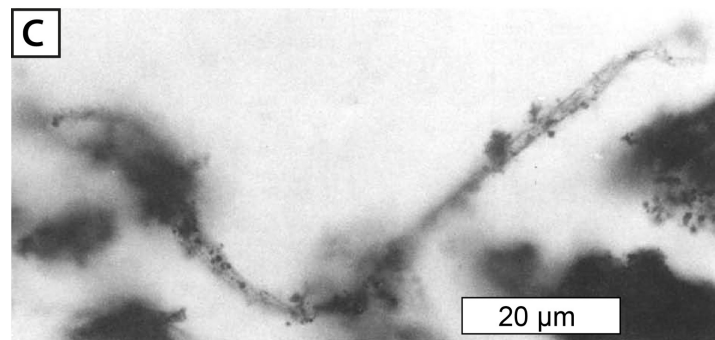
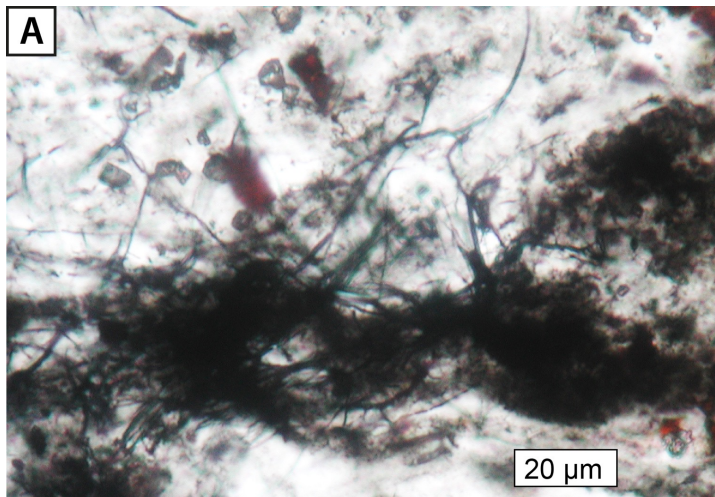
Carbonaceous biofilms



Alleon et al., 2021

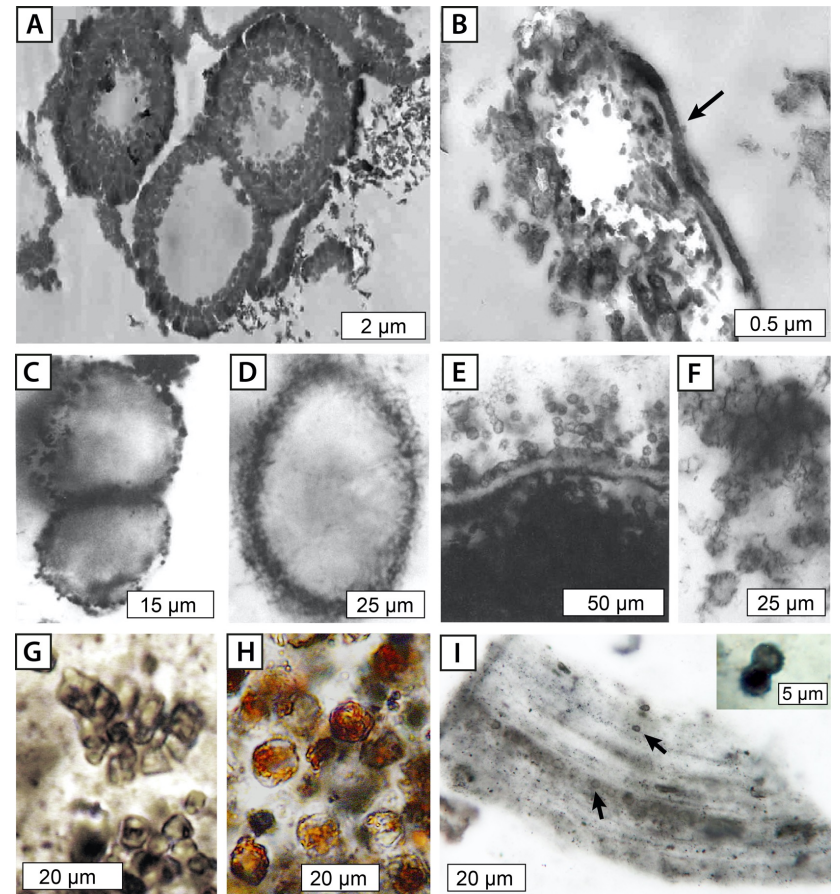
Microstructures of (non)biogenic origin

Filamentous structures



Walsh 1992; 2000

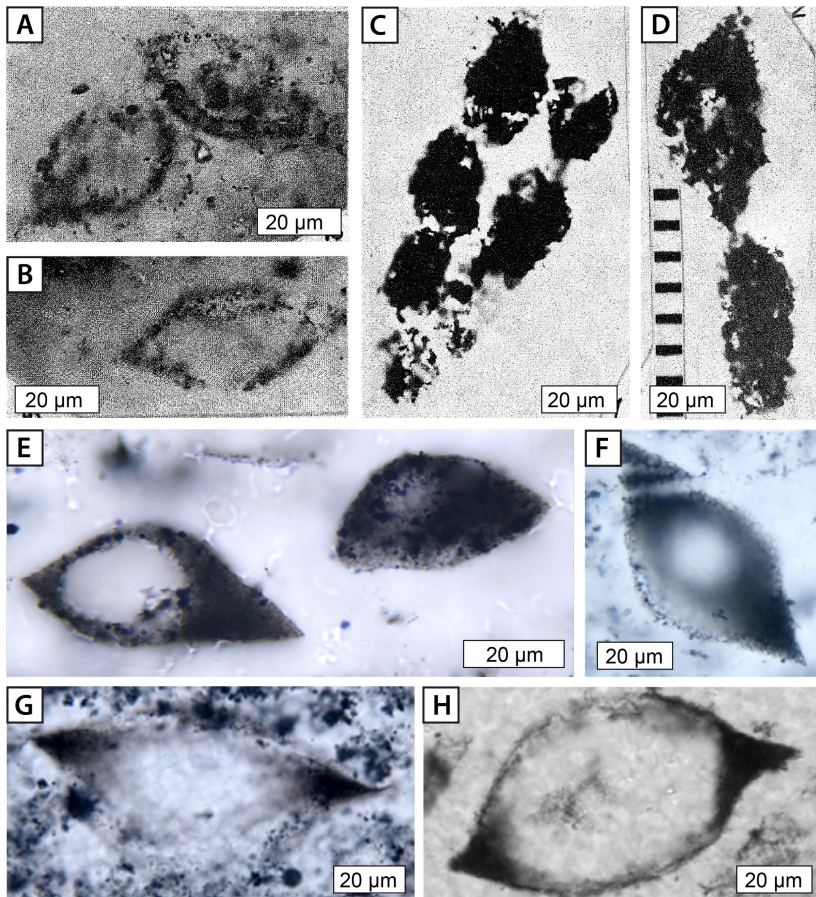
Spheroidal structures



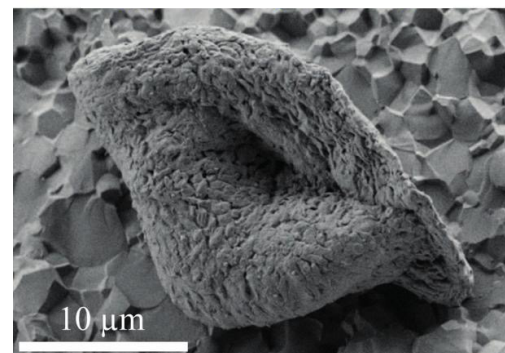
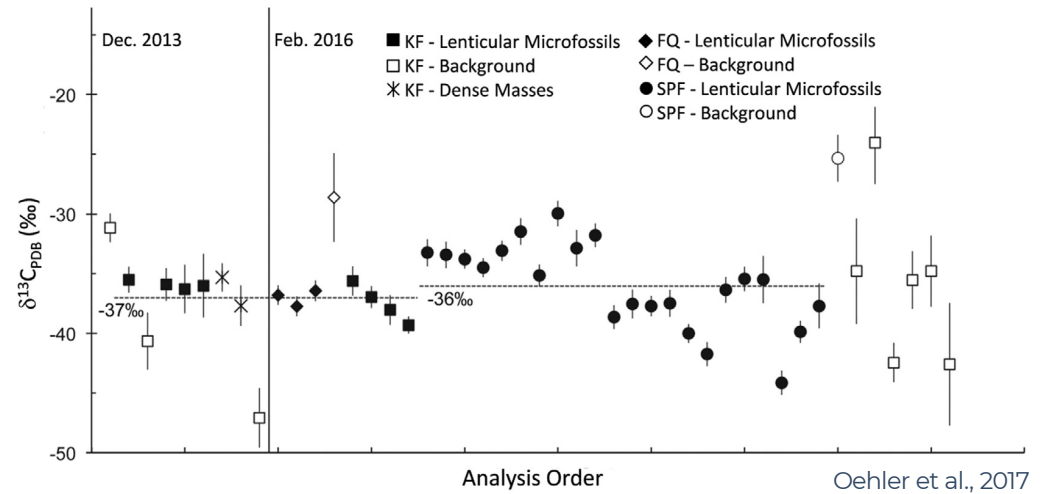
Glikson et al. (2008); Walsh (1992); Kremer and Kazmierczak (2017); Knoll and Barghoorn, 1977

