

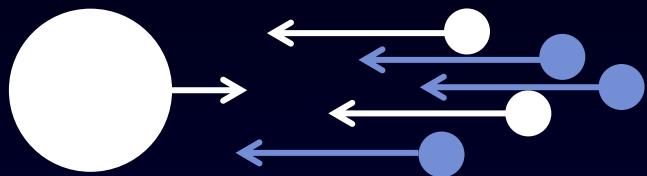
Physical properties of cometary analogue materials

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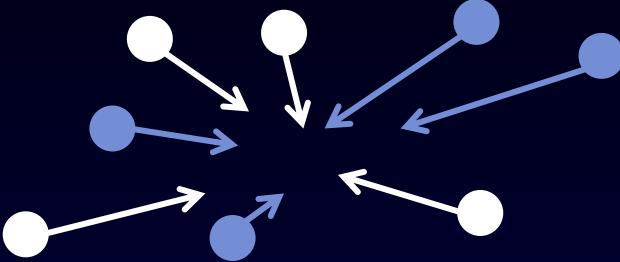
Mass transfer (MT)



(Windmark et al., 2012; Garaud et al., 2013)

- > Collision of the formed ~ cm-sized aggregates with „lucky winners”
- > High impact velocities
- > Fragmentation of impacting aggregates
- > Mass transfer leads to growth

Gravitational instability (GI)

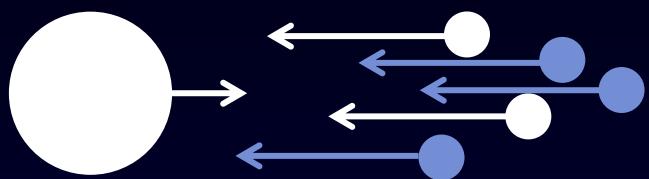


(Johansen et al., 2007)

- > Concentration of an ensemble of ~ cm-sized aggregates by, i.e., streaming instability*
- *: (Youdin and Godman, 2005)
- > Ensemble becomes gravitational unstable -> collapse
- > Relatively low collision velocities**
-> ~ 0.01 m/s - 0.1 m/s**
- > Aggregates are not destroyed**
**: (Wahlberg and Johansen, 2014)

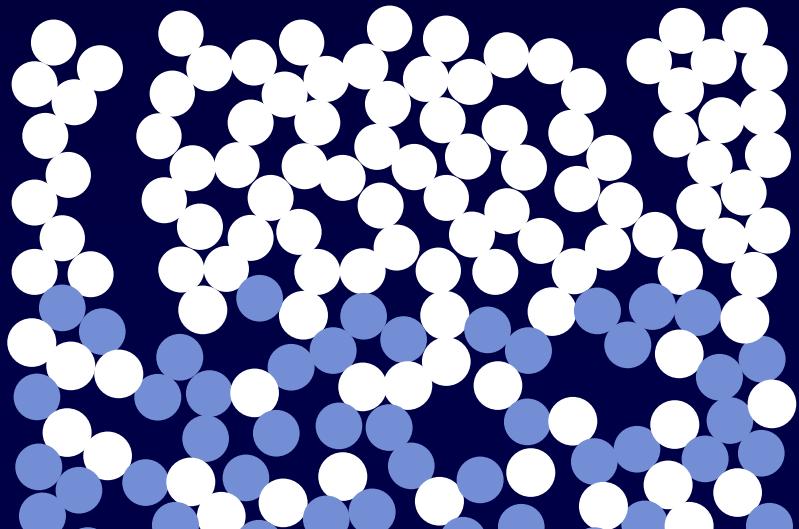
Formation scenarios

Mass transfer (MT)

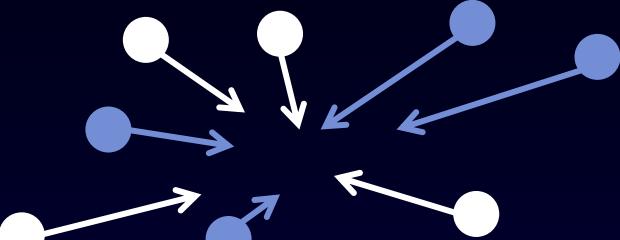


Cometary surface:

