

Climate and Radiation

Yuk Ling Yung

GPS Caltech

Presentation at Caltech

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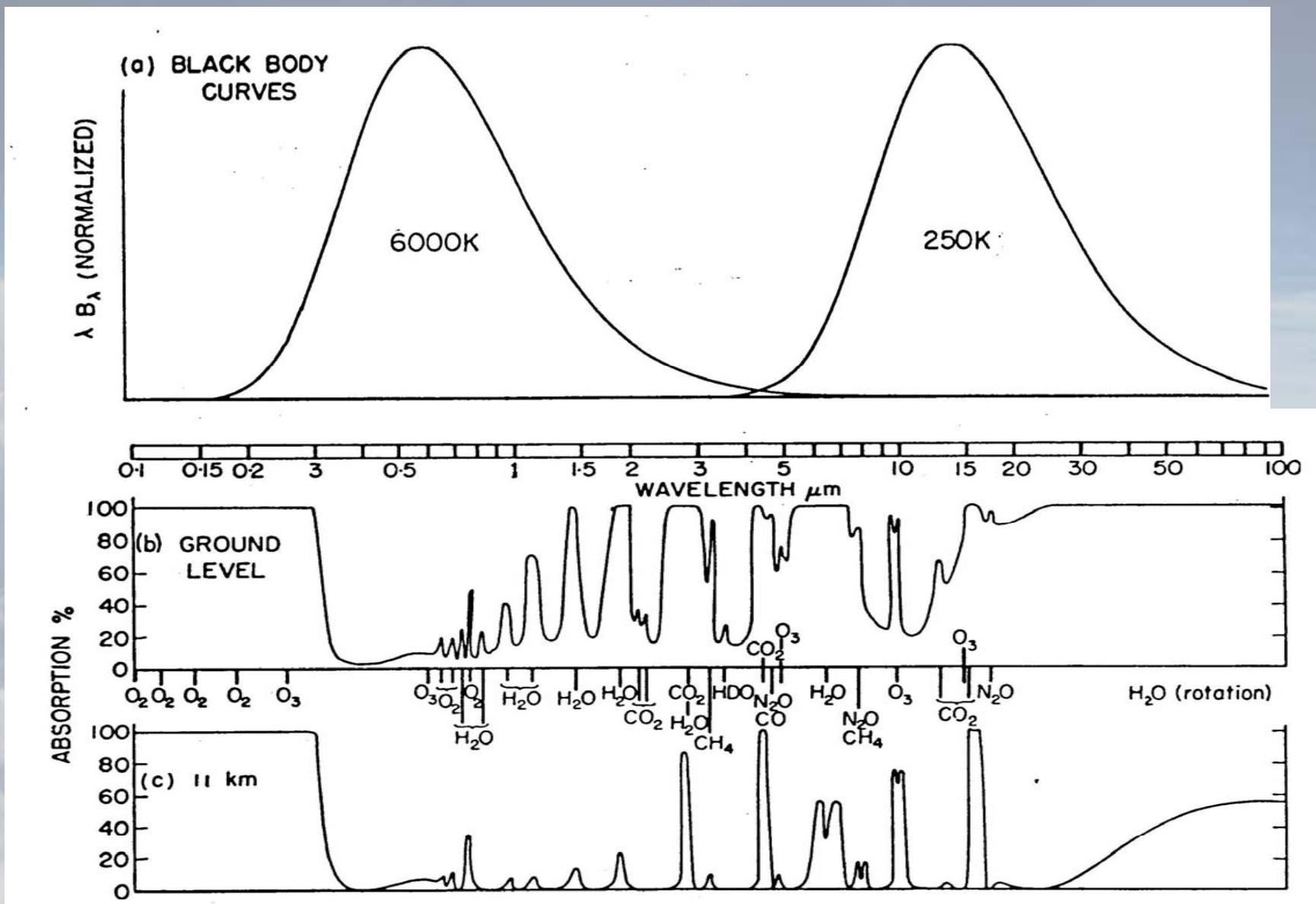
Today's Outline

- ❖ What is Radiation? (Review)
- ❖ What is the Greenhouse Effect? (Review)
- ❖ GCM Prediction of Future (IPCC)
- ❖ Use of Spectral Data
- ❖ Conclusions

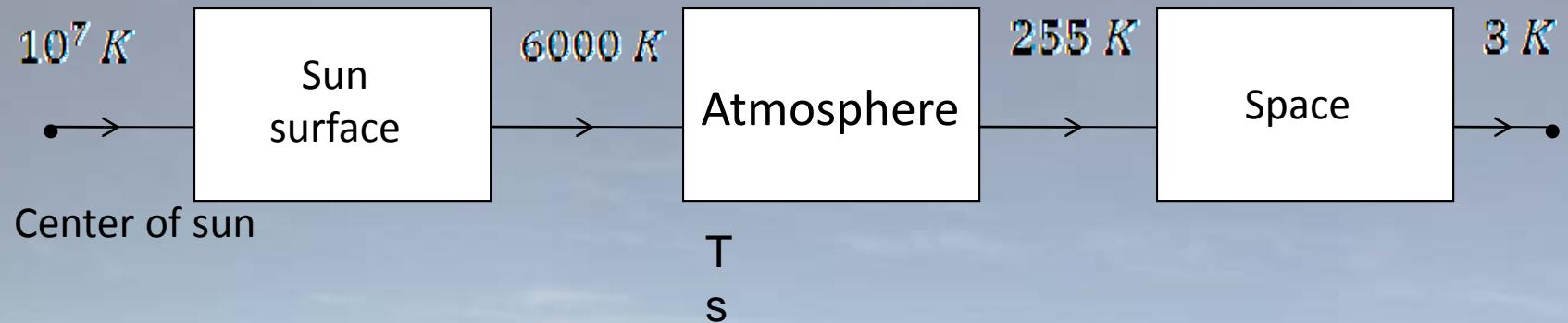
Blackbody Radiation

$$B_{\nu}(T) = \frac{2h\nu^3}{c^2} \frac{1}{e^{h\nu/kT} - 1}$$





Goody and Yung (1995)



Atmosphere : Processor of energy

2
8
Sunlight = driver of 8 { Chemistry
K dynamics
 biology

$$\overline{f_{\odot}} = 350 \text{ W} \cdot \text{m}^{-2}$$

Mean equilibrium temperature

$$f(\pi r^2)(1 - \alpha) = \sigma T^4(4\pi r^2)$$

$$T = \left(\frac{(1 - \alpha)f}{4\sigma} \right)^{1/4}$$

$$f = 1400 \text{ W/m}^2$$

$$\alpha = 0.31$$

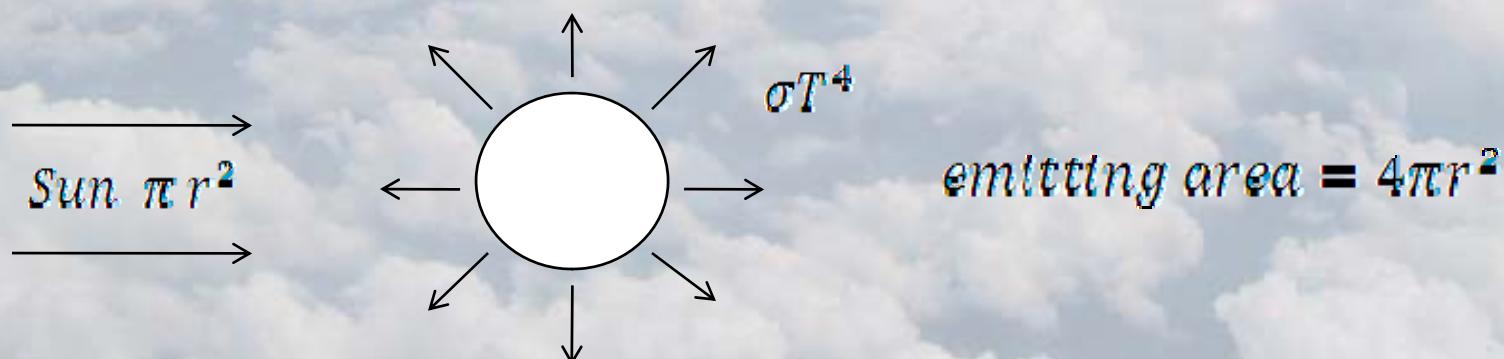
$$\sigma = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$$

(Appendix 1, p. 462, G + Y)

$$T = \left(\frac{(0.69)(1400)}{5.67 \times 10^{-8}} \right)^{1/4} = 255 \text{ K}$$