Microfossils in the 3.4 Ga Buck Reef Chert

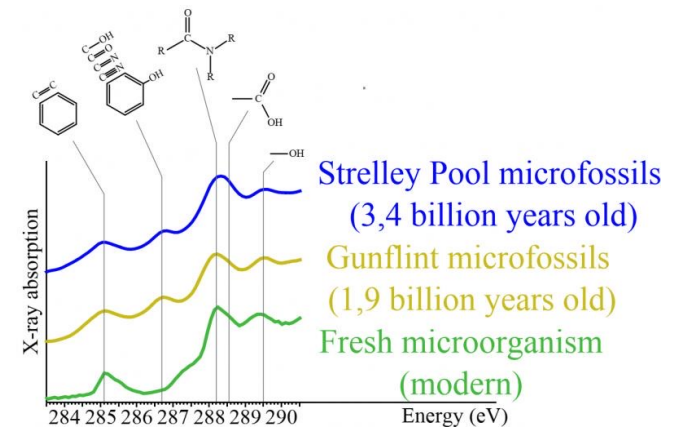
Lenticular microfossils



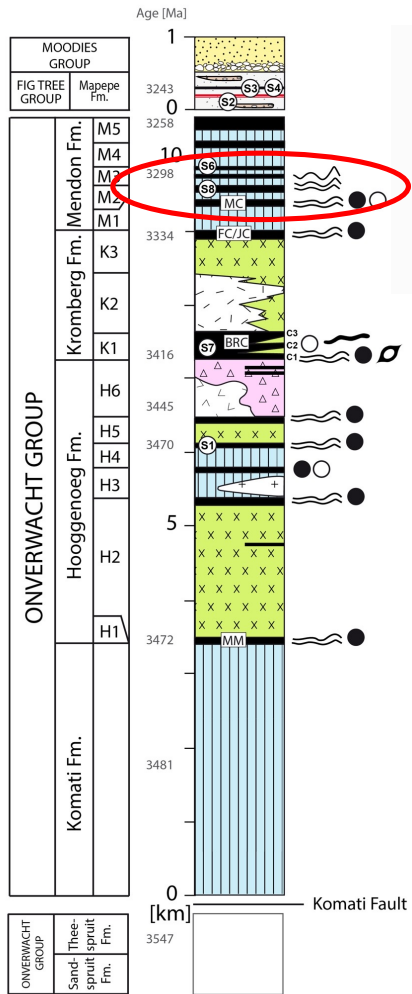
Pflug, 1966; Walsh 1992; Oehler et al., 2017



Alleon et al., 2018



Possible stromatolites in ~3.3 Ga cherts

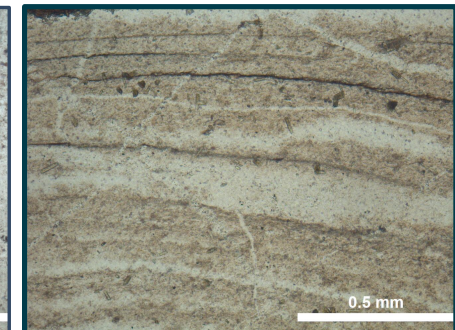
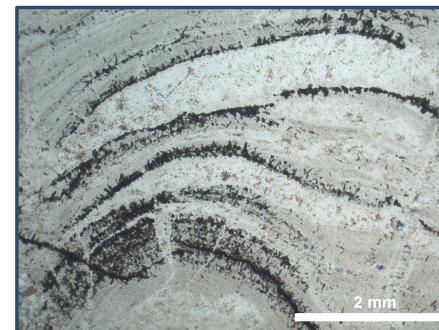
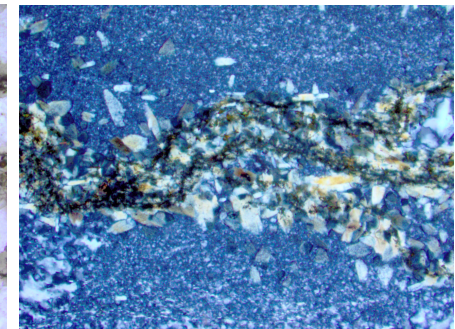
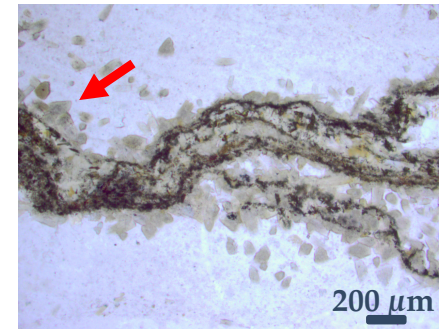
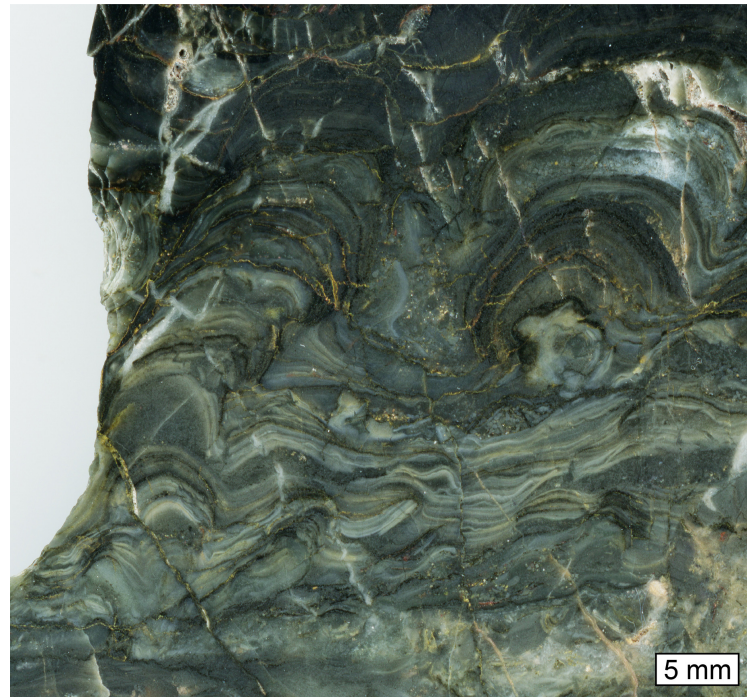
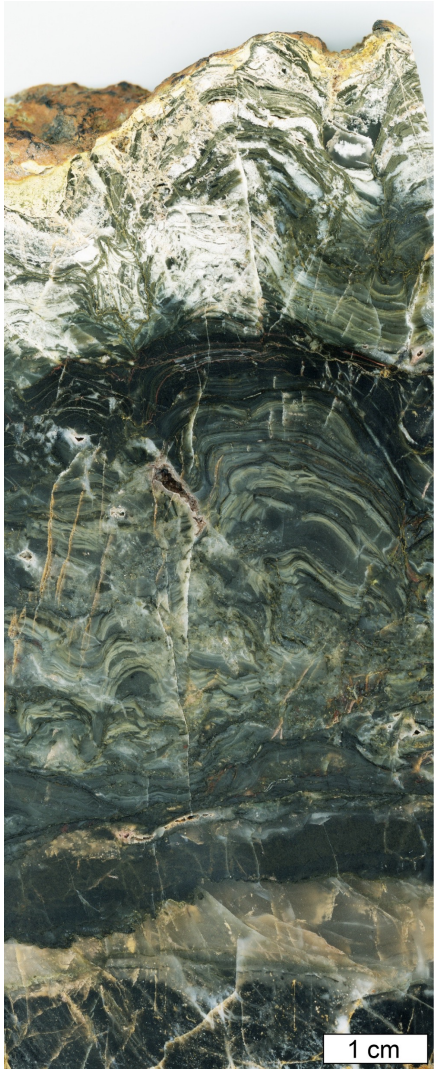


- Stratigraphic thickness: 1 - 20 cm within a 5 m thick chert layer
- Scattered outcrops for >10 km along strike

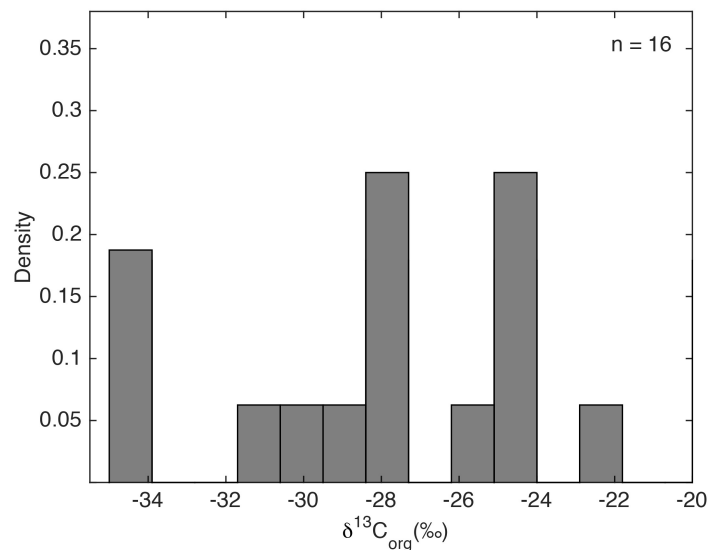


Stromatolites or silicious sinter deposits?