- > Relatively compact
- > Volume Filling Factor: ~ 0.4
- > μm -sized monomers

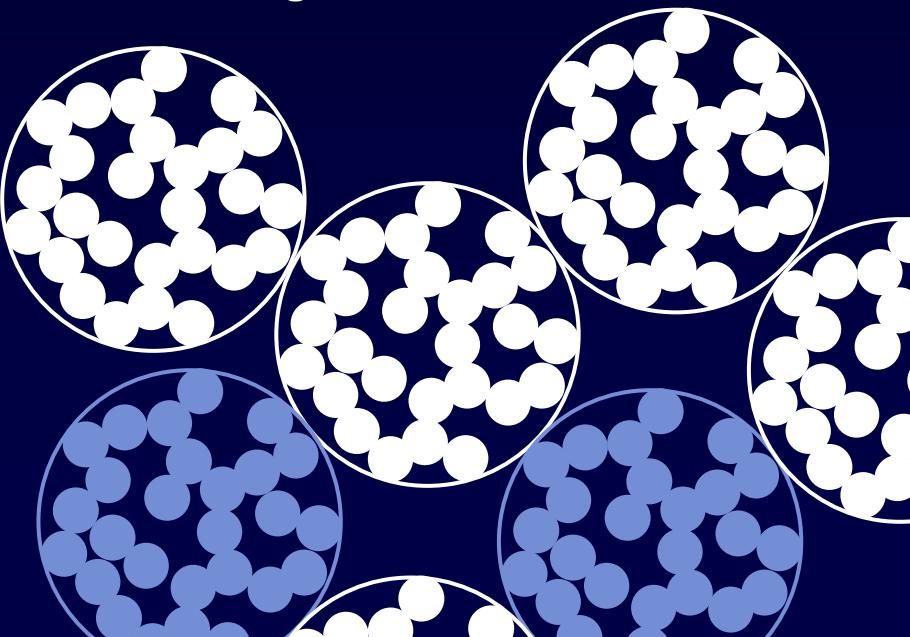


Gravitational instability (GI)



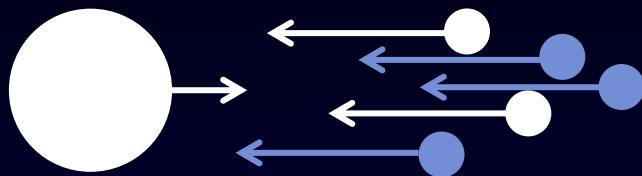
Cometary surface:

- > Loose arrangement of aggregates
- > Vol. filling factor: $\sim 0.35 \times 0.6 \approx 0.2$