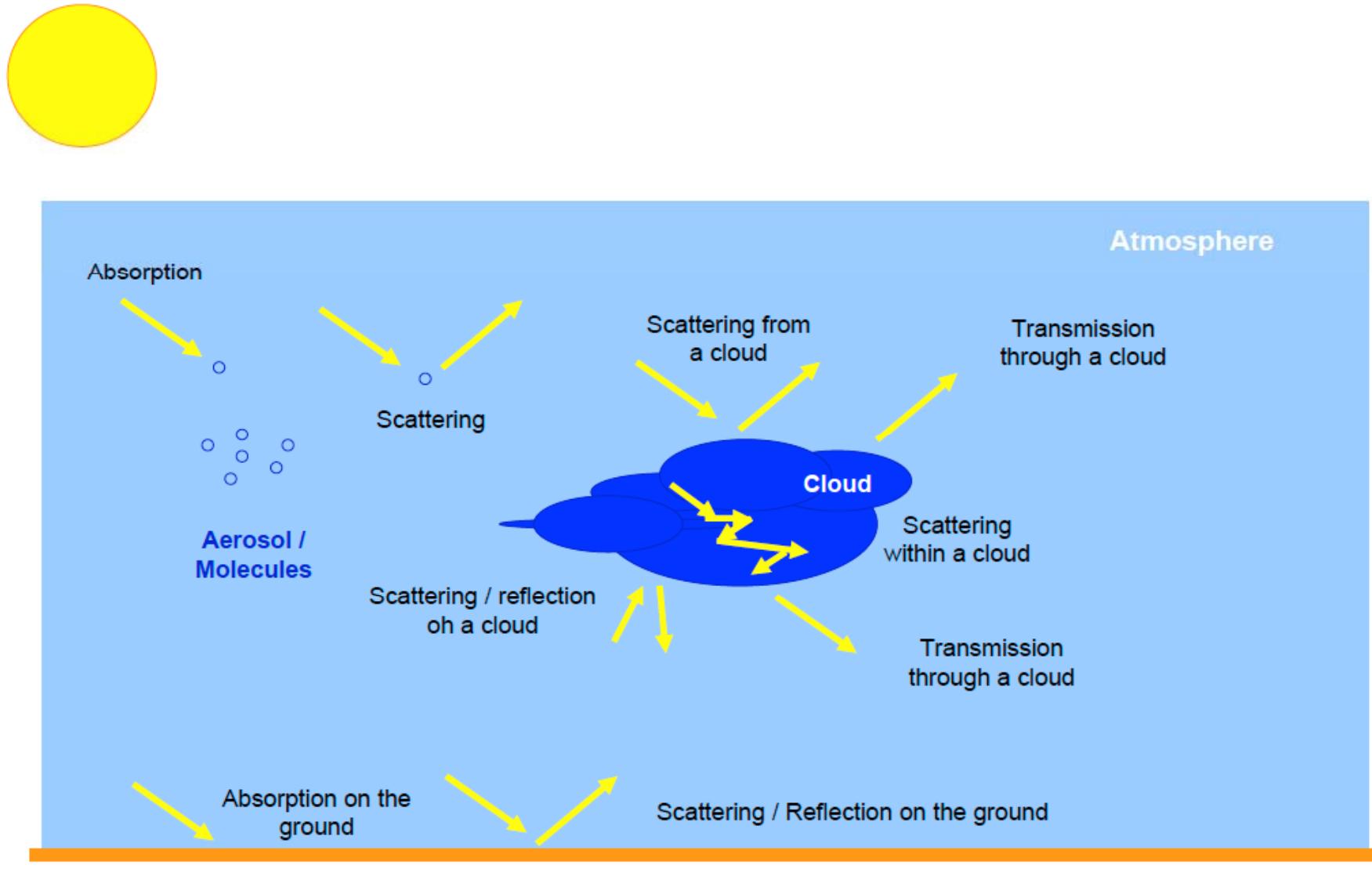
$$Ts = 288 \text{ K}$$

$$\therefore \Delta T = 33 \text{ K}$$

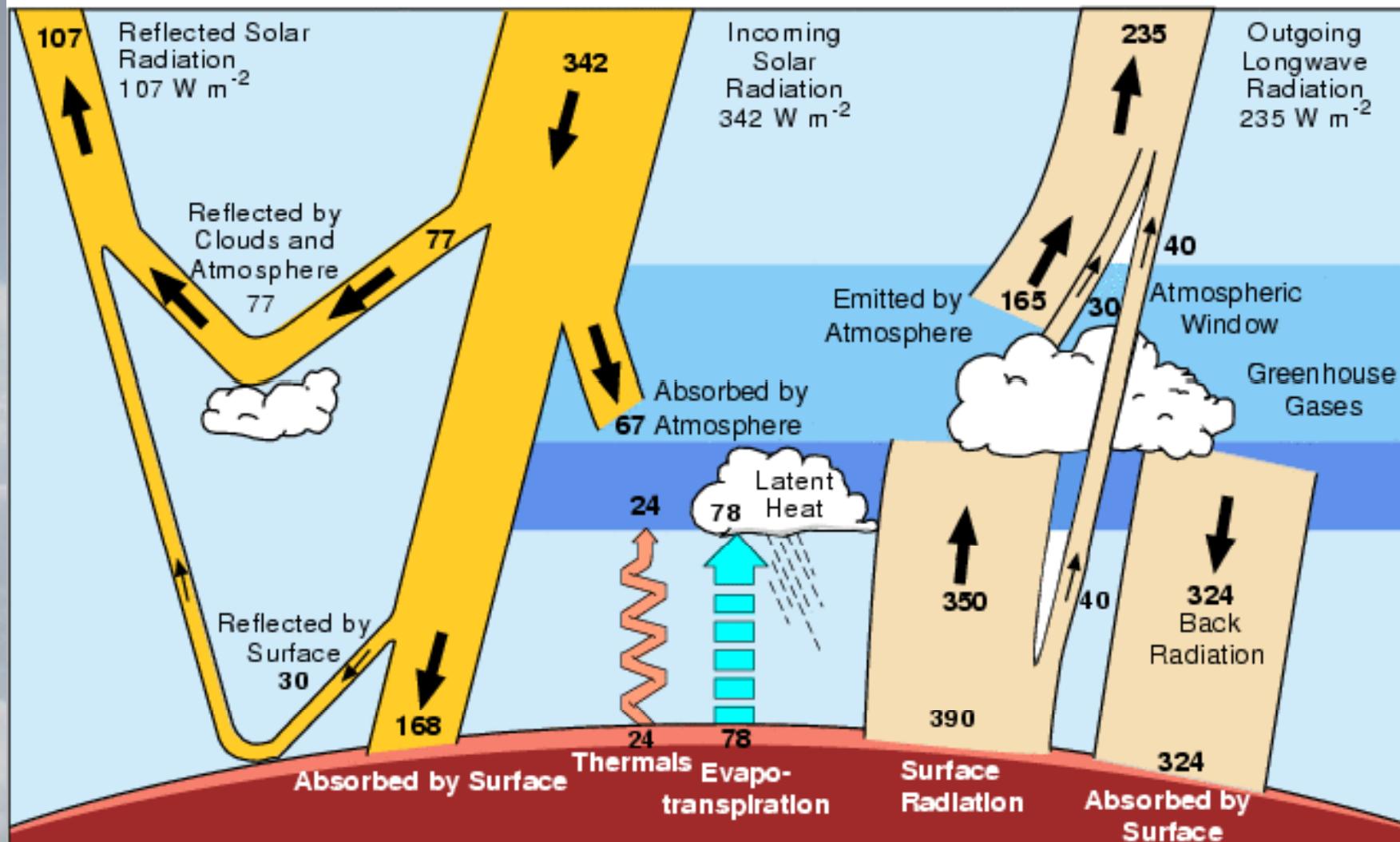


<u>Process</u>	<u>Total Power (Watts)</u>
Sunlight	2×10^{17}
Seawater evaporation	3×10^{16}
Photosynthesis	1×10^{14}
Geothermal flux	3×10^{13}
Tides	2×10^{13}
Energy use by man (1980)	1×10^{13}
UV energy to make O ₃ layer	1×10^{13}
Hydroelectric Power	3×10^{11}
Winds	2×10^{11}
Volcanoes	4×10^{10}
Earthquakes	1×10^{10}