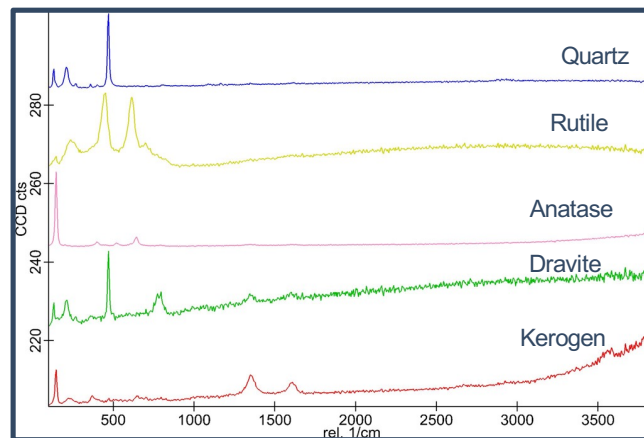
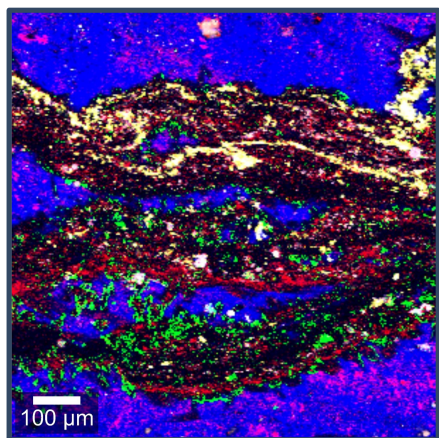
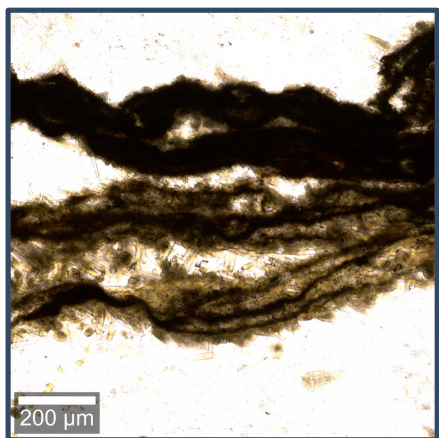
- Domal to pseudo columnar growth morphologies
- Crinkly, carbonaceous laminations contain Mg-Bo-Cr tourmaline
- Mineralization was likely driven by boron-rich hydrothermal fluids (Byerly et al., 1986; Lowe and Byerly, 2015)



Organic Carbon isotopes and Raman spectroscopy



- $\delta^{13}\text{C}_{\text{org}}$ values range between -34.5‰ and -22.1‰ (n=16)
- consistent with biogenic origin

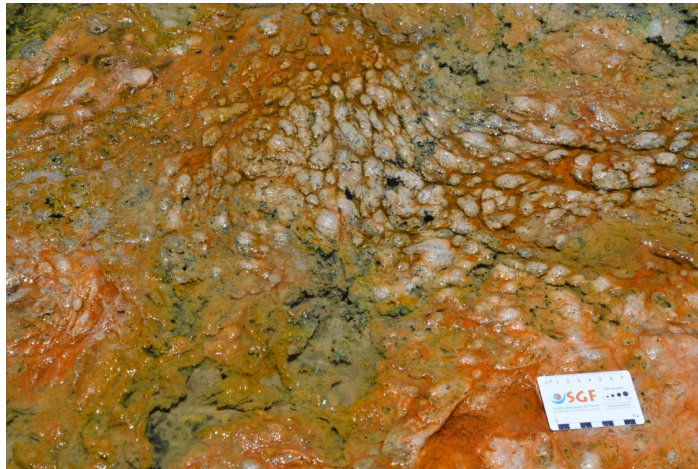


- Raman spectroscopy confirms presence of mature carbon
- T_{peak} : 440°C

El Tatio geothermal field, Chile

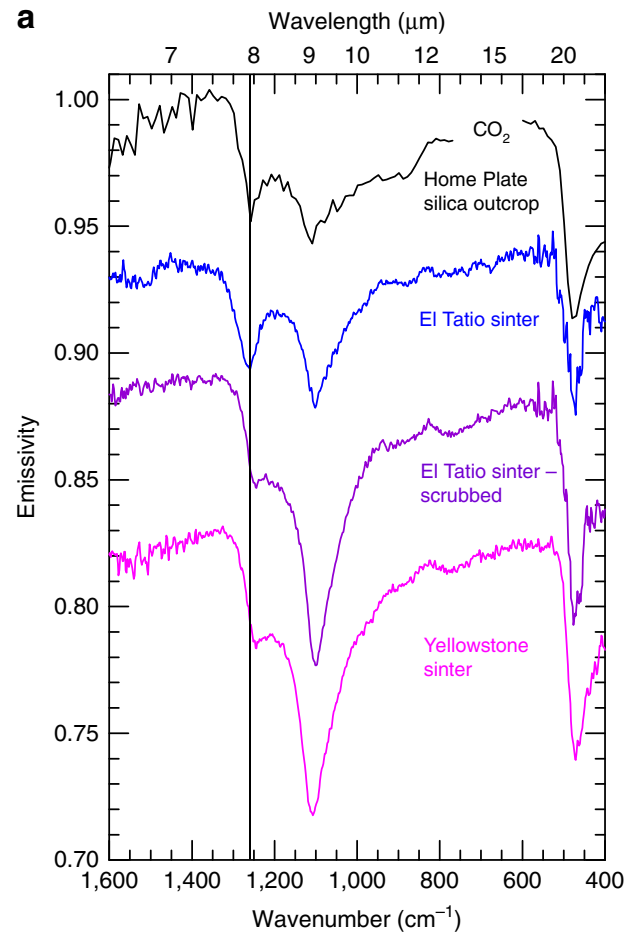
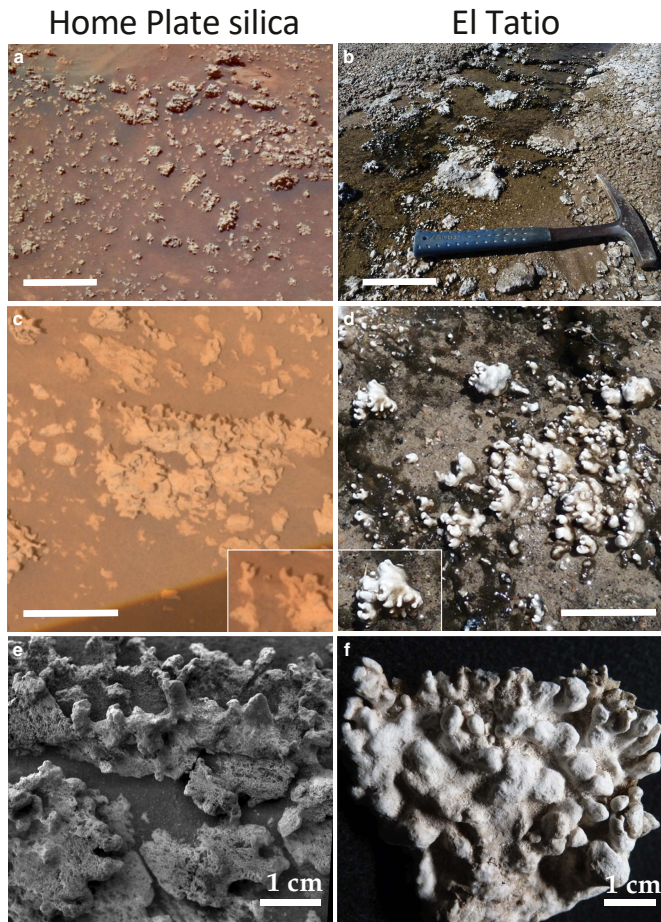


Microbial mats and evaporative silicification in hot spring silica sinters



Silica deposits on Mars with features resembling hot spring biosignatures at El Tatio in Chile

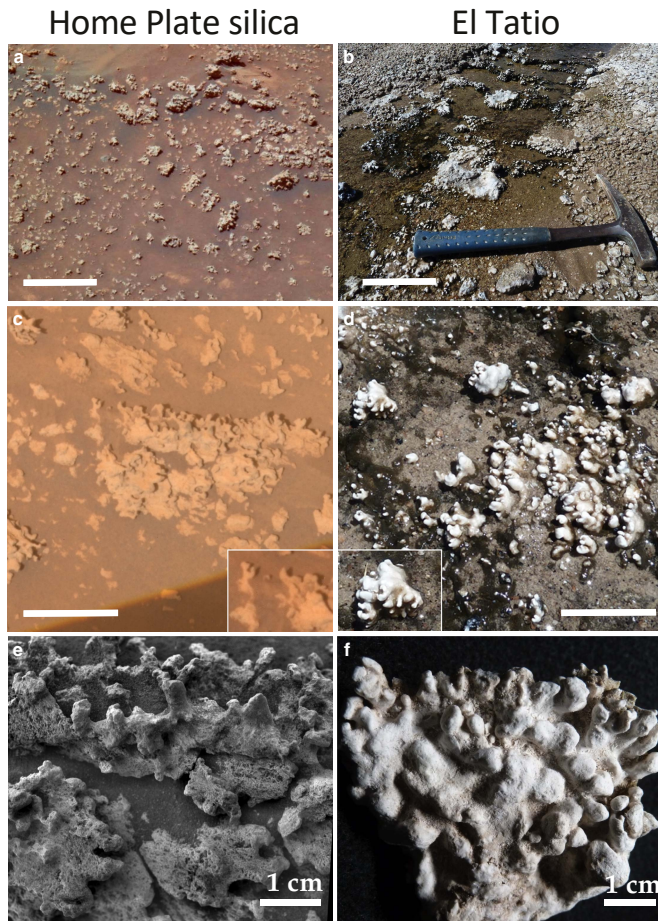
Steven W. Ruff¹ & Jack D. Farmer¹



Ruff and Farmer, 2016

Silica deposits on Mars with features resembling hot spring biosignatures at El Tatio in Chile