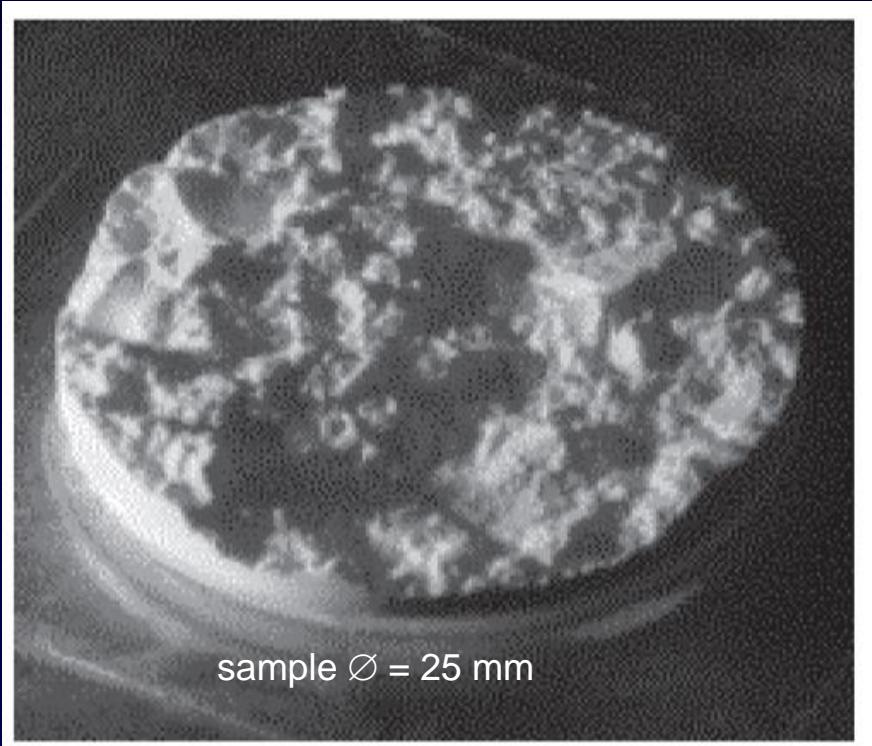


Formation scenarios

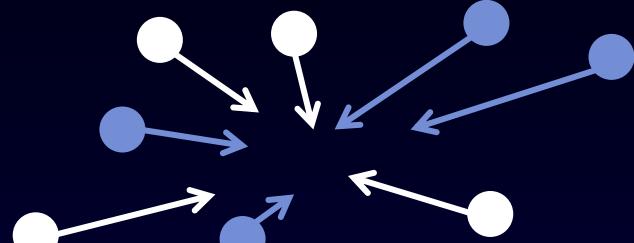
Mass transfer (MT)



Cometary surface:



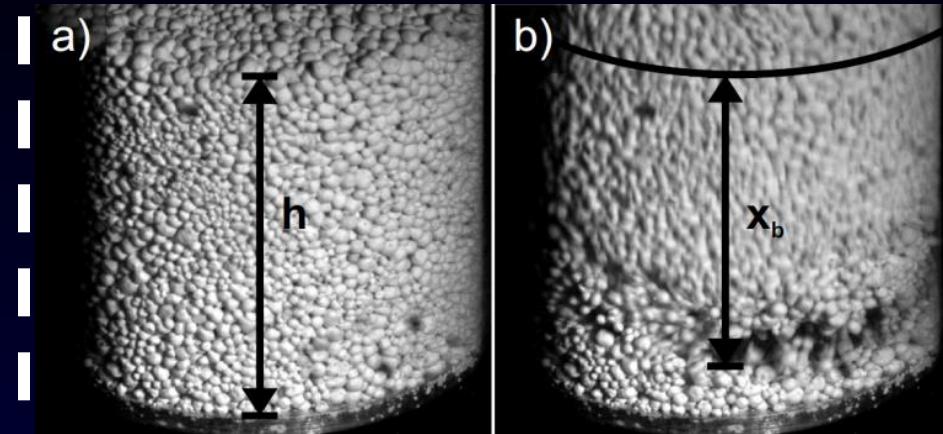
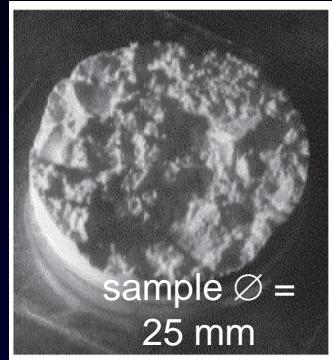
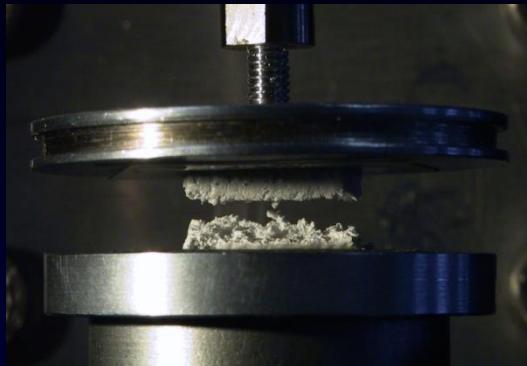
Gravitational instability (GI)



Cometary surface:



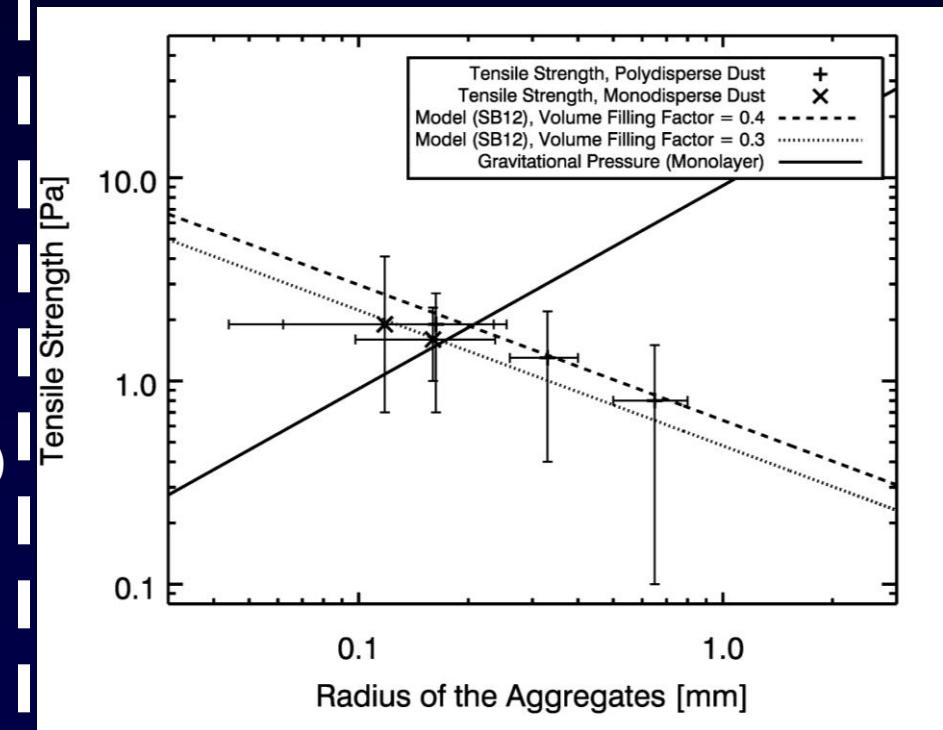
Tensile Strength



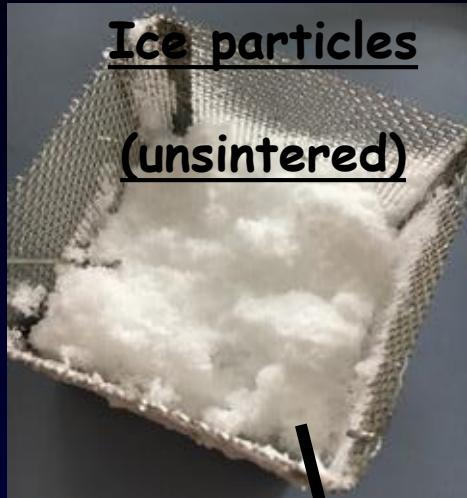
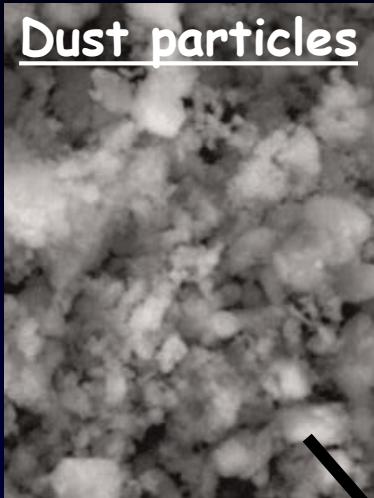
$$T(\phi) = 2400 \text{ Pa} (\phi = 0.41)$$