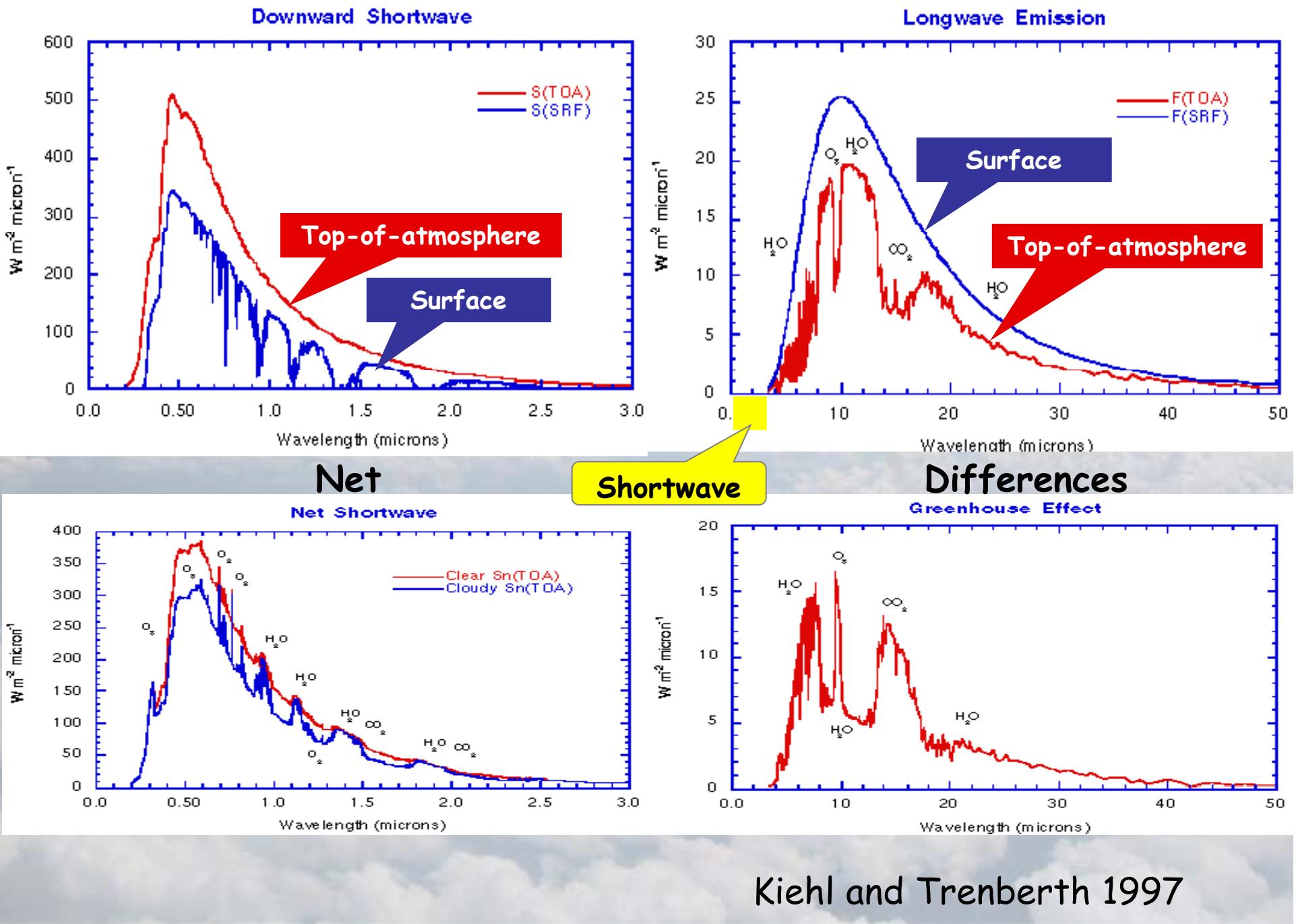
Radiative Transfer

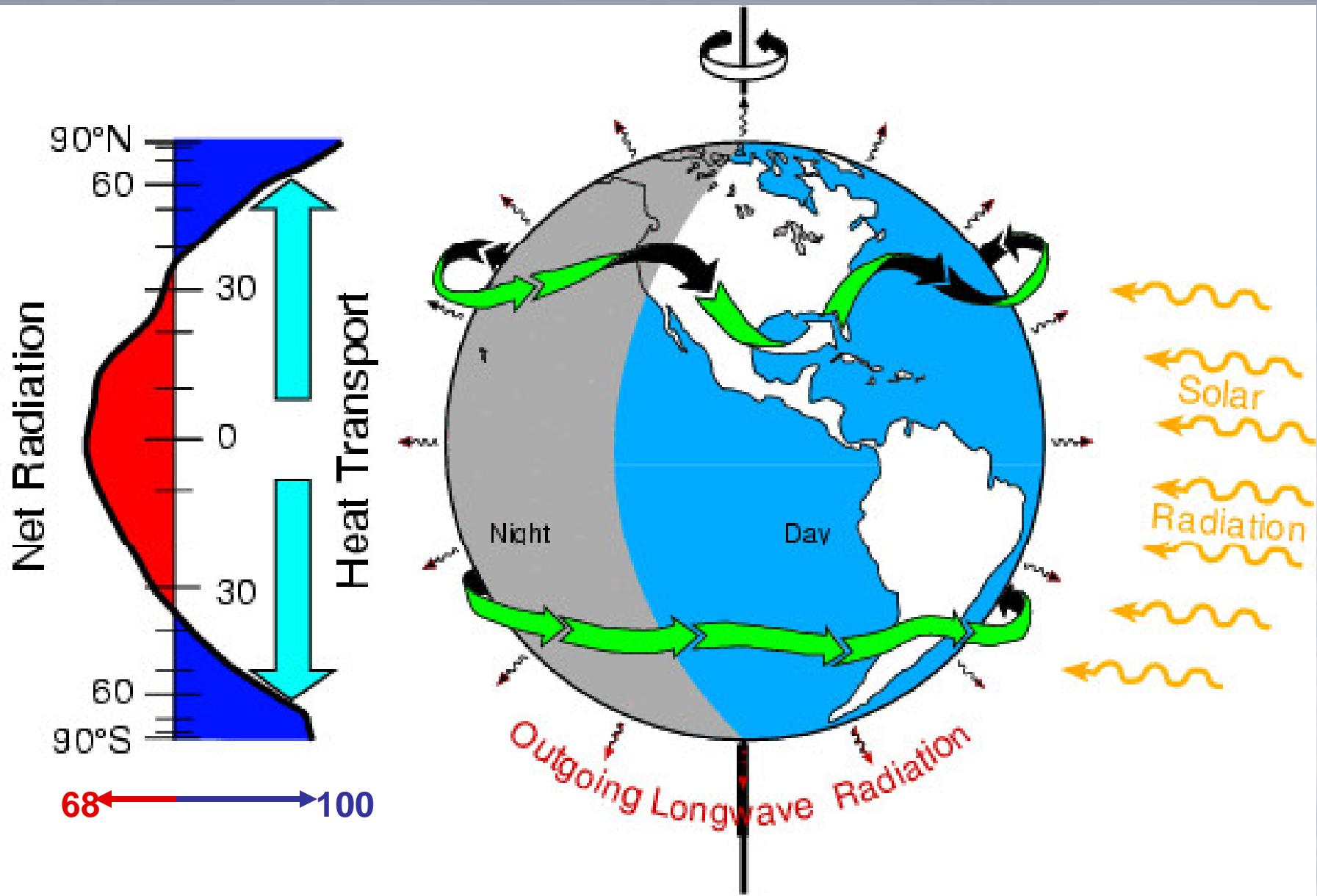


Global Heat Flows



Kiehl and Trenberth 1997





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Observational Dataset

- AIRS (Atmospheric Infrared Sounder)

Lifetime: May 2002 – Present

Spectral coverage: $650\text{-}2665\text{ cm}^{-1}$

Spectral resolution (FWHM): $\nu/1200\text{ cm}^{-1}$

Field of view: 13.5 km

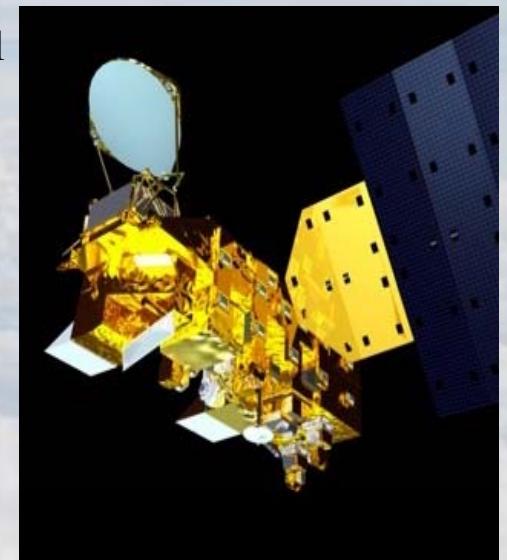
(~9 times of the FOVs per AMSU-A FOV)