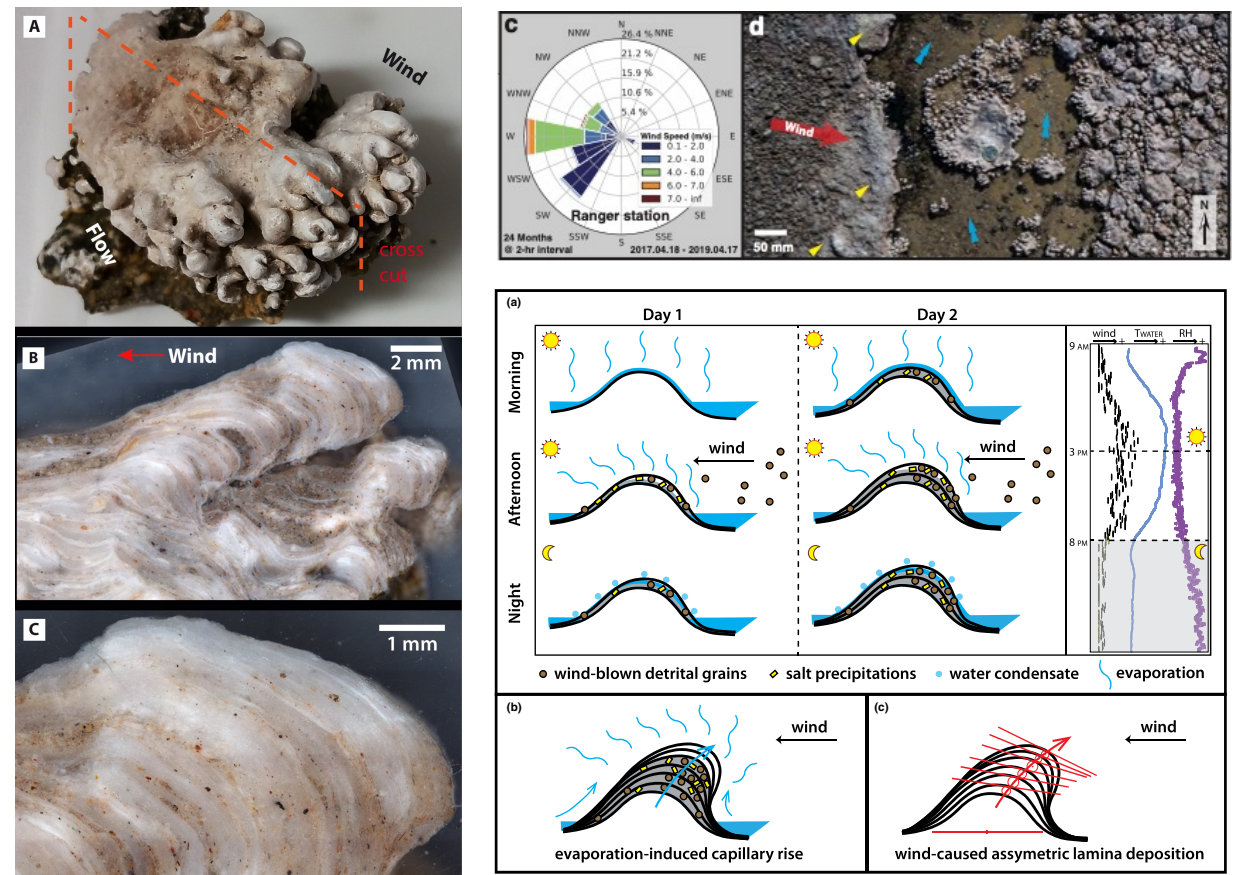
Steven W. Ruff¹ & Jack D. Farmer¹



Ruff and Farmer, 2016

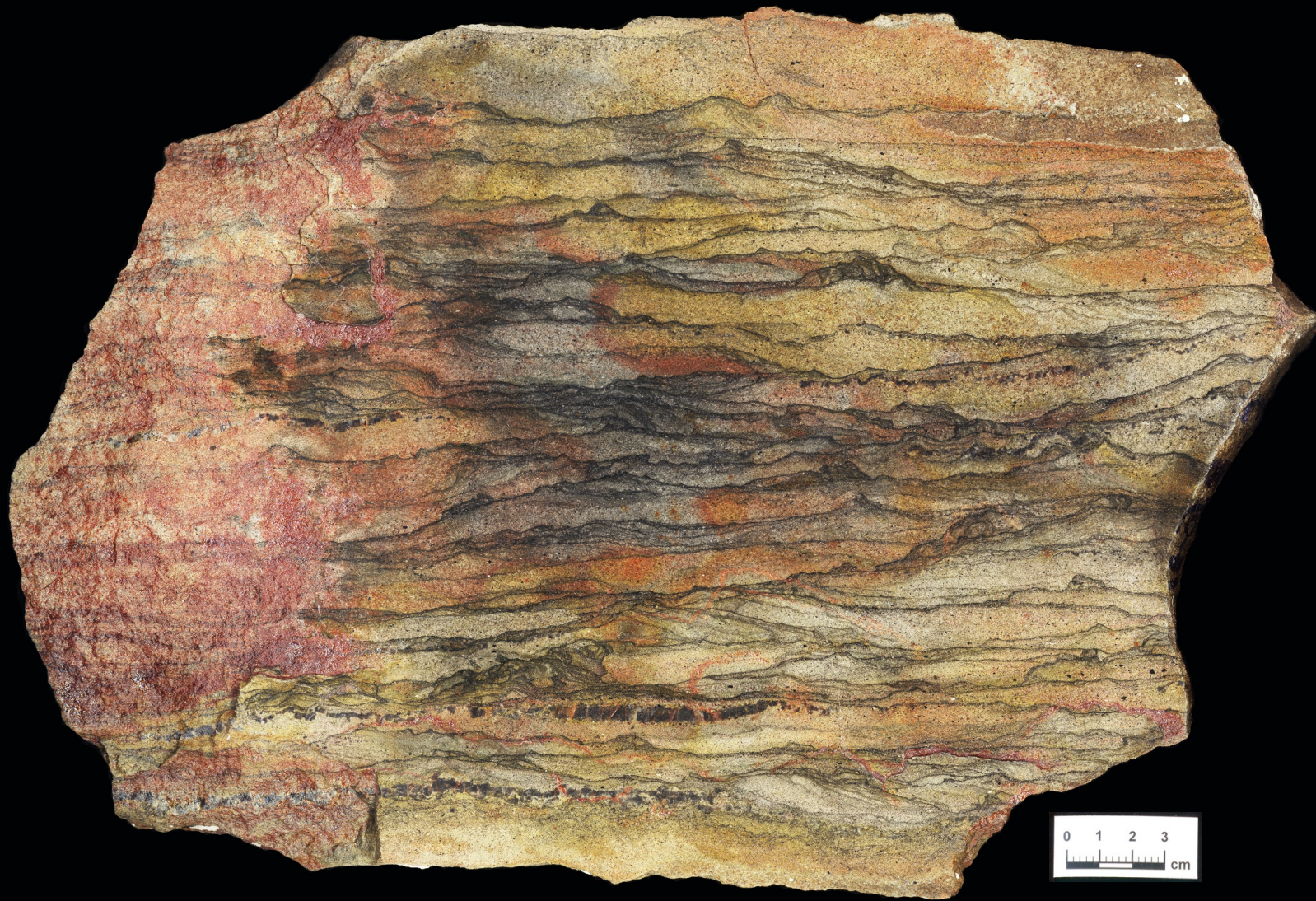
Morphogenesis of digitate structures in hot spring silica sinters of the El Tatio geothermal field, Chile

Jian Gong^{1,2} | Carolina Munoz-Saez³ | Dylan T. Wilmeth^{1,4} | Kimberly D. Myers¹ | Martin Homann⁵ | Gernot Arp⁶ | John R. Skok⁷ | Mark A. van Zuilen¹



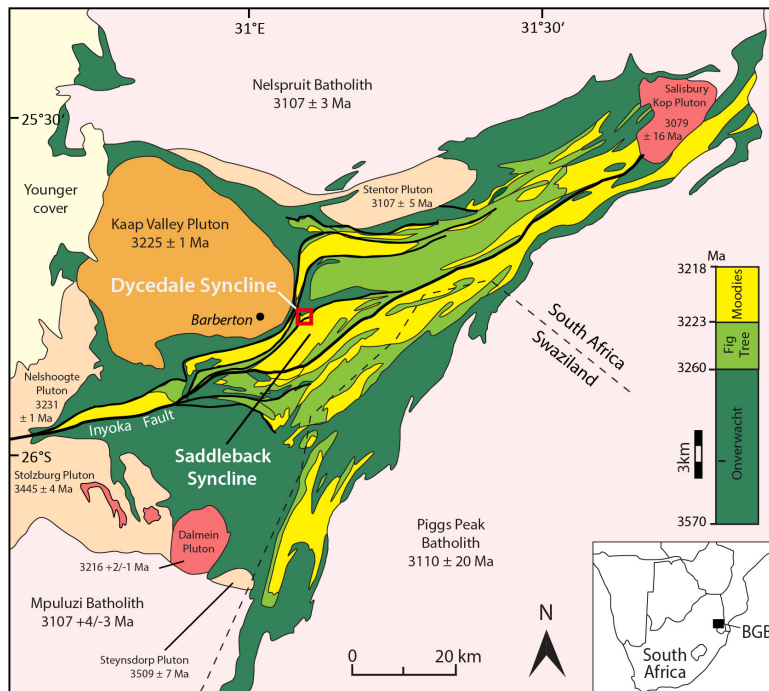
Gong et al., 2021

Microbial mats in the 3.22 Ga old Moodies Group

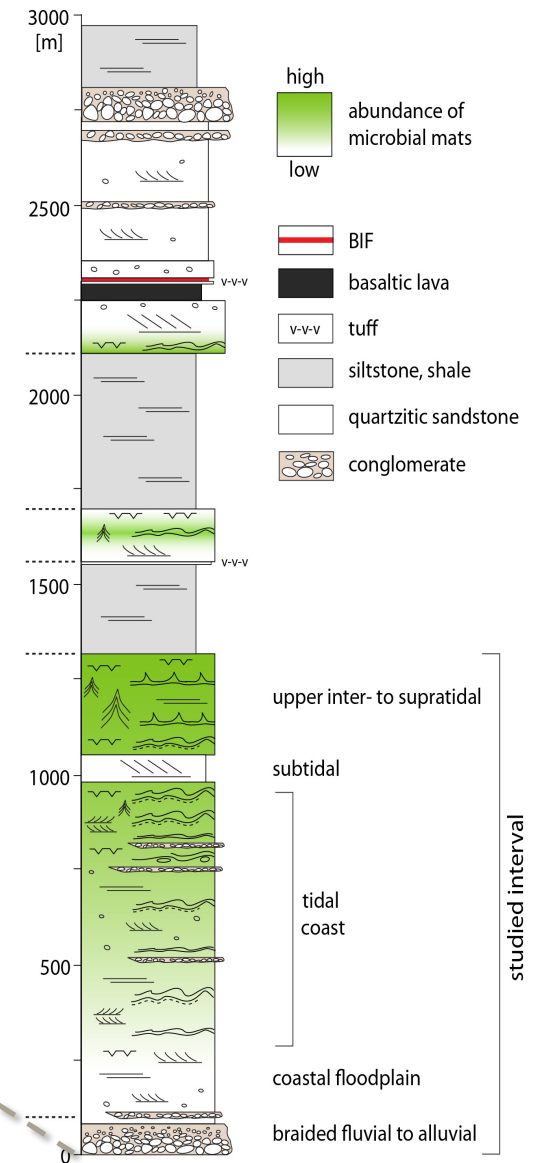


The 3.22 Ga Moodies Group

- Earth's oldest well-preserved continental to marine transition
- Very-high-resolution record (3.6 km deposited over 1 to 14 Myr)
- Paleoenvironment: alluvial to deltaic settings, with a dominance of coastal plains and tidal deltas