$$T(\phi) = 10^{2.8 + 1.48\phi} \text{ (Pa)}$$

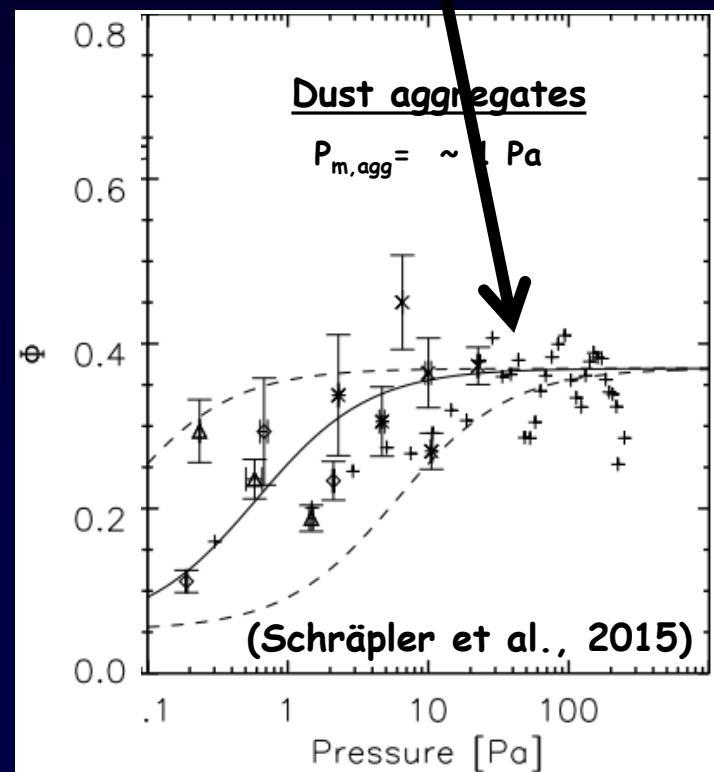
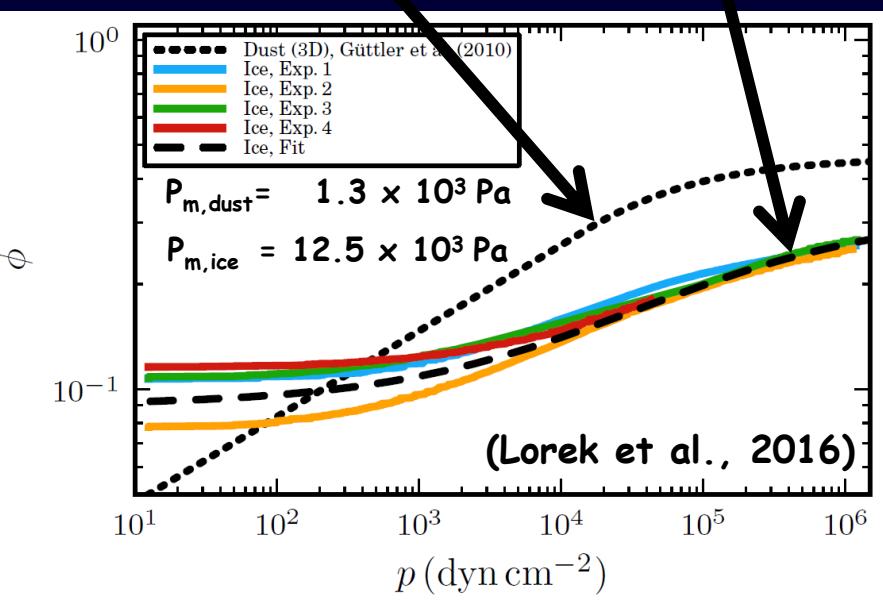
(Güttler et al., 2009)



(Data from Blum et al., 2014; Brisset, 2014)



Granular H_2O ice



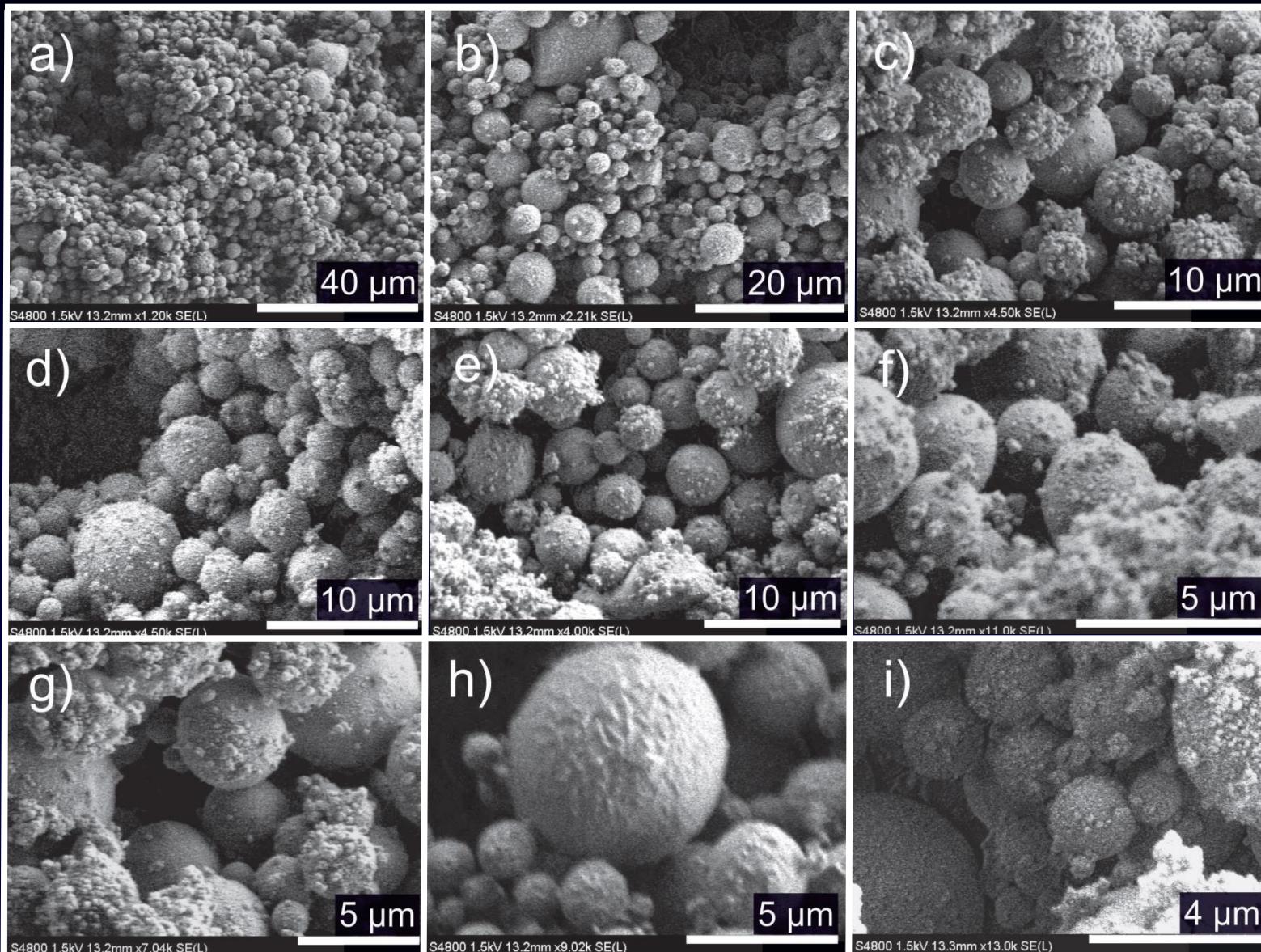
Overview (dust)