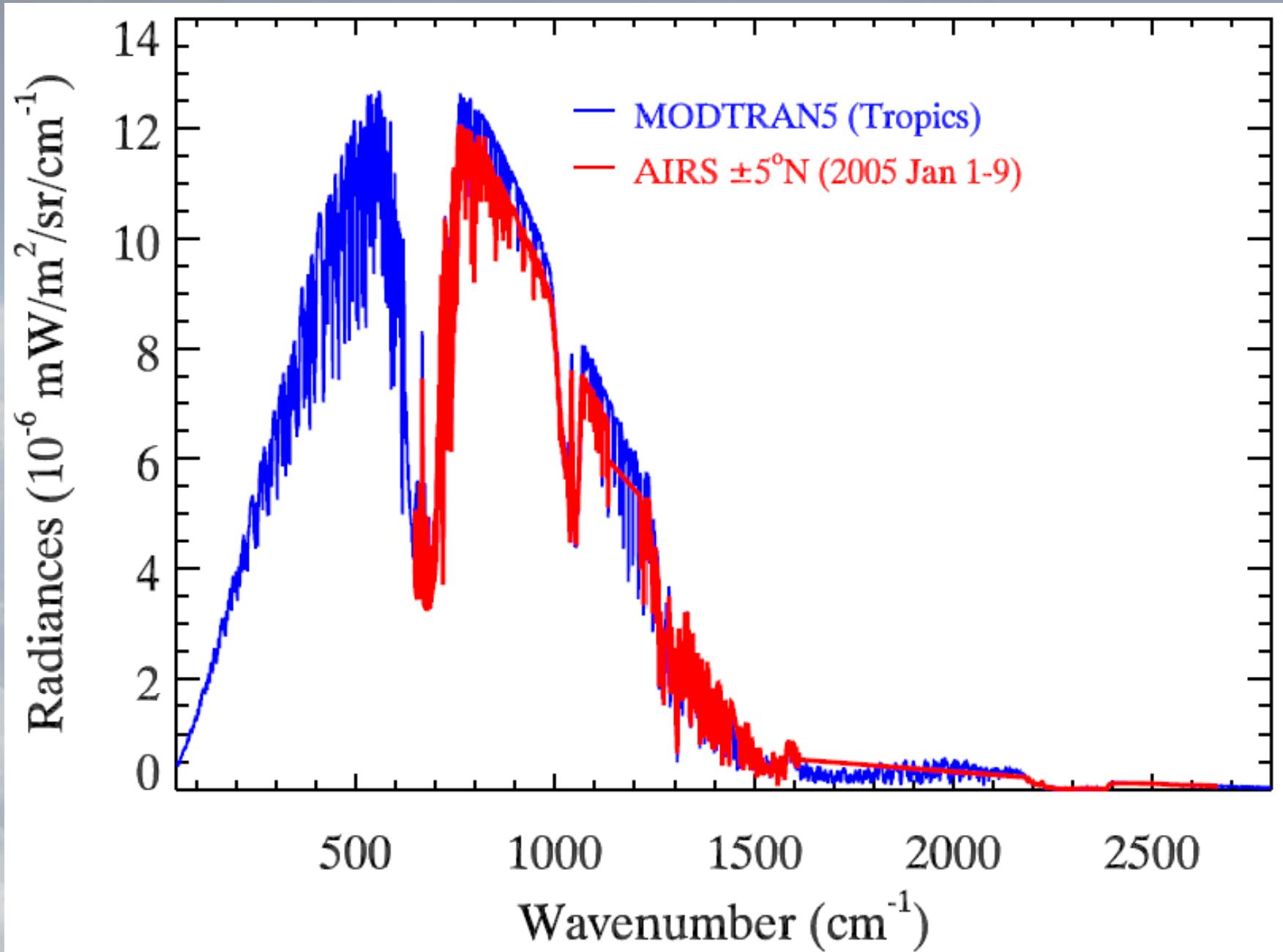
SNR: 600-4000

Calibration bias: 0.07-0.5 K in BT

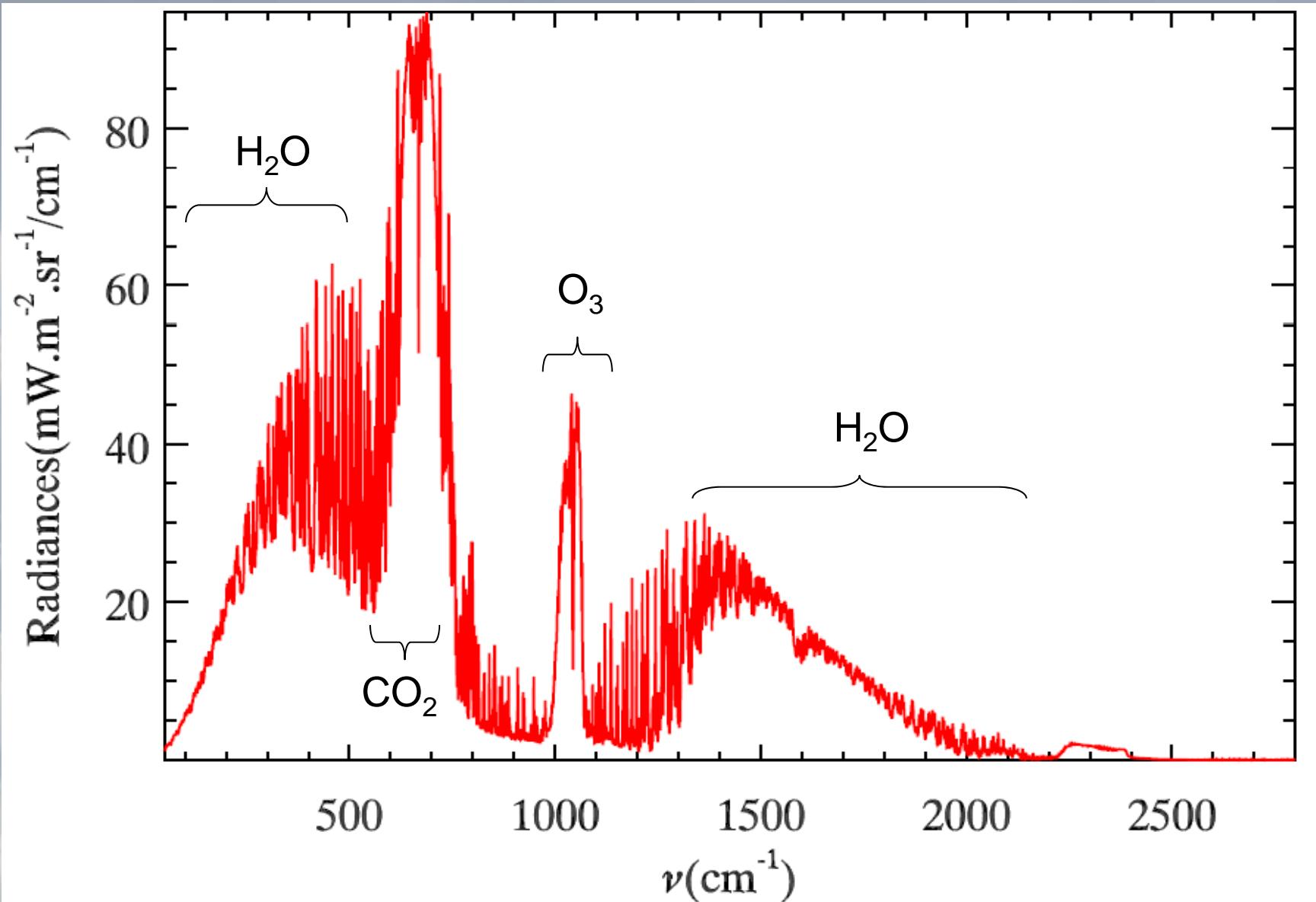
Orbit: Sun-synchronous, 1:30 PM sampling