Homann et al., 2015; 2016





Fossil microbial mats...

fossil mats

sandstone

fossil mats



...formed a primary microrelief

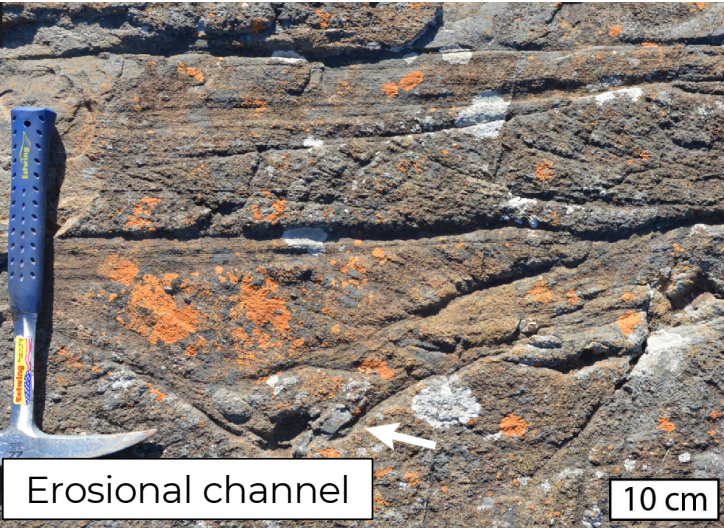
fossil mats



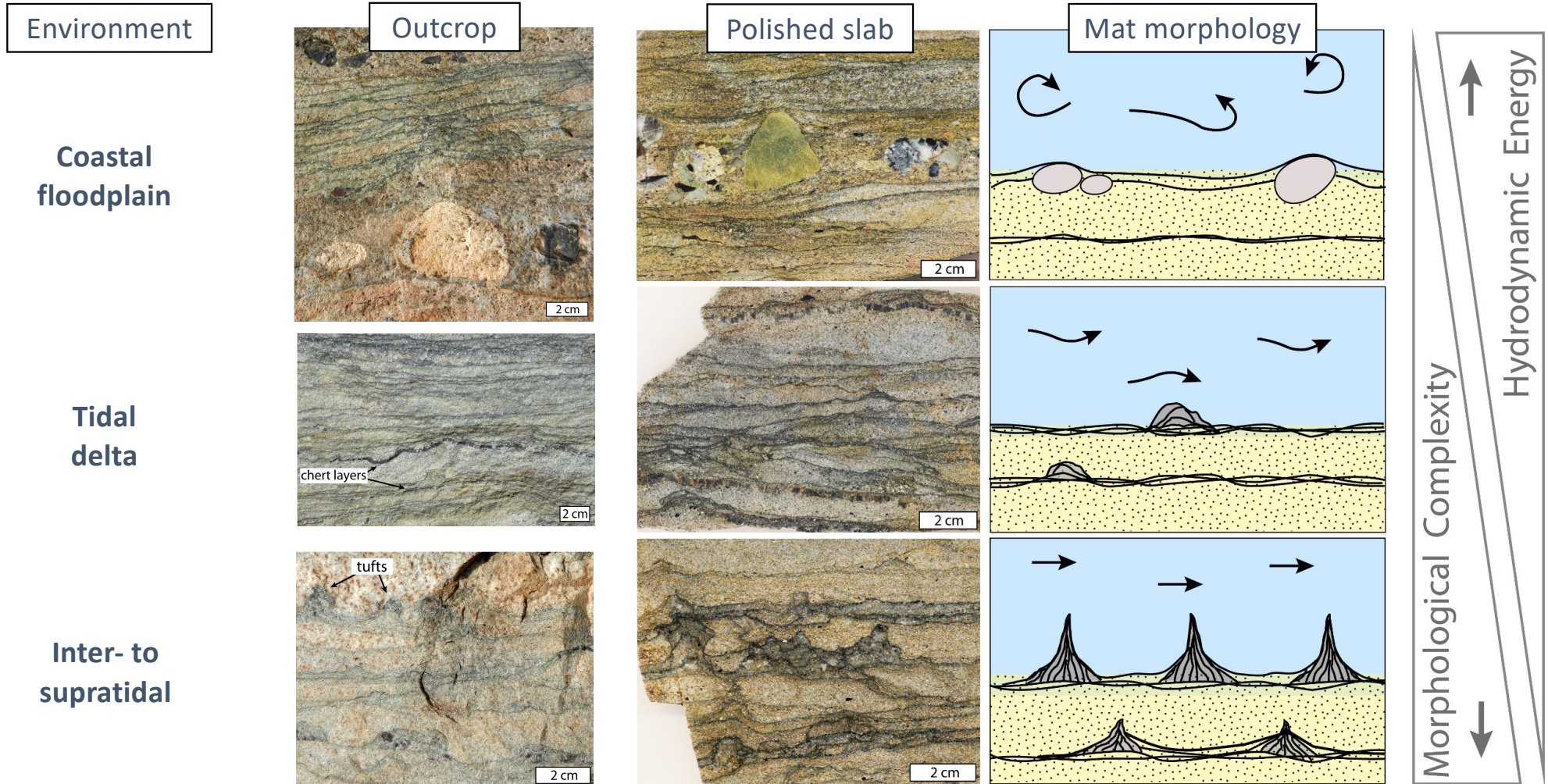
sand ripple



Exceptionally well-preserved sedimentary structures



Morphological adaptation to different hydrodynamic settings

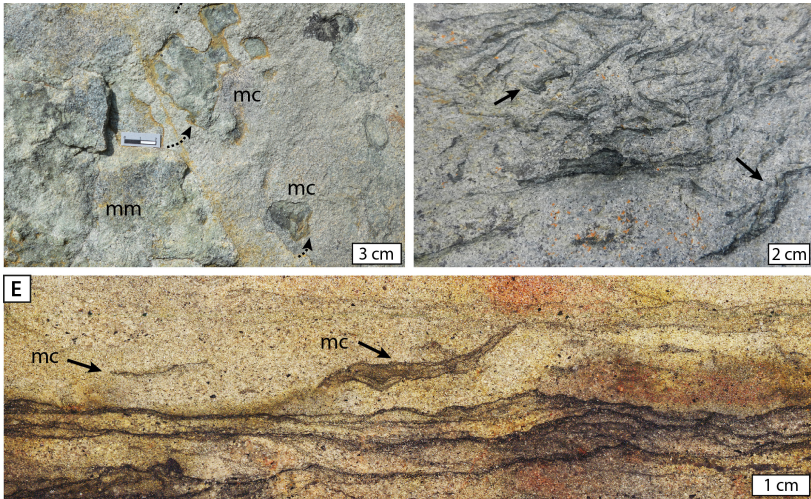


Microbially induced sedimentary structures