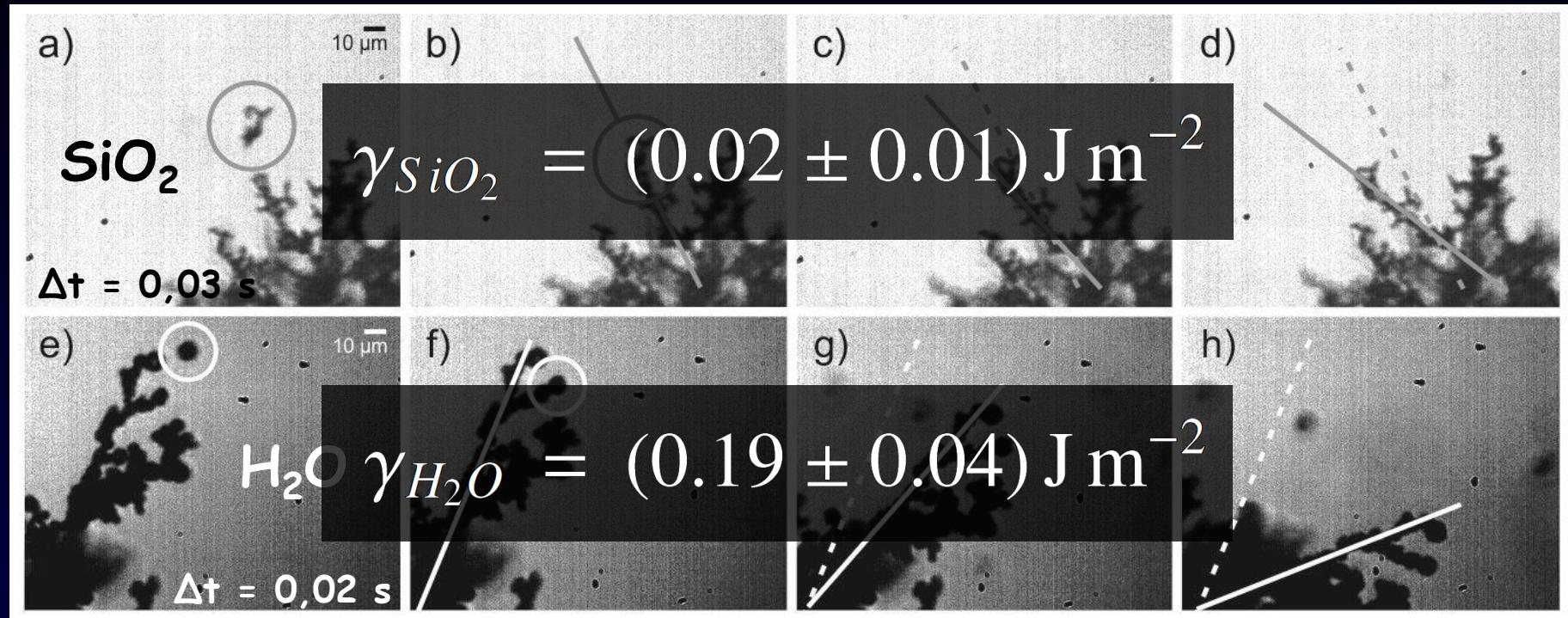
	Mass transfer (MT)	Gravitational instability (GI)
Volume filling factor	~ 0.4	~ $0.35 \times 0.6 \approx 0.2$
Tensile strength (surface)	~ 1000 Pa	~ 1 Pa
Compression curve (p_m @ $\Phi \approx 0.25$)	~ 1300 Pa (up to 10^6 Pa) *no ice (MUPUS: > MPa)	~ 1 Pa (up to 10^3 Pa) (MUPUS: > MPa)
Gas permeability	$\sim 1 \times 10^{-9} \text{ m}^4 \text{s}^{-1}$	$\sim 1 \times 10^{-6} \text{ m}^4 \text{s}^{-1}$
Thermal conductivity	$10^{-2} - 10^{-1} \text{ W m}^{-1} \text{ K}^{-1}$ (conduction)	$10^{-3} - 1 \text{ W m}^{-1} \text{ K}^{-1}$ (conduction/radiation)