Swath Width: 1650 km, twice daily swath



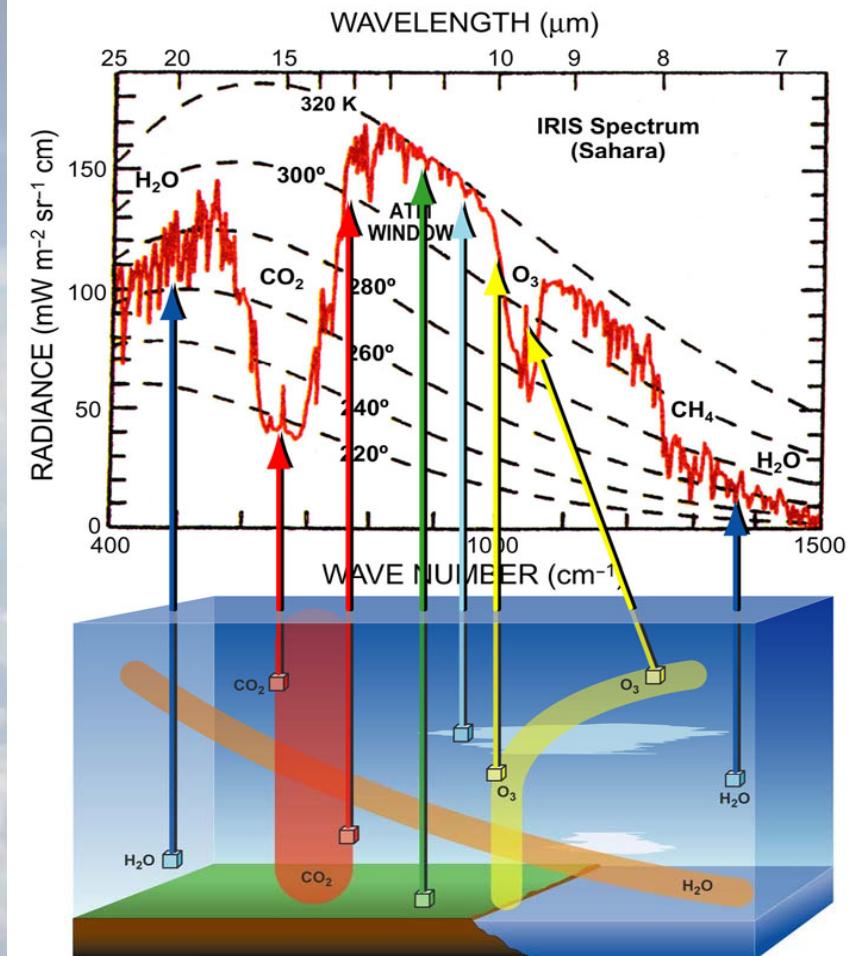


Absorption in Earth's Atmosphere

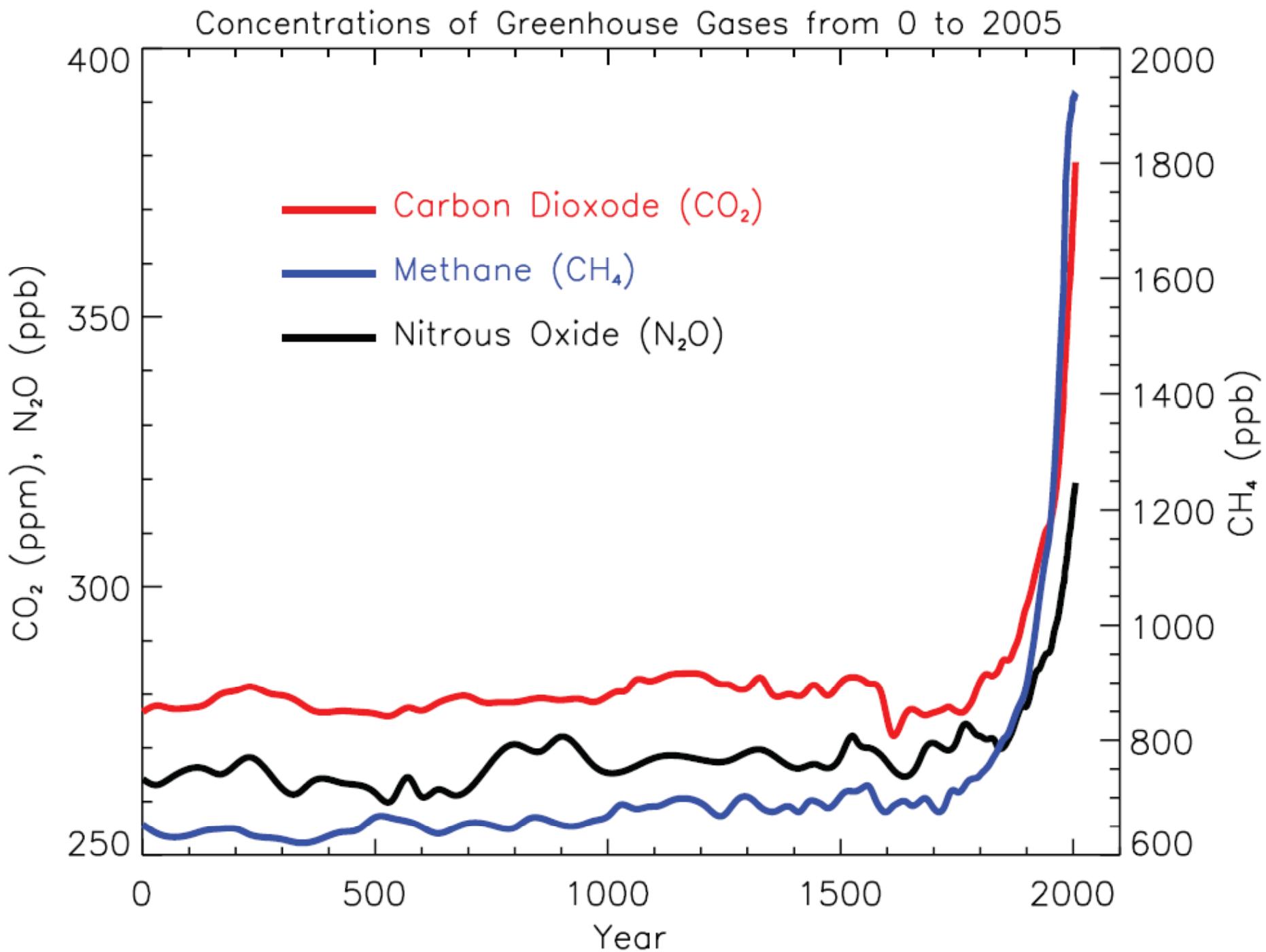


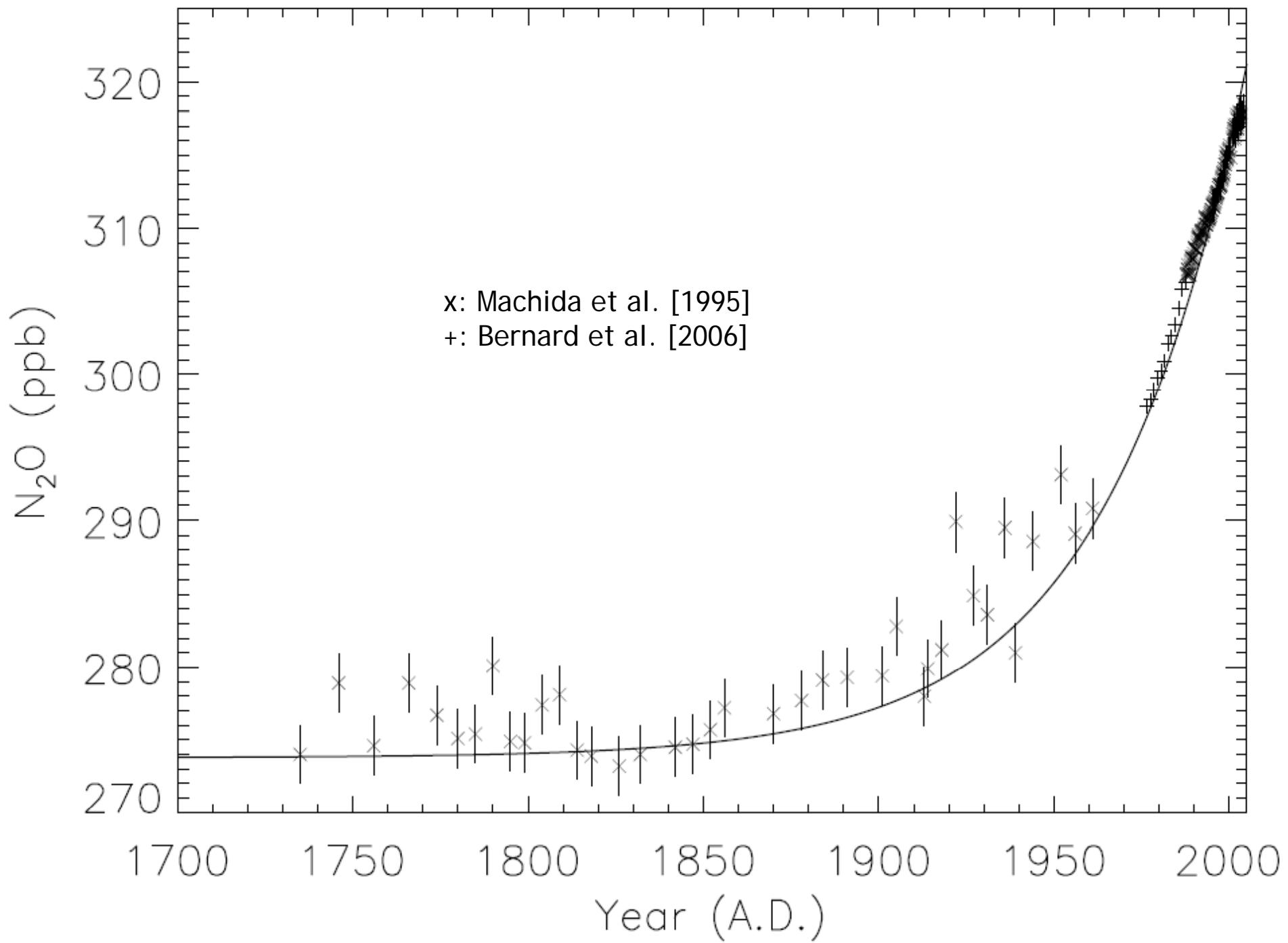
Greenhouse gases and radiance

- Outgoing thermal infrared spectra ($\text{W/m}^2/\text{cm}^{-1}/\text{sr}$) $I(\nu)$
 - Integrand of OLR, containing many pieces of useful information
 - Fundamental to climate diagnostics



By courtesy of John Dykema

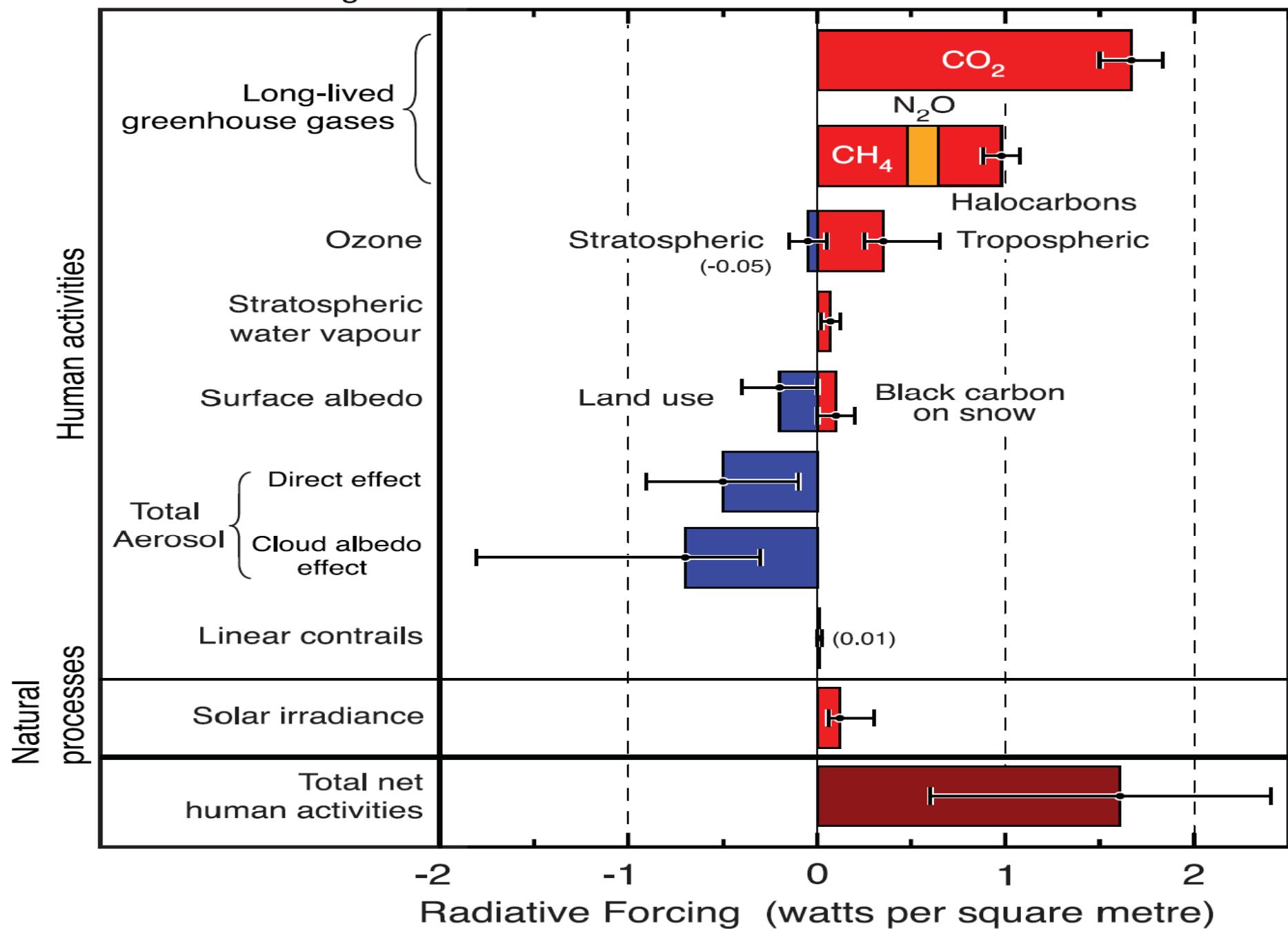




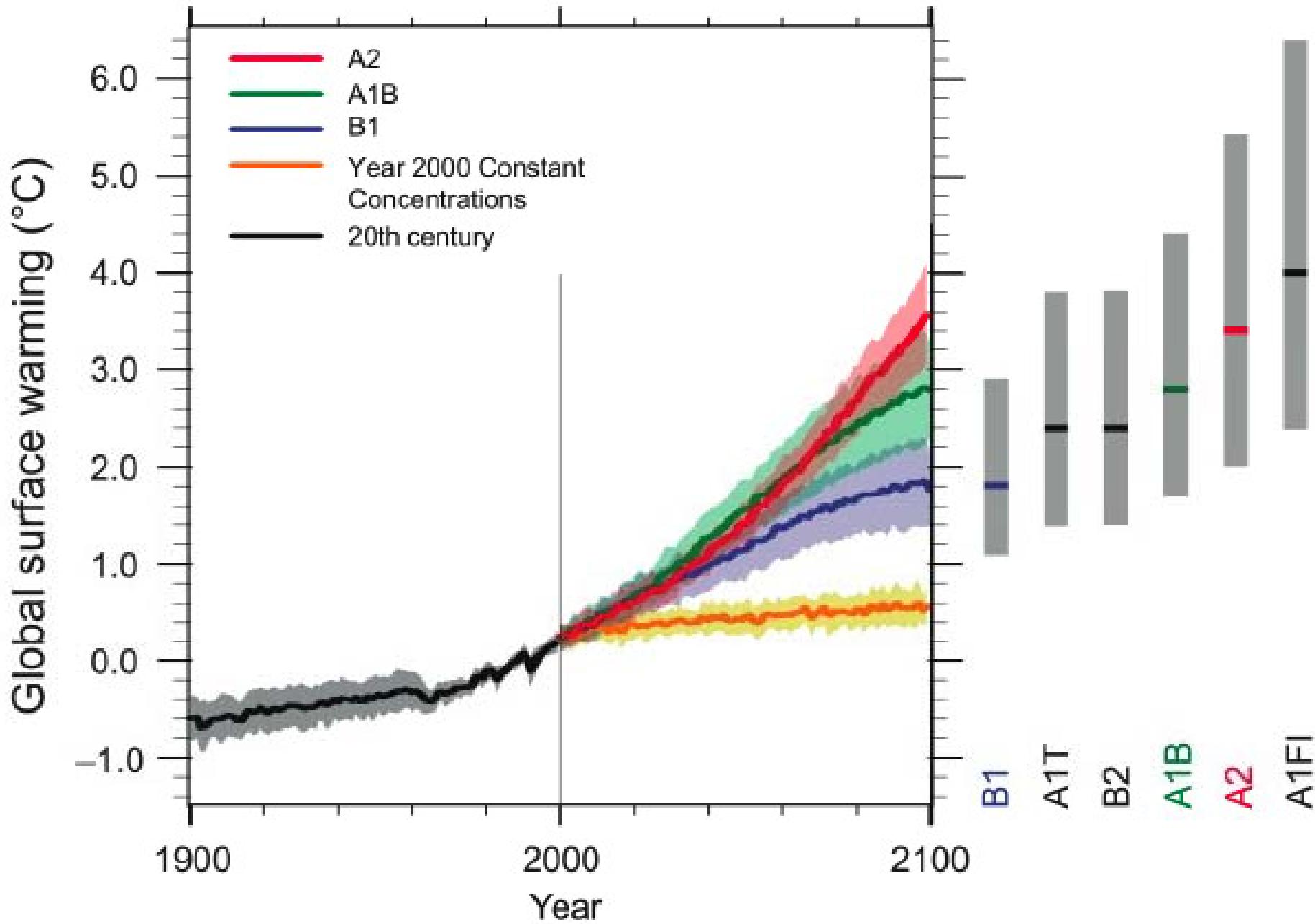
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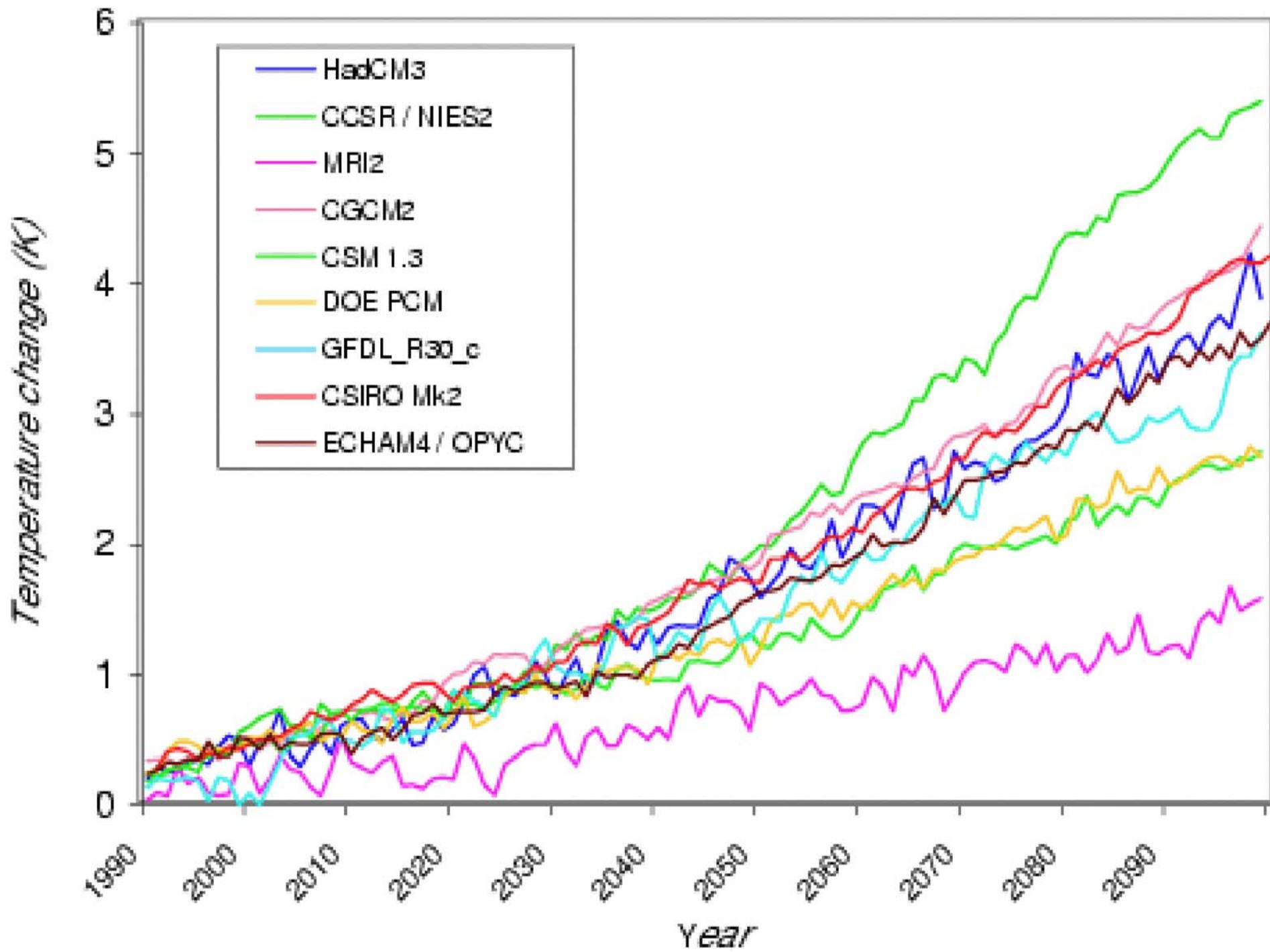
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Radiative Forcing Terms