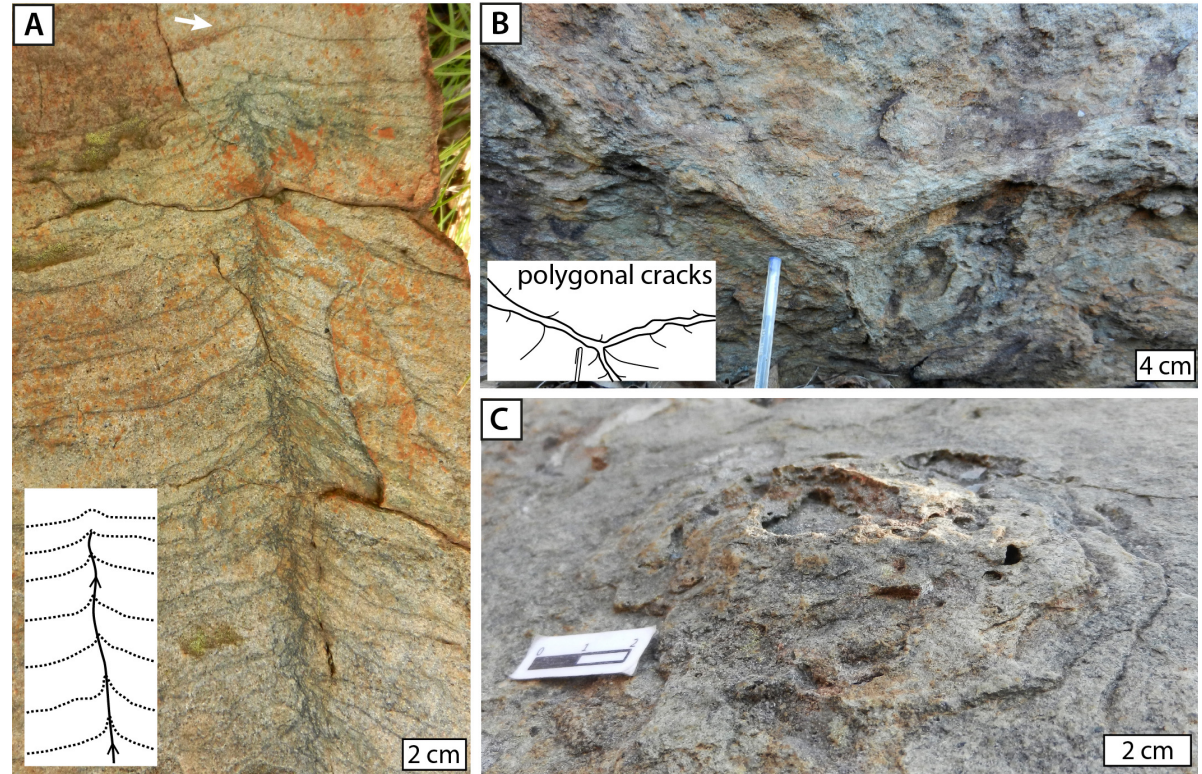
Shrinkage cracks



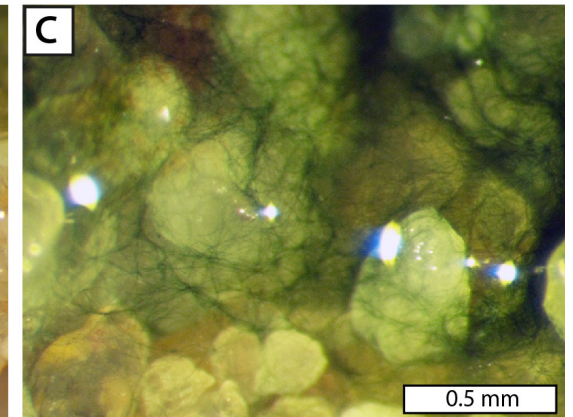
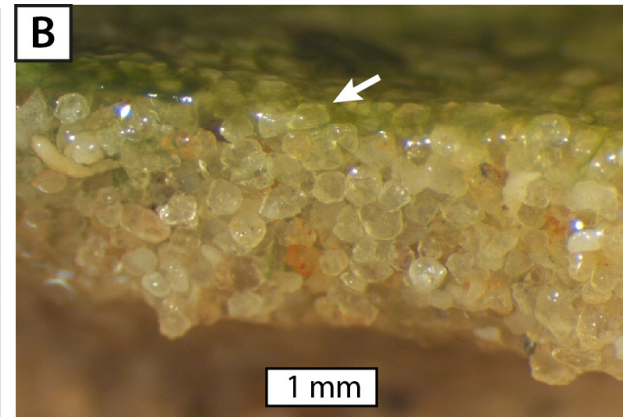
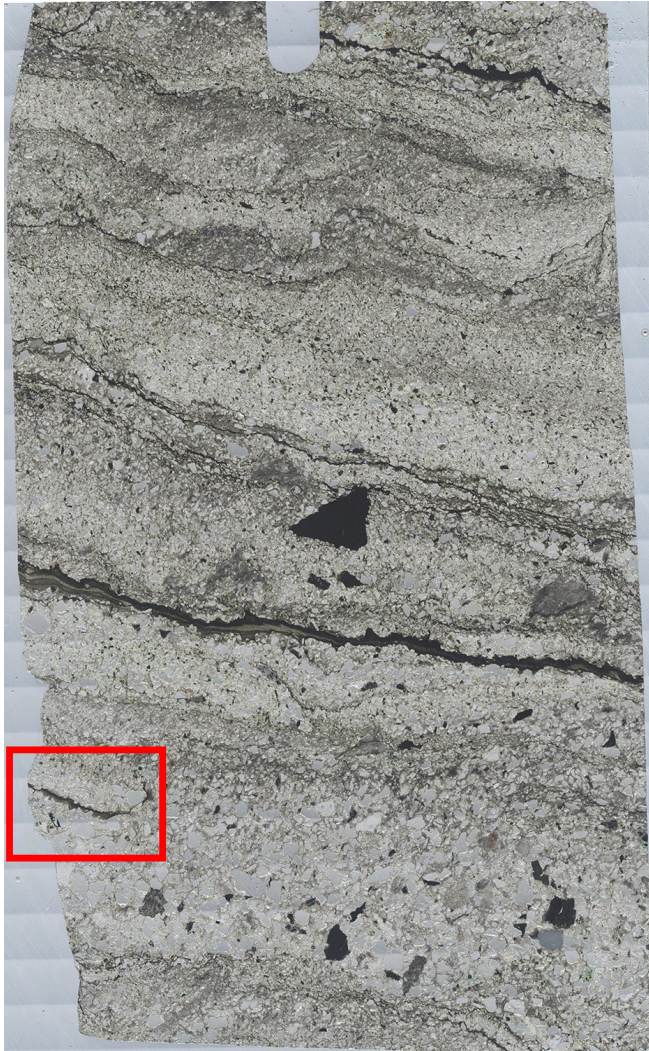
Mat fragments



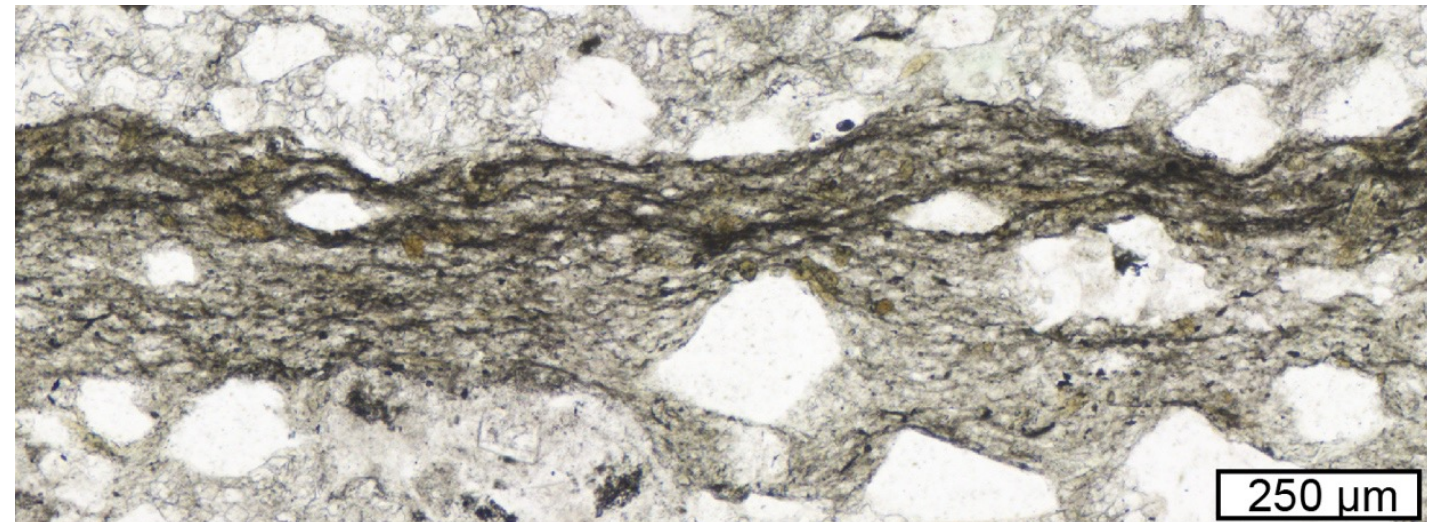
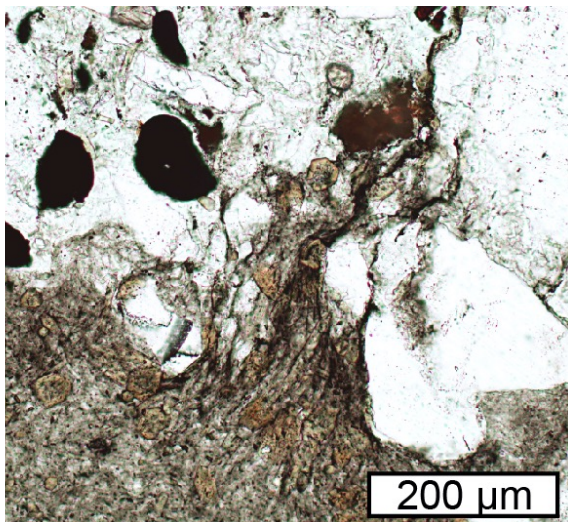
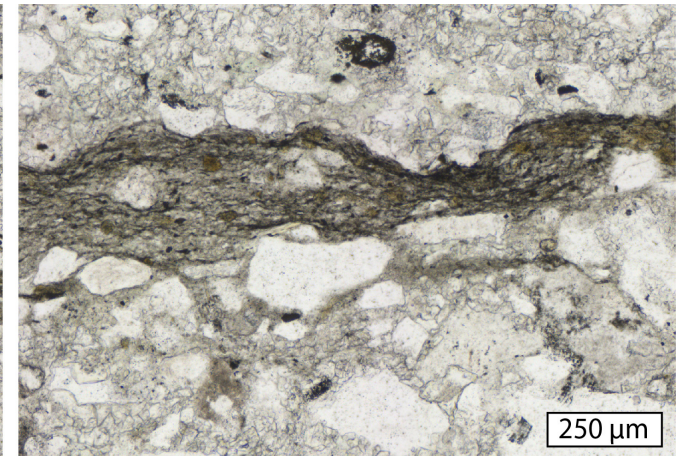
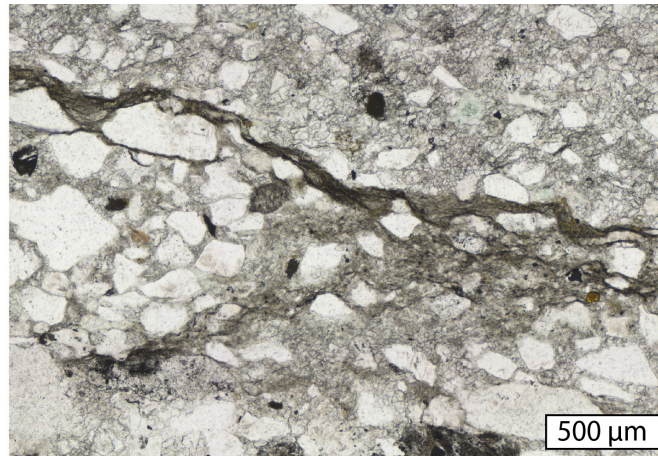
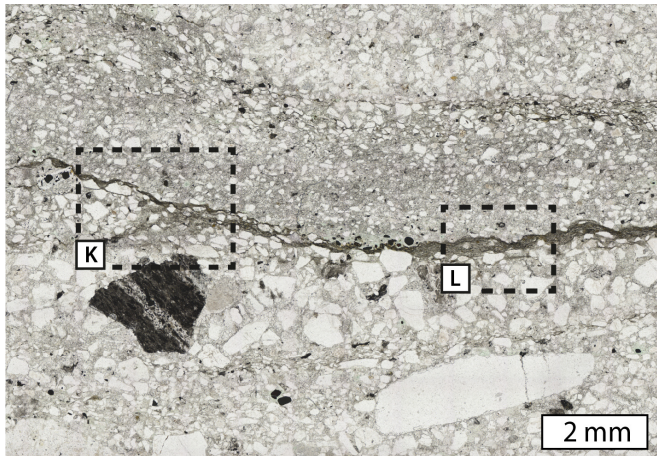
Fluid-escape structures