And KOSI...?

- The measured heat conductivity of the sample material ($1 \times 10^{-1} \text{ W m}^{-1} \text{ K}^{-1}$ to $6 \times 10^{-1} \text{ W m}^{-1} \text{ K}^{-1}$; Spohn et al., 1989) is lower than the bulk heat conductivity of water ice. However, the heat conductivity of the KOSI samples is much higher than the heat conductivity estimated for the surface material of Comet 9P/Tempel (heat conductivity: $\lambda \approx 1 \times 10^{-3} \text{ W m}^{-1} \text{ K}^{-1}$; Groussin et al., 2007).
- Irradiated ice-dust mixtures form a non-volatile dust layer, which decreases the gas production rate of the sample (see Fig. 1.12; Lämmenzahl et al., 1995).
- The compressive strength of the sample material is very low and increases during the experiment from $\sim 1 \times 10^5 \text{ N m}^{-2}$ to $\sim 5 \times 10^6 \text{ N m}^{-2}$ [Colangeli et al., 2004]. The increase of the compressive strength is probably caused by mass transport inside the sample due to the outgassing of the icy materials, or by sintering¹⁸ of the sample material.
- The activity of the icy constituents causes the emission of particles (Lämmenzahl et al., 1995).

Water ice





(Gundlach et al., 2011b)

$$F_{Roll} = m g_0 \frac{a_{cm}}{r} \quad \leftrightarrow \quad F_{Roll} = 6 \pi \gamma \xi$$

(Blum and Wurm, 2000)

(Dominik and Tielens, 1995)

m: Mass of the Aggregate

a_{cm} : Lever Arm (Horizontal Projektion)