Multi-model Averages and Assessed Ranges for Surface Warming

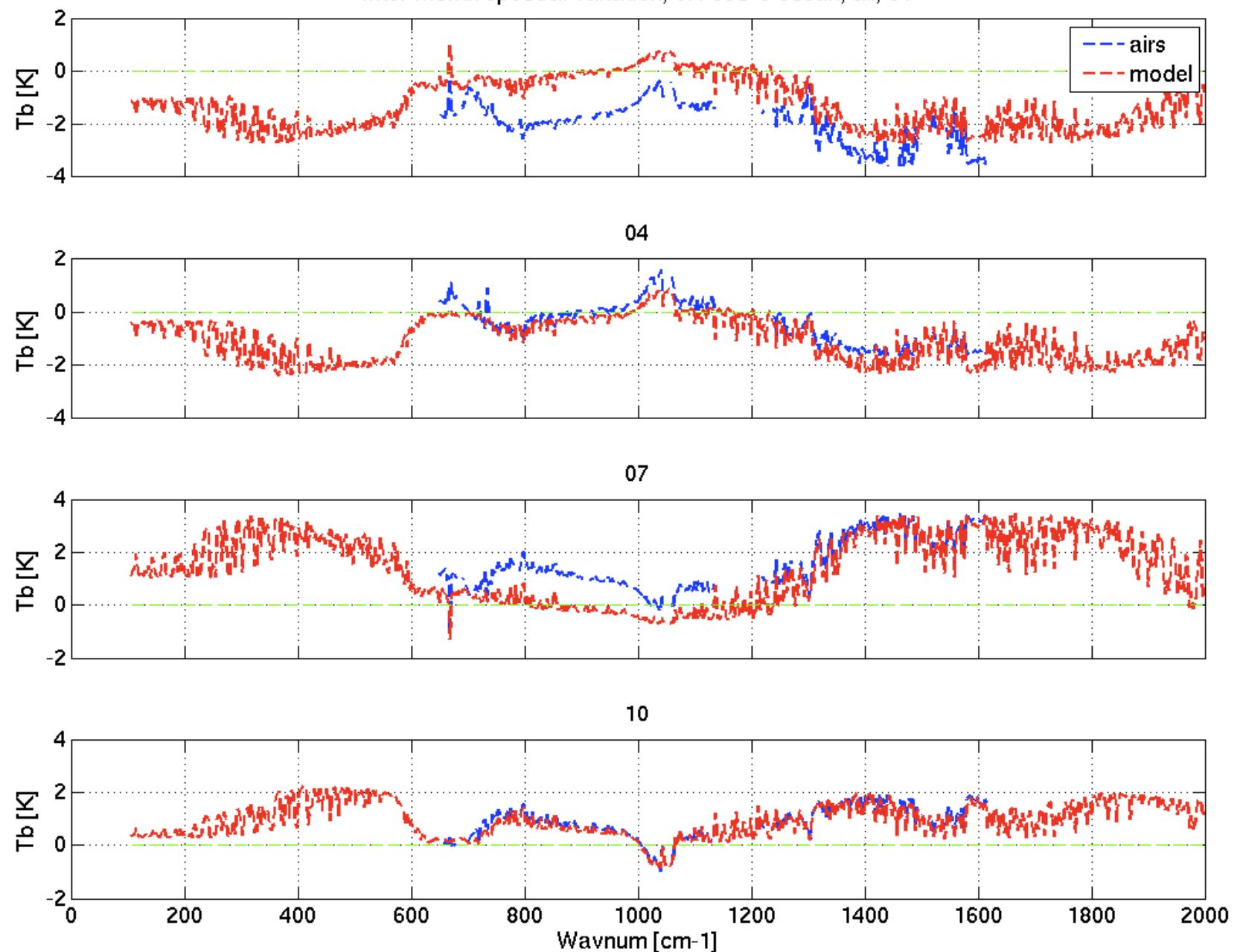




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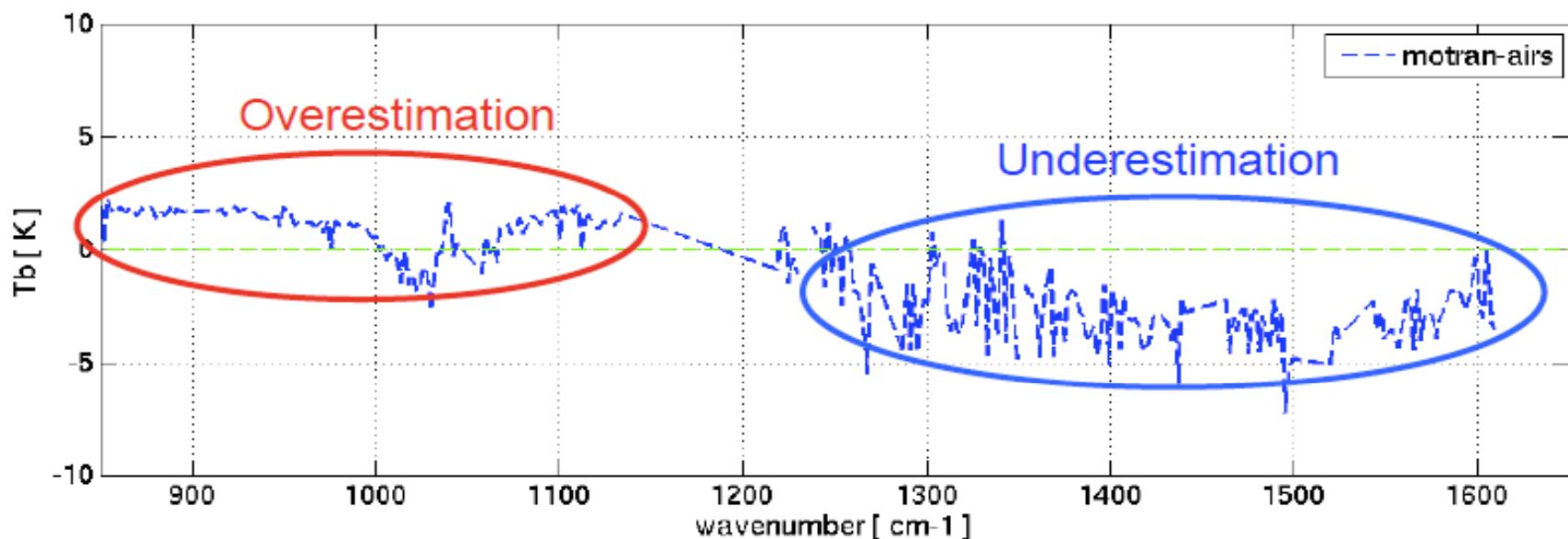
Inter-month spectral variation, 37: 30S-0 ocean, all, 01



Offsetting errors

The seemingly good agreement (-1.1 W m⁻² difference) of the total-sky OLR broadband flux may be fortuitous and arise due to a cancellation