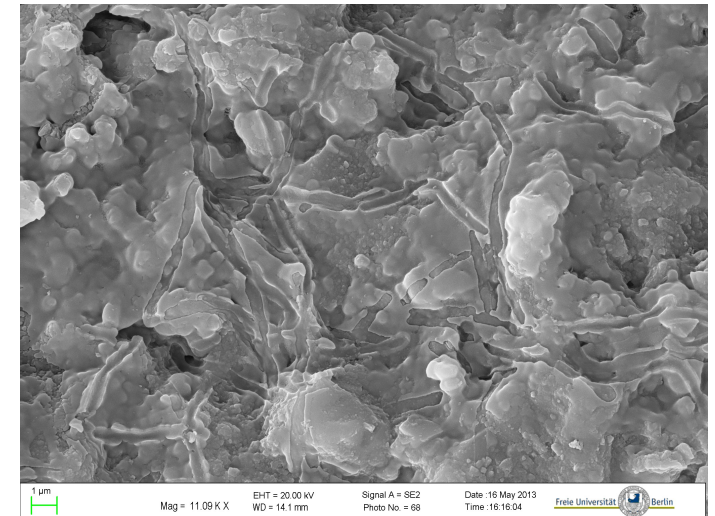
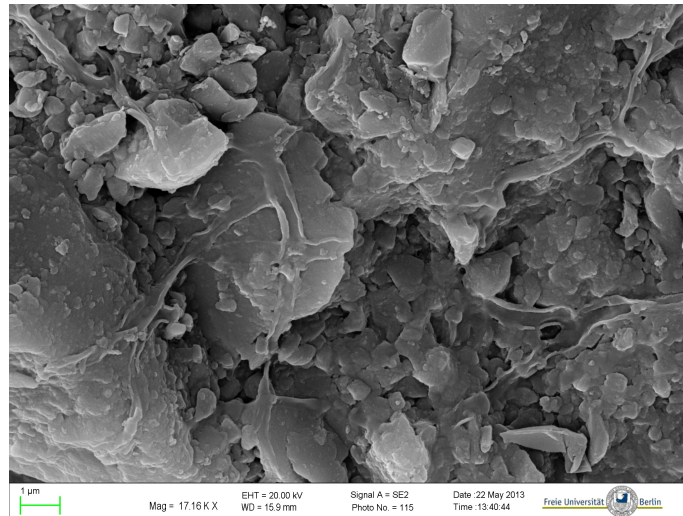
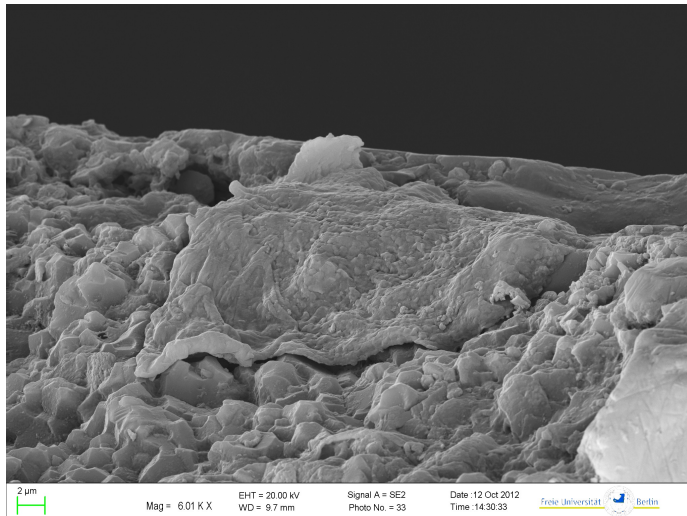
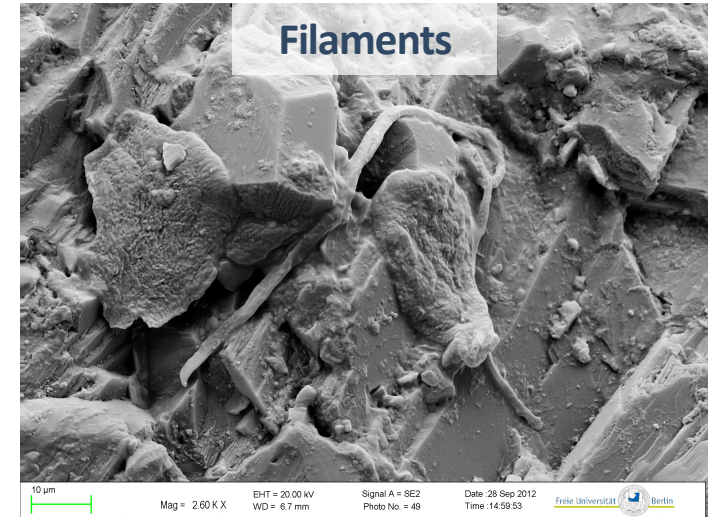
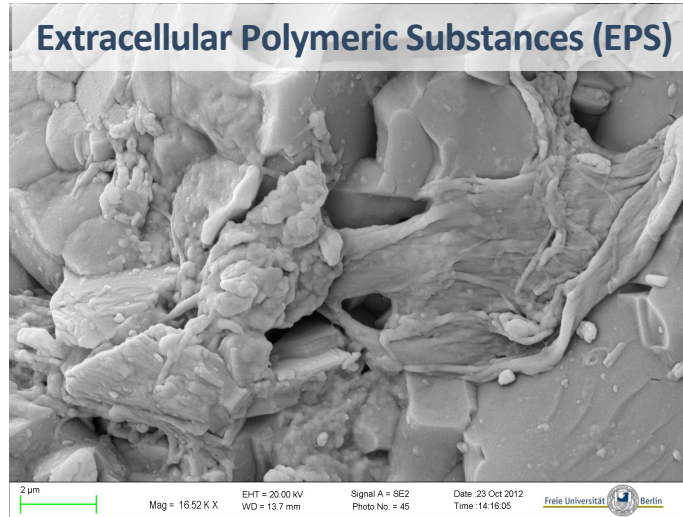
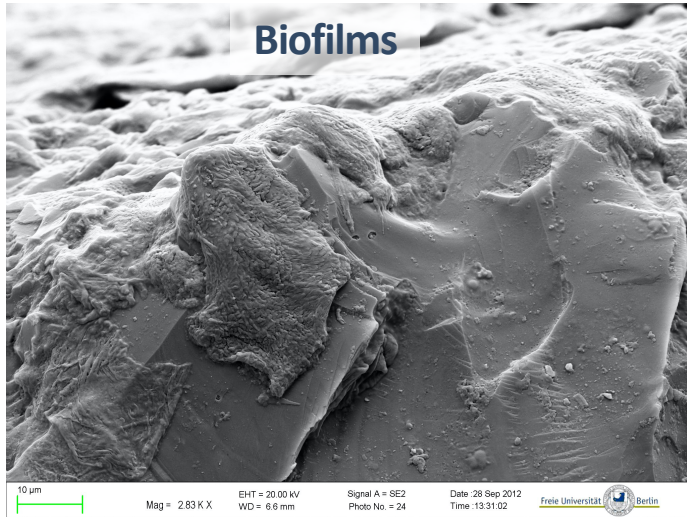
Fossil mats are preserved as organic-rich laminations