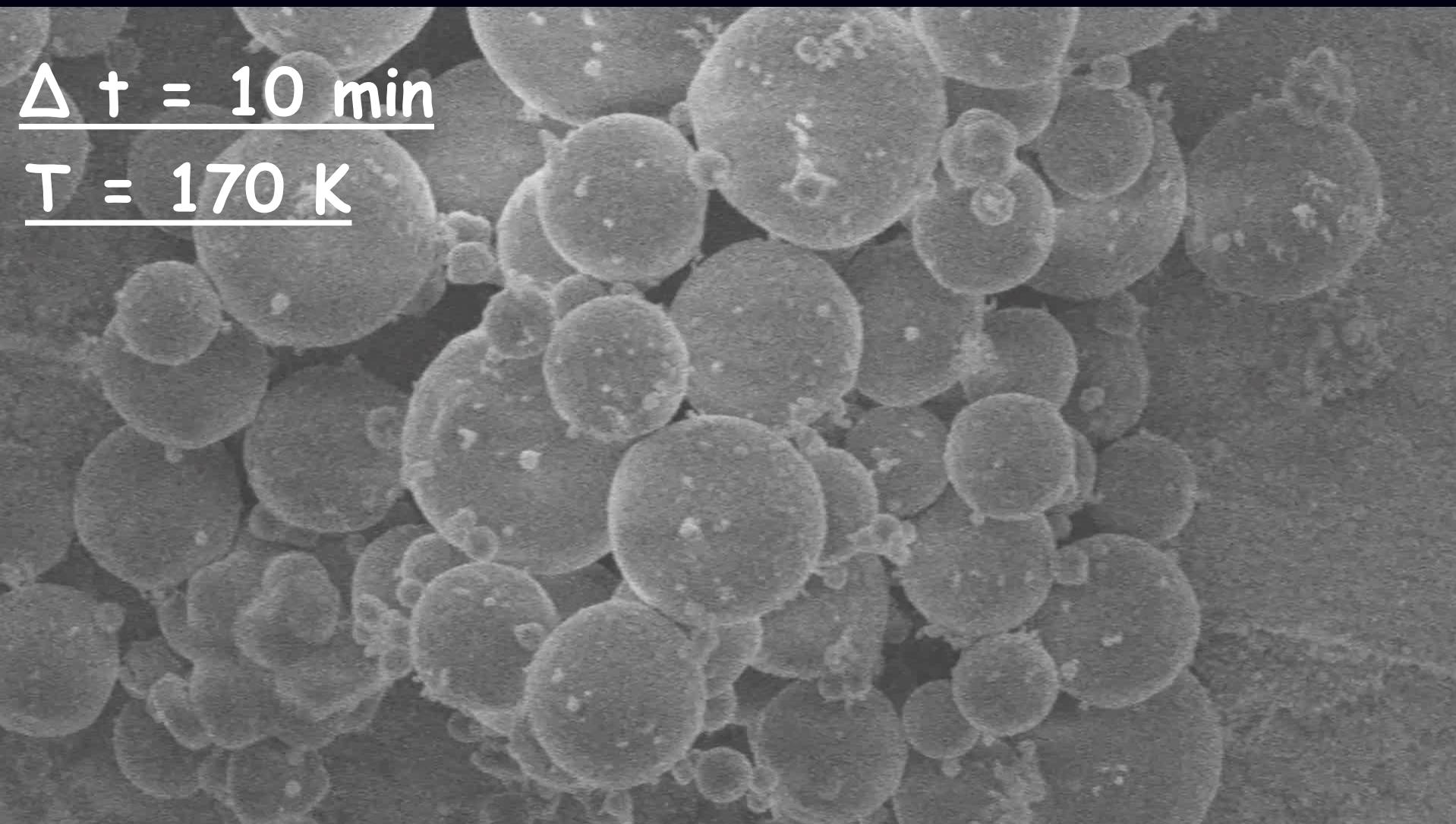
r: Particle Radius

γ : Specific Surface Energy

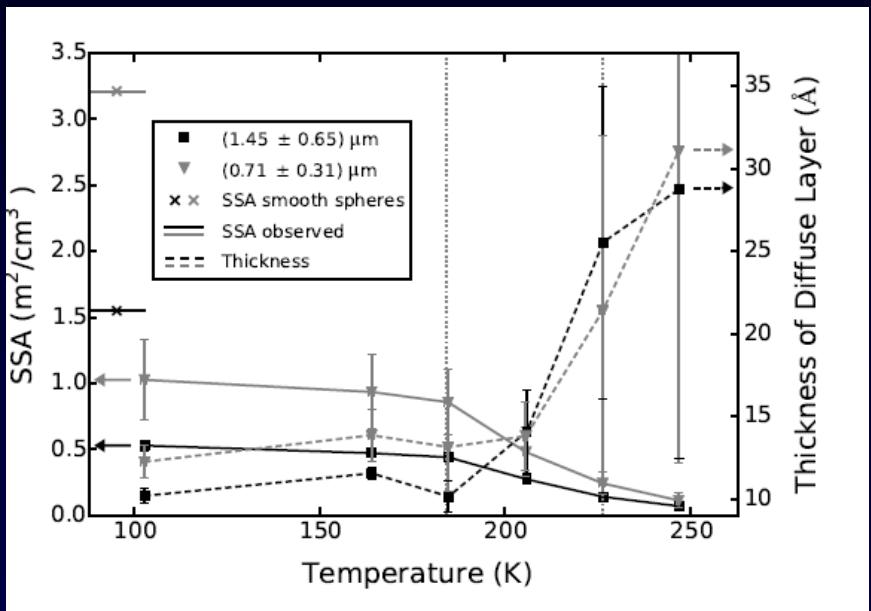
ξ : Critical Rolling Distance

Water ice

$\Delta t = 10 \text{ min}$
 $T = 170 \text{ K}$



Water ice (@ 200 K)



(Gärtner et al., submitted)



Formation of a solidified layer on comets?