Unit: W m ⁻²	OLR		Window band	
	Total sky	Clear sky	Total sky	Clear sky
CERES	241.73	275.87	66.94	83.28
AM2	240.63	263.43	73.99	87.56
AM2-CERES	-1.10	-12.44	7.05	4.28



Observational Dataset

- IRIS (Infrared Interferometer Spectrometer)

Lifetime: Apr 70 – Jan 71

Spectral coverage: $400\text{-}1600\text{ cm}^{-1}$

(in this study, $400\text{-}1400\text{ cm}^{-1}$)

Spectral resolution (FWHM): 2.8 cm^{-1}

Field of view: 95km in diameter

(~6.25 times of the FOV of ERBE, 15% of an AM2 grid box)

SNR: 20~100

Calibration bias: 0.25-0.75K in BT

Orbit: Sun-synchronous, 10:30AM/PM sampling

700,000 good spectra over 10 months

Observational Dataset

- AIRS (Atmospheric Infrared Sounder)

Lifetime: May 2002 – Present

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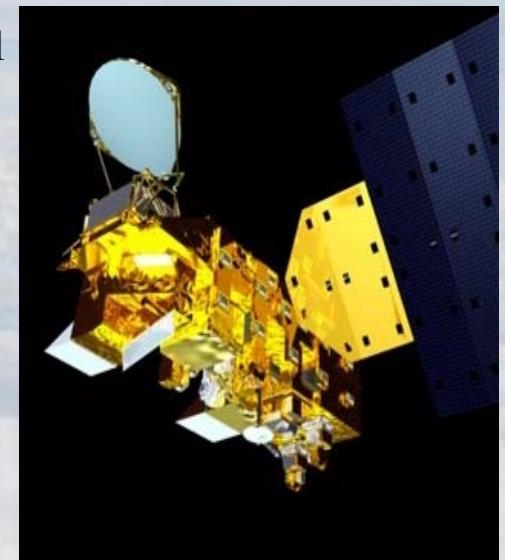
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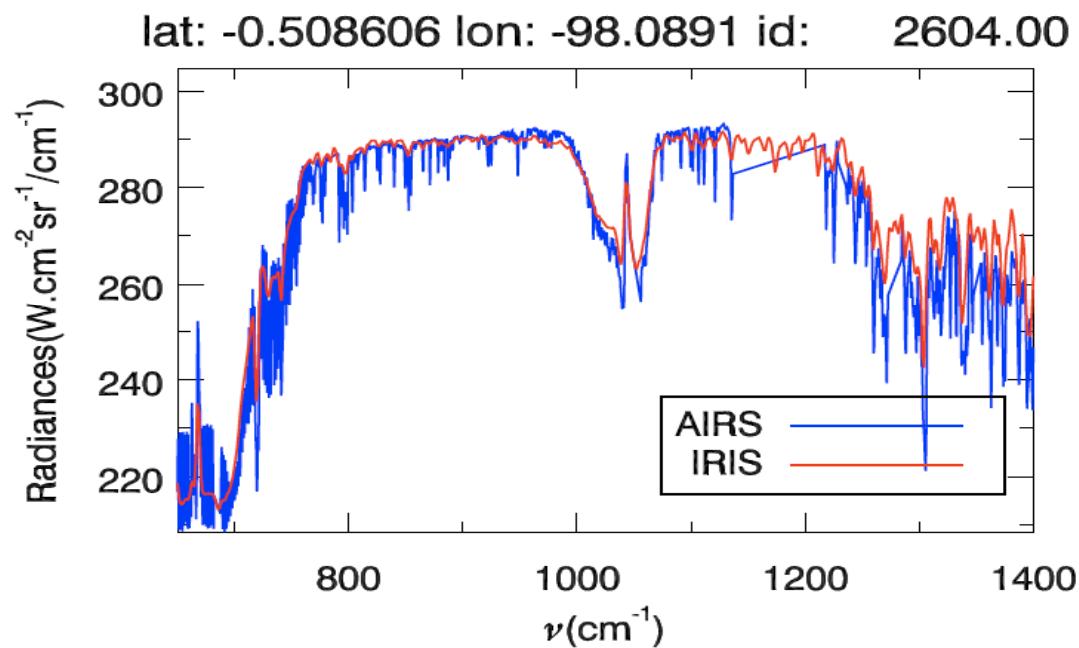
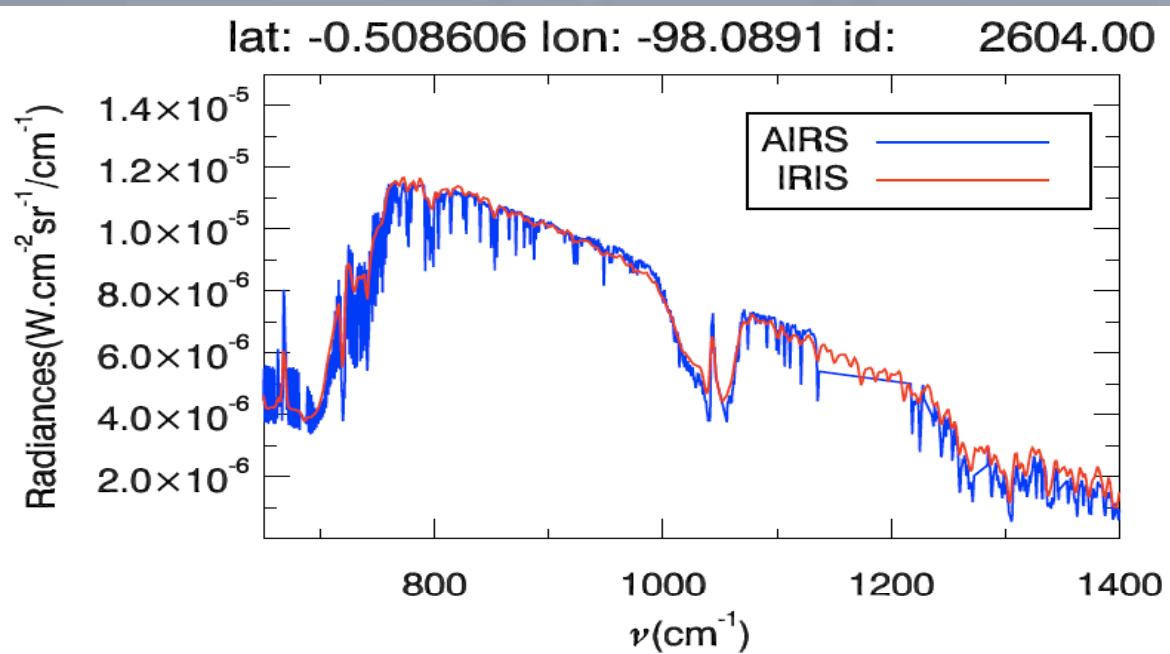
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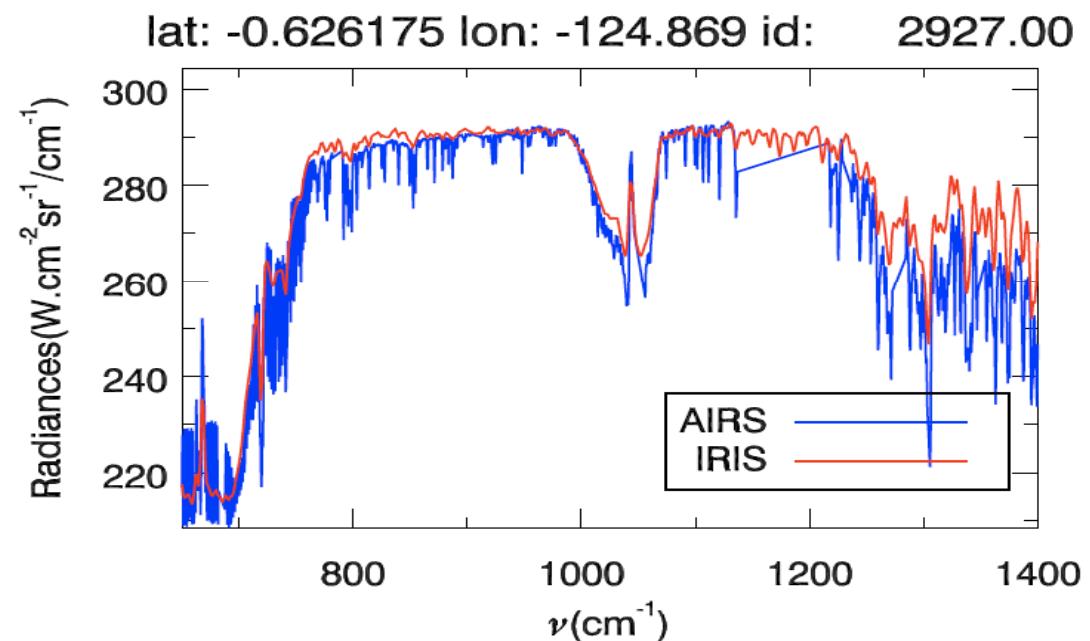
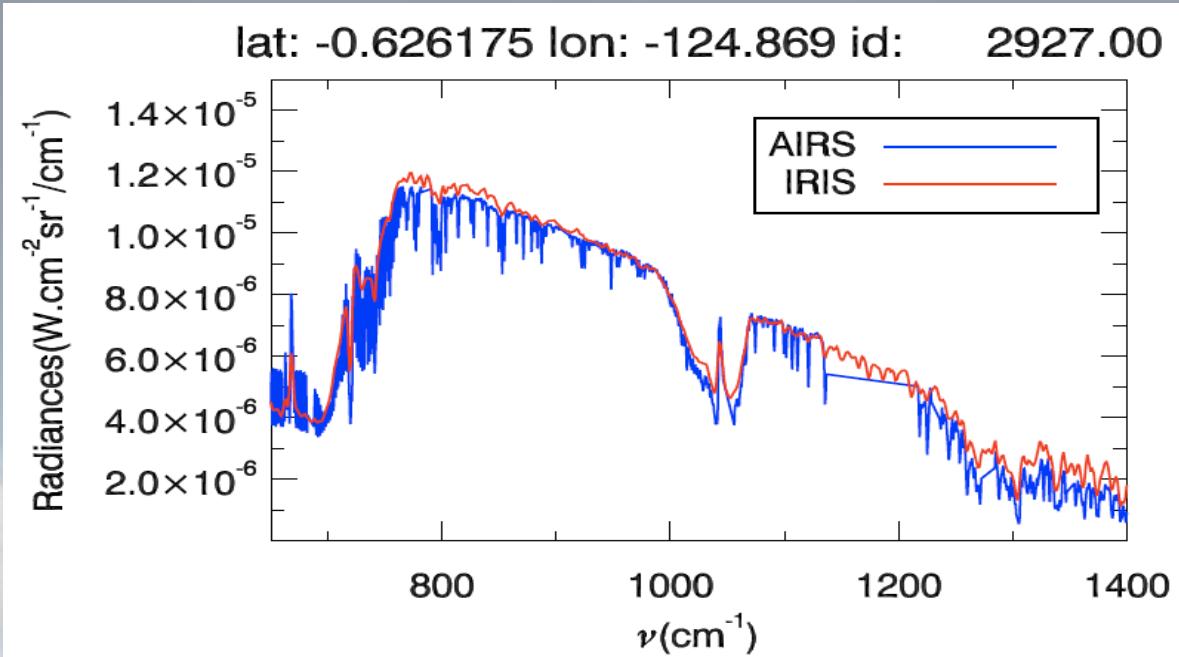
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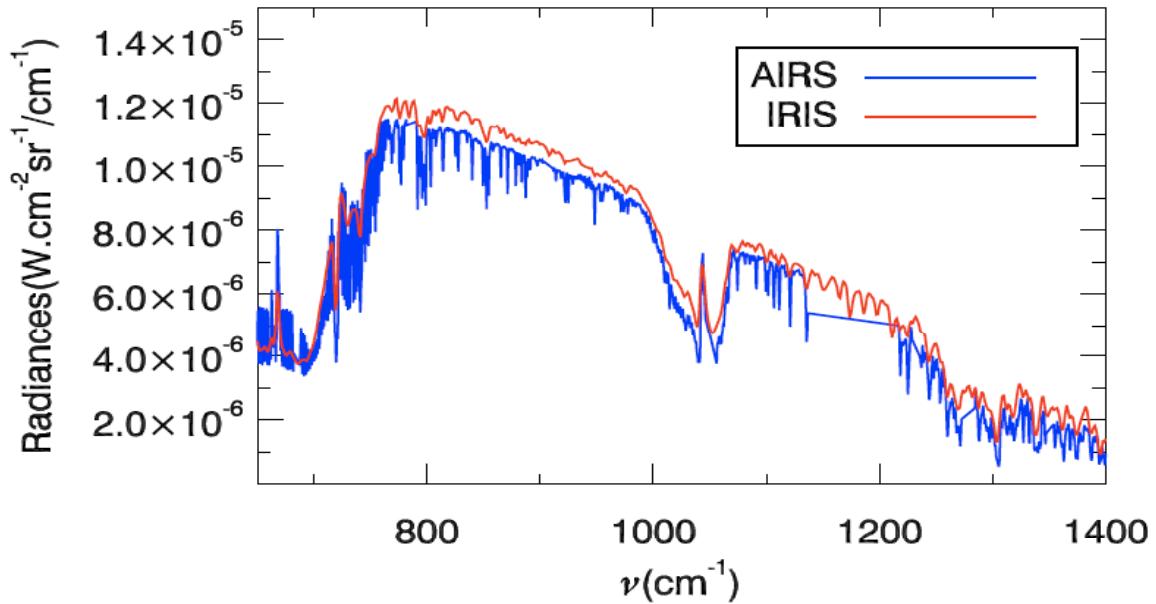
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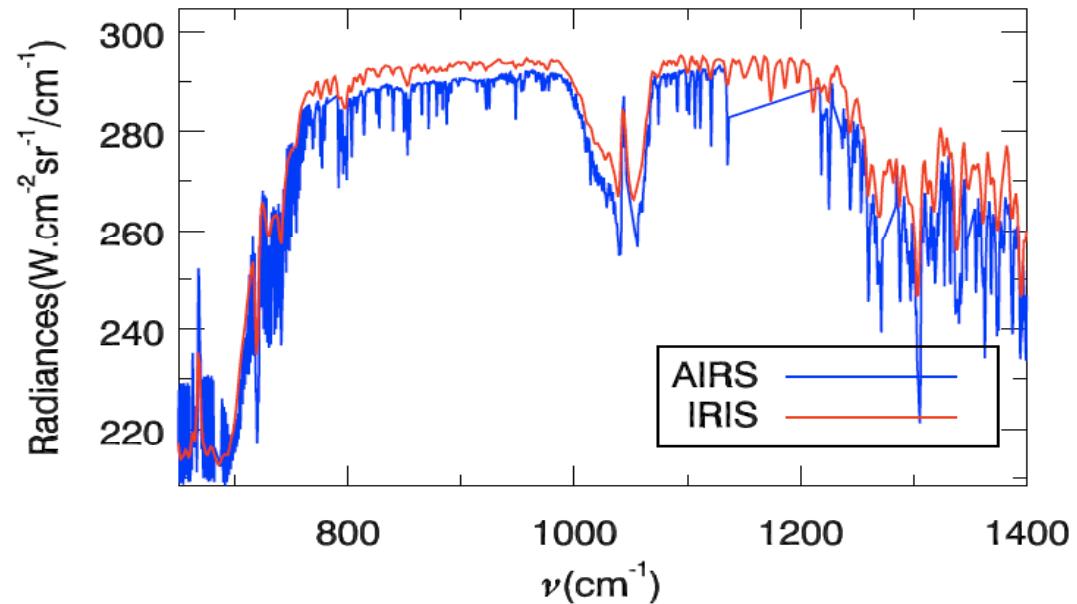