Laminae composed of filamentous microstructures

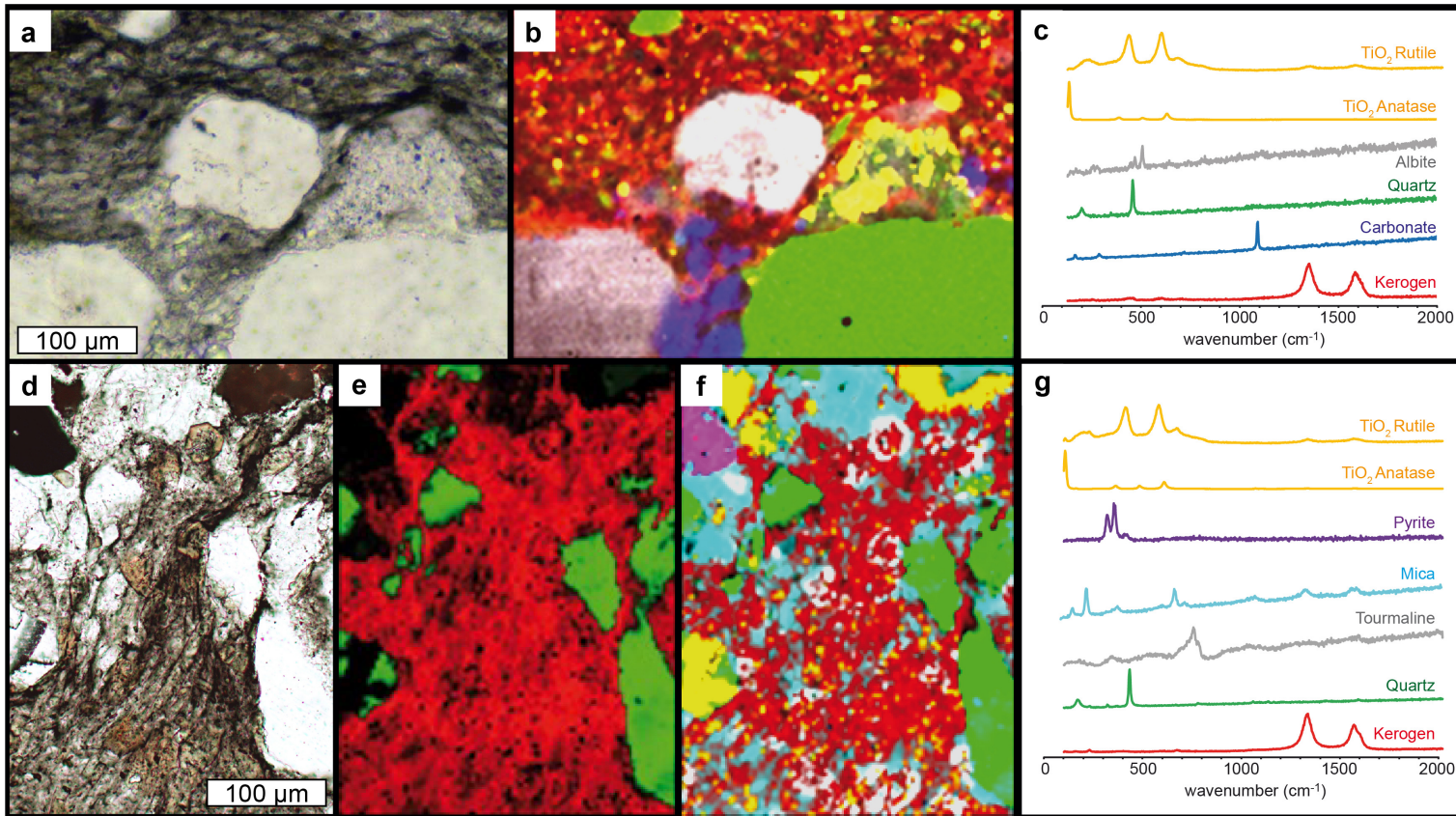


Scanning electron microscopy analysis

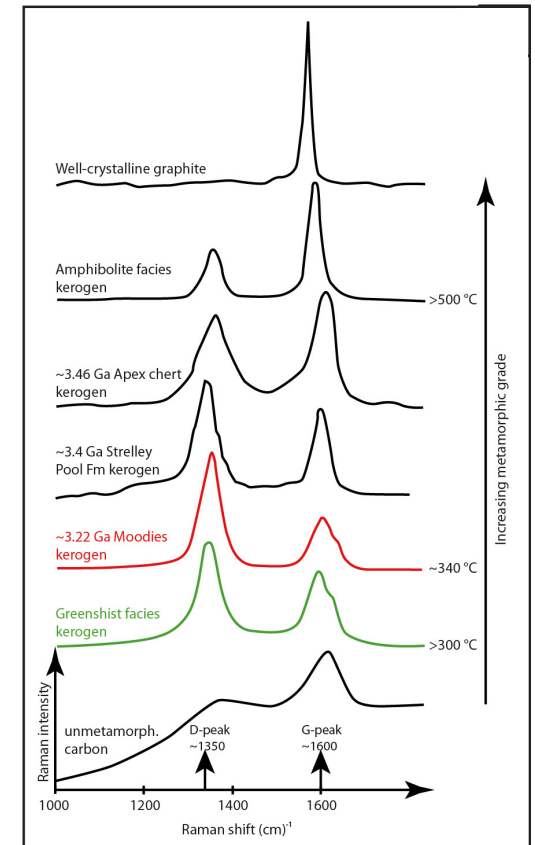


Raman analysis confirms:

(1) Carbonaceous composition of the laminae



(2) Syngenicity of the kerogen



Comparison of marine and fluvial mats

Marine facies



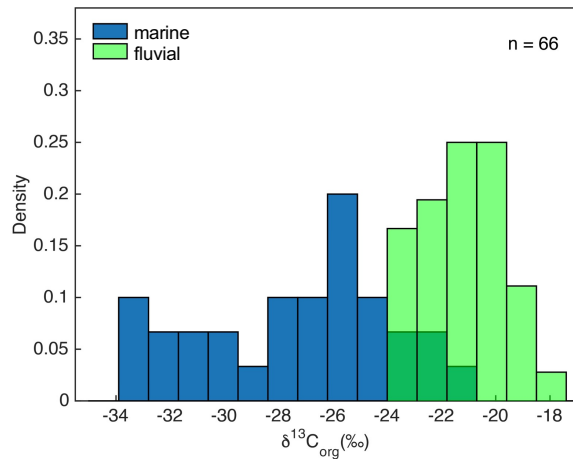
Fluvial facies



- ❖ Interbedded with medium- to coarse-grained sandstones
- ❖ Crinkly and tufted growth morphologies

- ❖ Interbedded with pebbly sandstones and conglomerates
- ❖ Mostly planar and generally thicker mats

Organic carbon and nitrogen isotope analysis

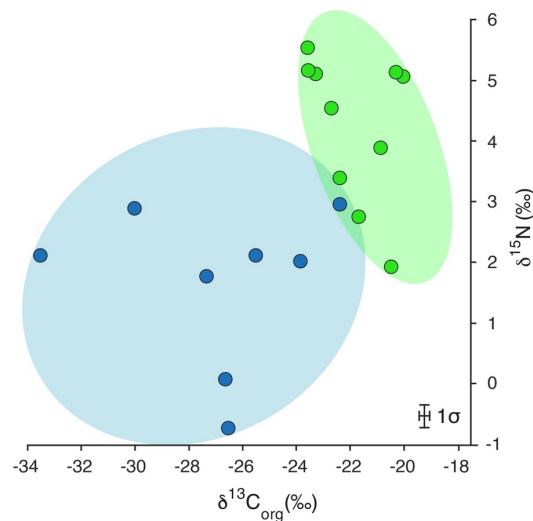


Fluvial mats: $-24 < \delta^{13}\text{C} < -18\text{‰}$

❖ consistent with autotrophic carbon fixation via the *Calvin-Benson cycle*

Marine mats: $-34 < \delta^{13}\text{C} < -21\text{‰}$

❖ best explained with carbon fixation via *Wood-Ljungdahl pathway*, which includes methanogens and sulfate reducers



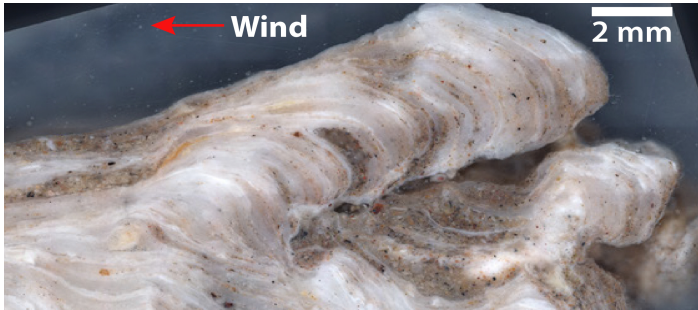
Fluvial mats: $+2 < \delta^{15}\text{N} < +5\text{‰}$

Marine mats: $0 < \delta^{15}\text{N} < +3\text{‰}$

❖ Biological mechanisms that produce biomass with $\delta^{15}\text{N} > +2\text{‰}$:

- 1) Partial assimilation of ammonium (NH_4^+)
- 2) Partial nitrification
- 3) Partial denitrification

Summary



- ❖ Multiple, independent lines of evidence are needed to establish biogenicity; morphology alone is not enough.



- ❖ Early diagenetic mineralization is key for the preservation and survivability of biosignatures.



- ❖ Silicified microbial mats are some of the most robust evidence for early life on Earth.



m.homann@ucl.ac.uk

@matworlds

Caltech

Acknowledgements

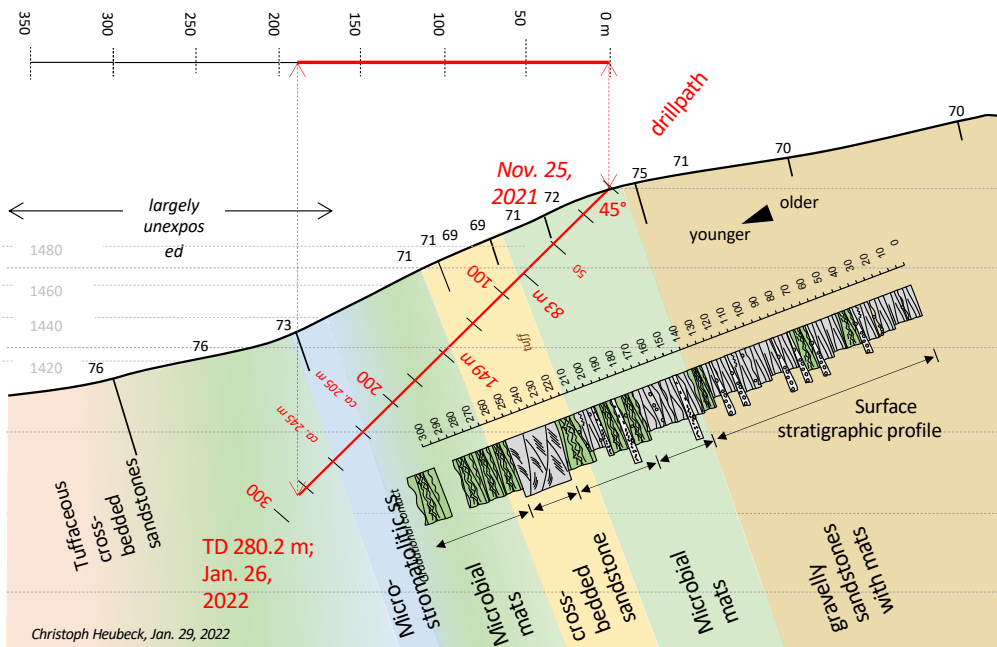
Christoph Heubeck
Stefan Lalonde
Pierre Sansjofre
Tomaso Bontognali
Mark Van Zuilen
Jian Gong
Mike Tice
Dylan Wilmeth
Alessandro Airo
Bryan Killingsworth
Raphael Baumgartner
Martin Van Kranendonk



ETH zürich



ICDP drilling project BASE



Site 3 (Saddleback Syncline): Post-drill geological cross section

Drill core diameter:

6.1 cm



Mars 2020

1.3 cm