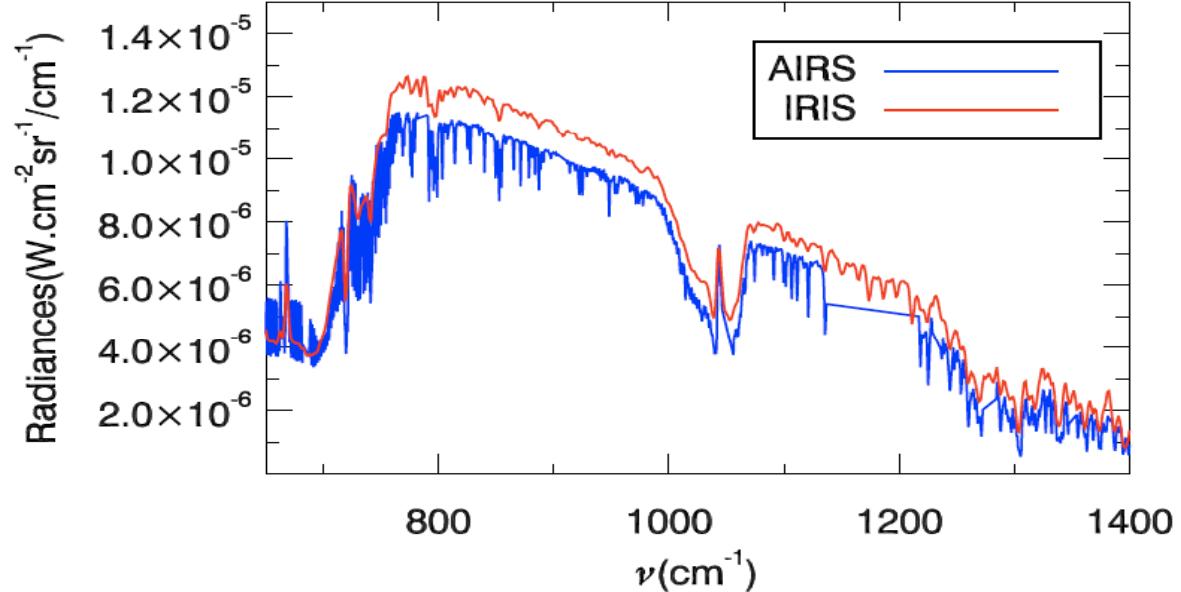
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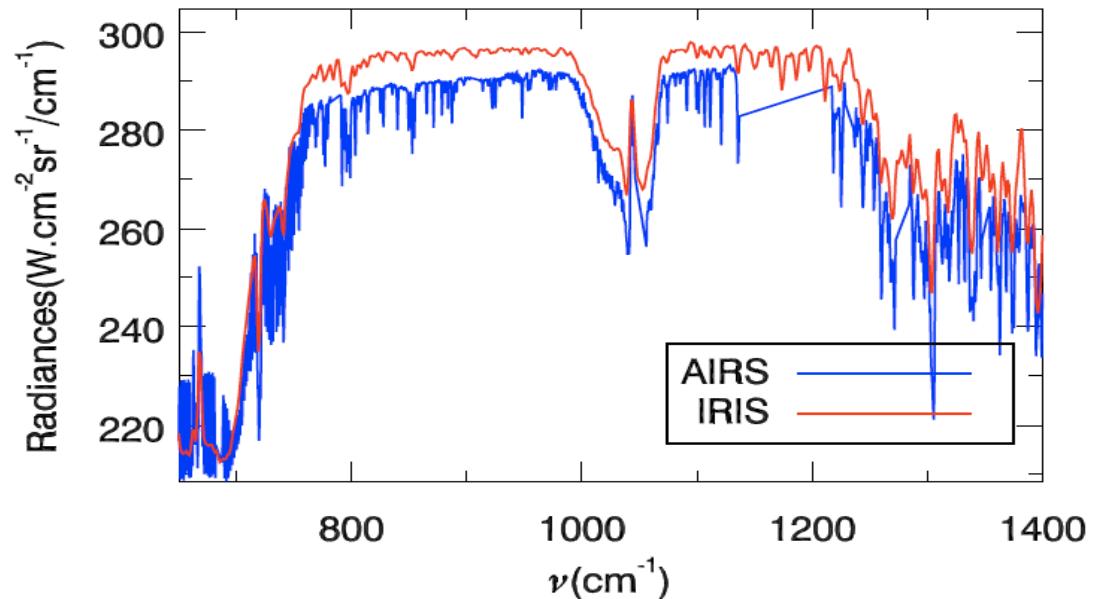
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lat: -0.861145 lon: -178.429 id: 3503.00



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Conclusions

Radiation plays a fundamental role as the source of energy

Spectral radiance provides detailed picture of energy exchange

Radiation as diagnostic of climate change (potential)

Thanks

- ❖ Yung's Group at Caltech
 - ❖ Trenberth (NCAR)
 - ❖ Prof. Xianglei Huang (U Michigan)
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- ❖ *Ref: Goody and Yung (1989)*
 - ❖ *Huang and Yung (2005) JGR 110*